# Accurate Establishment of Error Models for the Satellite Gravity Gradiometry Recovery and Requirements Analysis for the Future GOCE Follow-On Mission 

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## Abstract

Firstly, the new single and combined error models applied to estimate the cumulative geoid height error are efficiently produced by the dominating error sources consisting of the gravity gradient of the satel-lite-equipped gradiometer and the orbital position of the space-borne GPS/GLONASS receiver using the power spectral principle. At degree 250 , the cumulative geoid height error is $1.769 \times 10^{-1} \mathrm{~m}$ based on the new combined error model, which preferably accords with a recovery accuracy of $1.760 \times 10^{-1} \mathrm{~m}$ from the GOCE-only Earth gravity field model GO_CONS_GCF_2_TIM_R2 released in Germany. Therefore, the new combined error model of the cumulative geoid height is correct and reliable in this study. Secondly, the requirements analysis for the future GOCE Follow-On satellite system is carried out in respect of the preferred design of the matching measurement accuracy of key payloads

[^0]comprising the gravity gradient and orbital position and the optimal selection of the orbital altitude of the satellite. We recommend the gravity gradient with an accuracy of $10^{-13}-10^{-15} / \mathrm{s}^{2}$, the orbital position with a precision of $1-0.1 \mathrm{~cm}$ and the orbital altitude of $200-250 \mathrm{~km}$ in the future GOCE Follow-On mission.

Key words: GOCE Follow-On, single and combined error models, requirements analysis, power spectral principle, satellite gravity gradiometry recovery.

## 1. INTRODUCTION

The successful implementation of the satellite gravity gradiometry (SGG) mission is another breakthrough innovation in the interdisciplinary scientific fields, e.g., space geodesy, geophysics, etc., after the Global Positioning System (GPS) is constructed by the U.S. Department of Defense (DOD) (Hsu 2001). The European Space Agency (ESA) independently developed and successfully launched the Gravity Field and Steady-State Ocean Circulation Explorer (GOCE) satellite on 17 March 2009. As shown in Table 1, the GOCE satellite flies in an almost-circular, sun-synchronous and low-altitude orbit, and adopts a combination of the Satellite-to-Satellite Tracking in the High-Low mode (Zheng et al. 2011a, 2012a,b, 2015a) with the Satellite Gravity Gradiometry mode (SST-HL/SGG). Apart from the accurate tracking and positioning of the low-orbiting GOCE satellite by the high-orbit

Table 1
A comparison of the current GOCE and future GOCE Follow-On mission

| Parameters | Gravity satellites |  |  |
| :---: | :---: | :---: | :---: |
|  | GOCE | GOCE Follow-On |  |
| Orbit | Orbital altitude | 250 km | $200-250 \mathrm{~km}$ |
|  | Orbital <br> inclination | $96.5^{\circ}$ | $90-95^{\circ}$ |
|  | Orbital <br> eccentricity | $10^{-3}$ | $10^{-3}-10^{-4}$ |
|  | Tracking model | SST-HL/SGG | SST-HL/SGG |
| Gravity gradient <br> error | $3 \times 10^{-12} / \mathrm{s}^{2}$ <br> $($ Electrostatic suspension <br> gravity gradiometer) | $10^{-13}-10^{-15} / \mathrm{s}^{2}$ <br> $($ Cold-atom interferometric <br> gravity gradiometer) |  |
|  | Orbital position <br> error | $1-0.1 \mathrm{~cm}$ <br> (GPS/GLONASS) | GPS/GLONASS/ <br> Compass/Galileo) |

GPS/GLONASS system (Bobojc and Drozyner 2011), the secondderivatives of the Earth's gravitational potential at the satellite orbit are precisely measured through the space-based electrostatic suspension gravity gradiometer located at the center of mass of the GOCE satellite (Bian and Ji 2006, Eshagh 2010), and the non-conservative forces acting on the GOCE system are compensated in real time by the satellite-borne Drag-Free Control System (DFCS).

At present, the recovery methods for the satellite gravity gradiometry mainly include the space-wise method (Reguzzoni and Tselfes 2009, Migliaccio et al. 2010, 2011; Pertusini et al. 2010, Reguzzoni et al. 2010, Sanso et al. 2011, Liu et al. 2012), the time-wise method (Milani et al. 2005, Pail et al. 2010, 2011a; Eshagh 2011, Goiginger et al. 2011), the space-timewise method (Zheng et al. 2011b), and the direct method (Bruinsma et al. 2010, Pail et al. 2011b), and so on. On 10 November 2013, the GOCE satellite has ended its extended mission to map the Earth's gravitational field. As displayed in Table 1, many international research organizations are actively carrying through the intensive and extensive investigations concerning the requirements analysis and payload development for the future GOCE Fol-low-On satellite mission (Bender et al. 2003, Rummel 2003, Zheng et al. 2012c, 2013, 2015b) for the sake of meeting the urgent requirements for the related interdisciplines on further improving the accuracy of the Earth's gravitational field determination. In the interest of efficiently getting rid of the existing shortcomings containing the complex computational process, slow computing speed, etc., in the current methods for the satellite gravity gradiometry recovery, we first founded the single and combined error models for the gravity recovery depending on the errors in gravity gradient measurements of the satellite-equipped gravity gradiometer and the errors in orbital position observations of the satellite-borne GPS/GLONASS receiver by the power spectral principle (Welch 1967, Jenkins and Watts 1968, Press et al. 1992a, b), and proposed the matching accuracy indexes of the pivotal sensors and the appropriate orbital parameter of the satellite for the future GOCE Follow-On mission.

## 2. POWER SPECTRUM OF GRAVITY GRADIENT SIGNALS

The spherical harmonic series expansion of the Earth's gravitational potential in the Earth-fixed coordinate system is defined as (Ditmar et al. 2003)

$$
\begin{equation*}
V(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}} \sum_{l=0}^{L}\left(\frac{\mathrm{R}_{\mathrm{e}}}{r}\right)^{l+1} \sum_{m=0}^{l}\left(\bar{C}_{l m} \cos m \lambda+\bar{S}_{l m} \sin m \lambda\right) \overline{\mathrm{P}}_{l m}(\cos \theta) \tag{1}
\end{equation*}
$$

where $G M$ is the product of the gravitational constant $G$ and the Earth's mass $M ; R_{\mathrm{e}}$ denotes the mean radius of the Earth; $L$ represents the maximum de-


Fig. 1. Local North-Stabilized (LNS) frame $O-X Y Z$ and Earth-Centered Inertial (ECI) frame o-xyz.
gree of the spherical harmonic expansion; $r, \theta$, and $\lambda$ are the geocentric radius, geocentric colatitude and geocentric longitude, respectively; $\overline{\mathrm{P}}_{l m}(\cos \theta)$ indicates the normalized associated Legendre polynomials of degree $l$ and order $m$; and $\bar{C}_{l m}, \bar{S}_{l m}$ express the estimated normalized geopotential coefficients.

As shown in Fig. 1, the transformation formula between the Local NorthStabilized (LNS) frame ( $X, Y, Z$ ) and the Earth-Centered Inertial (ECI) frame $(x, y, z)$ can be given by
$\left[\begin{array}{l}x \\ y \\ z\end{array}\right]=\left[\begin{array}{ccc}-\cos (\pi-\lambda) \cos (\pi / 2+\theta) & \sin (\pi-\lambda) & \cos (\pi-\lambda) \sin (\pi / 2+\theta) \\ -\sin (\pi-\lambda) \cos (\pi / 2+\theta) & -\cos (\pi-\lambda) & \sin (\pi-\lambda) \sin (\pi / 2+\theta) \\ \sin (\pi / 2+\theta) & 0 & \cos (\pi / 2+\theta)\end{array}\right]\left[\begin{array}{c}X \\ Y \\ Z+r\end{array}\right]$.
$O-X Y Z$ represents the local north-stabilized frame, where the origin $O$ is located at the center of mass of the satellite, the $X$ - and $Y$-axis are respectively directed to the north and west, and the $Z$-axis completes a right-handed triad with $X$ - and $Y$-axis. Analytic formulas for the second-order derivatives of the Earth's gravitational potentials $V(r, \theta, \lambda)$ at a point with spherical coordinates ( $r, \theta, \lambda$ ) are greatly simple in the local north-stabilized frame (Moritz 1971, Tscherning 1976). The second-order gradients of $V(r, \theta, \lambda)$ with respect to $X, Y, Z$ in the local north-stabilized frame yields

$$
\Gamma=\left[\begin{array}{lll}
V_{X X} & V_{X Y} & V_{X Z}  \tag{3}\\
V_{X X} & V_{Y Y} & V_{Y Z} \\
V_{Z X} & V_{Z Y} & V_{Z Z}
\end{array}\right],
$$

where $\Gamma$ is a symmetrical traceless matrix, $V_{X X}+V_{Y Y}+V_{Z Z}=0$, with five independent components of nine gravity gradient tensors

$$
\left\{\begin{array}{l}
\boldsymbol{V}_{X X}(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}^{3}} \sum_{l=2}^{L}\left(\frac{R_{\mathrm{e}}^{3}}{r}\right)^{l+3} \sum_{m=0}^{l}\left(\overline{\boldsymbol{C}}_{l m} \cos m \lambda+\overline{\boldsymbol{S}}_{l m} \sin m \lambda\right) \boldsymbol{H}^{X X}(\theta) \\
\boldsymbol{V}_{Y Y}(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}^{3}} \sum_{l=2}^{L}\left(\frac{R_{\mathrm{e}}^{3}}{r}\right)^{l+3} \sum_{m=0}^{l}\left(\overline{\boldsymbol{C}}_{l m} \cos m \lambda+\overline{\boldsymbol{S}}_{l m} \sin m \lambda\right) \boldsymbol{H}^{Y Y}(\theta) \\
\boldsymbol{V}_{Z Z}(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}^{3}} \sum_{l=2}^{L}\left(\frac{R_{\mathrm{e}}^{3}}{r}\right)^{l+3} \sum_{m=0}^{l}\left(\overline{\boldsymbol{C}}_{l m} \cos m \lambda+\overline{\boldsymbol{S}}_{l m} \sin m \lambda\right) \boldsymbol{H}^{Z Z}(\theta)  \tag{4}\\
\boldsymbol{V}_{X Y}(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}^{3}} \sum_{l=2}^{L}\left(\frac{R_{\mathrm{e}}^{3}}{r}\right)^{l+3} \sum_{m=0}^{l}\left(-\overline{\boldsymbol{C}}_{l m} \sin m \lambda+\overline{\boldsymbol{S}}_{l m} \cos m \lambda\right) \boldsymbol{H}^{X Y}(\theta) \\
\boldsymbol{V}_{X Z}(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}^{3}} \sum_{l=2}^{L}\left(\frac{R_{\mathrm{e}}^{3}}{r}\right)^{l+3} \sum_{m=0}^{l}\left(\overline{\boldsymbol{C}}_{l m} \cos m \lambda+\overline{\boldsymbol{S}}_{l m} \sin m \lambda\right) \boldsymbol{H}^{X Z}(\theta) \\
\boldsymbol{V}_{Y Z}(r, \theta, \lambda)=\frac{G M}{R_{\mathrm{e}}^{3}} \sum_{l=2}^{L}\left(\frac{R_{\mathrm{e}}^{3}}{r}\right)^{l+3} \sum_{m=0}^{l}\left(-\overline{\boldsymbol{C}}_{l m} \sin m \lambda+\overline{\boldsymbol{S}}_{l m} \cos m \lambda\right) \boldsymbol{H}^{Y Z}(\theta)
\end{array}\right.
$$

where

$$
\left\{\begin{array}{l}
\boldsymbol{H}^{X X}(\theta)=\overline{\boldsymbol{P}}_{l m}^{\prime \prime}(\cos \theta)-(l+1) \overline{\boldsymbol{P}}_{l m}(\cos \theta) \\
\boldsymbol{H}^{Y Y}(\theta)=\tan ^{-1} \theta \overline{\boldsymbol{P}}_{l m}^{\prime}(\cos \theta)-\left(l+1+m^{2} \sin ^{-2} \theta\right) \overline{\boldsymbol{P}}_{l m}(\cos \theta) \\
\boldsymbol{H}^{Z Z}(\theta)=(l+1)(l+2) \overline{\boldsymbol{P}}_{l m}(\cos \theta) \\
\boldsymbol{H}^{X Y}(\theta)=m \sin ^{-1} \theta\left[\overline{\boldsymbol{P}}_{l m}^{\prime}(\cos \theta)-\tan ^{-1} \theta \overline{\boldsymbol{P}}_{l m}(\cos \theta)\right] \\
\boldsymbol{H}^{X Z}(\theta)=(l+2) \overline{\boldsymbol{P}}_{l m}^{\prime}(\cos \theta) \\
\boldsymbol{H}^{Y Z}(\theta)=m(l+2) \sin ^{-1} \theta \overline{\boldsymbol{P}}_{l m}(\cos \theta)
\end{array}\right.
$$

the zero-order, first-order, and second-order derivatives of the Legendre functions are, respectively, represented as

$$
\left\{\begin{array}{l}
\overline{\boldsymbol{P}}_{l m}(\cos \theta)=\gamma_{m} 2^{-l} \sin ^{m} \theta \sum_{k=0}^{[(l-m) / 2]}(-1)^{k} \frac{(2 l-2 k)!}{k!(l-k)!(l-m-2 k)!}(\cos \theta)^{l-m-2 k} \quad(m \leq l) \\
\overline{\boldsymbol{P}}_{l m}^{\prime}(\cos \theta)=\frac{1}{\sin \theta}\left[(l+1) \cos \theta \overline{\boldsymbol{P}}_{l m}(\cos \theta)-(l-m-1) \overline{\boldsymbol{P}}_{l+1, m}(\cos \theta)\right] \\
\overline{\boldsymbol{P}}_{l m}^{\prime \prime}(\cos \theta)=-l \overline{\boldsymbol{P}}_{l m}(\cos \theta)+l \cos \theta \overline{\boldsymbol{P}}_{l-1, m}^{\prime}(\cos \theta)+\frac{l}{4} \cos ^{2} \theta\left[\overline{\boldsymbol{P}}_{l-1, m+1}^{\prime}(\cos \theta)-4 \overline{\boldsymbol{P}}_{l-1, m-1}^{\prime}(\cos \theta)\right]
\end{array}\right.
$$

where

$$
\gamma_{m}=\left\{\begin{array}{ll}
\sqrt{2(2 l+1) \frac{(l-|m|)!}{(l+|m|)!}} & (m \neq 0) \\
\sqrt{2 l+1} & (m=0)
\end{array} .\right.
$$

According to the Parseval's theorem of the spherical harmonic function, the power spectrum of the satellite gravity gradient $V_{a b}$ in the local northstabilized frame is denoted as

$$
\begin{equation*}
P^{2}\left(V_{a b}\right)=\frac{1}{4 \pi} \iint\left[V_{a b}(r, \phi, \lambda)\right]^{2} \cos \phi \mathrm{~d} \phi \mathrm{~d} \lambda, \tag{5}
\end{equation*}
$$

where $a, b=X, Y, Z$.
Combining Eqs. 2, 4, and 5 and the orthogonality of the spherical harmonic function, the power spectrum of the gravity gradient signals in the Earth-centered inertial frame is indicated as (Meng and Liu 1993, van Gelderen and Koop 1997, Zheng et al. 2012c)

$$
\begin{equation*}
P^{2}\left(V_{i j}\right)=\left(\frac{G M}{R_{\mathrm{e}}^{3}}\right)^{2} \sum_{l=0}^{L} A_{i j}^{2}\left(\frac{R_{\mathrm{e}}}{R_{\mathrm{e}}+H}\right)^{2 l+6} \sum_{m=0}^{l}\left(\bar{C}_{l m}^{2}+\bar{S}_{l m}^{2}\right), \tag{6}
\end{equation*}
$$

where $A_{i j}$ is the sensitivity coefficient $(i, j=x, y, z), H$ is the orbital altitude of the satellite, and $\bar{C}_{l m}, \bar{S}_{l m}$ are derived from the Earth gravity field model GO_CONS_GCF_2_TIM_R2 released by the Technical University of Munich (TUM), Germany (Pail et al. 2011a).

Based on Eq. 6, the total power spectrum $P^{2}\left(V_{x y z}\right)$ of the gravity gradient tensor signals is represented as

$$
\begin{equation*}
P^{2}\left(V_{x z z}\right)=P^{2}\left(V_{x x}\right)+P^{2}\left(V_{y y}\right)+P^{2}\left(V_{z z}\right)+P^{2}\left(V_{x z}\right), \tag{7}
\end{equation*}
$$

where $V_{x y z}$ denotes the total signals of the gravity gradient tensors consisting of $V_{x x}, V_{y y}, V_{z z}$, and $V_{x z}$.

By the Kaula's rule (Kaula 1966), the power spectrum $P_{\mathrm{K}}^{2}\left(V_{i j}\right)$ of the satellite gravity gradient tensor signals is shown as

$$
\begin{equation*}
P_{\mathrm{K}}^{2}\left(V_{i j}\right)=\left(\frac{G M}{R_{\mathrm{e}}^{3}}\right)^{2} \sum_{l=0}^{L} A_{i j}^{2}\left(\frac{R_{\mathrm{e}}}{R_{\mathrm{e}}+H}\right)^{2 l+6}(2 l+1) \frac{10^{-10}}{l^{4}}, \tag{8}
\end{equation*}
$$

where the subscript K means "Kaula's rule" in $P_{\mathrm{K}}^{2}\left(V_{i j}\right)$.

Table 2
Sensitivity coefficients of power spectrum from satellite gravity gradient tensors

| Power spectrum of gradients | Sensitivity coefficients $A_{i j}$ |
| :---: | :---: |
| $P^{2}\left(V_{x x}\right)$ | $A_{x x}=-\sqrt{\frac{(l+1)^{3}(l+2)(2 l+3)}{3(2 l+1)}}$ |
| $P^{2}\left(V_{y y}\right)$ | $A_{y y}=-\left[(l+1)(l+2)-\sqrt{\frac{(l+1)^{3}(l+2)(2 l+3)}{3(2 l+1)}}\right]$ |
| $P^{2}\left(V_{z z}\right)$ | $A_{z z}=(l+1)(l+2)$ |
| $P^{2}\left(V_{x z}\right)$ | $A_{x z}=-(l+1) / 5$ |
| $P^{2}\left(V_{x y z}\right)$ | $A_{x y z}=\sqrt{A_{x x}^{2}+A_{y y}^{2}+A_{z z}^{2}+A_{x z}^{2}}$ |



Fig. 2. Sensitivity coefficients $\left|A_{i j}\right|$ of power spectrum from satellite gravity gradient tensors at every degree.

Table 2 lists the expressions for sensitivity coefficients $\left(A_{x x}, A_{y y}, A_{z z}, A_{x z}\right.$, $A_{x y z}$ ) of the gravity gradient tensors ( $V_{x x}, V_{y y}, V_{z z}, V_{x z}, V_{x y z}$ ). Figure 2 shows the sensitivity coefficients $\left|A_{i j}\right|$ of the gravity gradient tensors at every degree, and the statistical results are displayed in Table 3. The research results
are as follows. Firstly, the vertical tensor $V_{z z}$ of the satellite gravity gradients is the uppermost component, which is highly sensitive to the accuracy of the gravity field recovery. Secondly, the horizontal tensors $V_{x x}, V_{y y}$ of the satellite gravity gradients are very important for guaranteeing the accuracy of the Earth's gravitational field measurement. Finally, the cross tensor $V_{x z}$ of the satellite gravity gradients makes a minor contribution to the determination accuracy of the Earth's gravitational field.

Table 3
Statistical results of sensitivity coefficients $\left|A_{i j}\right|$

| Parameters | Sensitivity coefficients $\left\|A_{i j}\right\|$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Degree | Degree | Degree | Degree | Degree | Degree |
|  | 20 | 50 | 100 | 150 | 200 | 250 |
| $A_{x x}$ | 267 | 1531 | 5948 | 13251 | 23442 | 36519 |
| $A_{y y}$ | 195 | 1121 | 4354 | 9701 | 17160 | 26733 |
| $A_{z z}$ | 462 | 2652 | 10302 | 22952 | 40602 | 63252 |
| $A_{x z}$ | 4 | 10 | 20 | 30 | 40 | 50 |
| $A_{x y z}$ | 568 | 6261 | 12668 | 28222 | 49925 | 77776 |



Fig. 3. Signal amplitude spectrum of satellite gravity gradient tensors at every degree.

As illustrated in Fig. 3, the slim dashed, bold solid, bold dashed, and slim solid lines, respectively, represent the signal amplitude spectrum of the Kaula's vertical gravity gradient tensor $V_{\mathrm{K} z z}$ in Eq. 8 and the satellite gravity gradient tensors ( $V_{x x}, V_{y y}, V_{z z}, V_{x z}, V_{x y z}$ ) in Eq. 6, where the orbital altitude of the satellite is $H=250 \mathrm{~km}$, the gravitational constant is $G M=$ $3.986004415 \times 10^{14} \mathrm{~m}^{3} / \mathrm{s}^{2}$, and the mean radius of the Earth is $R_{\mathrm{e}}=6378 \mathrm{~km}$; the statistical results are listed in Table 4. The research results show: (i) the correctness of the power spectrum Eq. 6 of the satellite gravity gradient signals is validated by the conformance of the signal amplitude spectrum between $V_{\mathrm{K} z z}$ and $V_{z z}$; (ii) the vertical component $V_{z z}$ of the gravity gradient signals is stronger, the horizontal components $V_{x x}, V_{y y}$ of the gravity gradient signals are the second, and the cross component $V_{x z}$ of the gravity gradient signals is weaker; (iii) the diagonal components $V_{x x}, V_{y y}, V_{z z}$ of the gravity gradient tensors are absolutely necessary for precisely mapping the Earth's gravitational field with high spatial resolution.

Table 4
Statistics of signal amplitude spectrum of gravity gradients at every degree

| Para- <br> meter | Amplitude spectrum of gradient signals $\left[1 / \mathrm{s}^{2}\right]$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Degree 20 | Degree 50 | Degree 100 | Degree 150 | Degree 200 | Degree 250 |
| $V_{K z z}$ | $4.692 \times 10^{-11}$ | $2.135 \times 10^{-11}$ | $4.276 \times 10^{-12}$ | $7.579 \times 10^{-13}$ | $1.273 \times 10^{-13}$ | $2.075 \times 10^{-14}$ |
| $V_{x x}$ | $1.622 \times 10^{-11}$ | $1.192 \times 10^{-11}$ | $3.053 \times 10^{-12}$ | $5.335 \times 10^{-13}$ | $8.196 \times 10^{-14}$ | $6.273 \times 10^{-15}$ |
| $V_{y y}$ | $1.186 \times 10^{-11}$ | $8.725 \times 10^{-11}$ | $2.235 \times 10^{-12}$ | $3.905 \times 10^{-13}$ | $5.999 \times 10^{-14}$ | $4.592 \times 10^{-15}$ |
| $V_{z z}$ | $2.807 \times 10^{-11}$ | $2.065 \times 10^{-11}$ | $5.288 \times 10^{-12}$ | $9.239 \times 10^{-13}$ | $1.419 \times 10^{-13}$ | $1.086 \times 10^{-14}$ |
| $V_{x z}$ | $2.552 \times 10^{-13}$ | $7.941 \times 10^{-14}$ | $1.036 \times 10^{-14}$ | $1.216 \times 10^{-15}$ | $1.405 \times 10^{-16}$ | $8.623 \times 10^{-18}$ |
| $V_{x y z}$ | $3.452 \times 10^{-11}$ | $2.539 \times 10^{-11}$ | $6.502 \times 10^{-12}$ | $1.136 \times 10^{-12}$ | $1.745 \times 10^{-13}$ | $1.336 \times 10^{-14}$ |

## 3. ERROR MODELS OF SATELLITE GRAVITY GRADIOMETRY

### 3.1 Single error model of the gravity gradient

Combining Eqs. 2, 4, and 5, the power spectrum of the satellite gravity gradient tensor error $\delta V_{i j}$ in the Earth-centered inertial frame is denoted as

$$
\begin{equation*}
P^{2}\left(\delta V_{i j}\right)=\left(\frac{G M}{R_{\mathrm{e}}^{3}}\right)^{2} \sum_{l=0}^{L} A_{i j}^{2}\left(\frac{R_{\mathrm{e}}}{R_{\mathrm{e}}+H}\right)^{2 l+6} \sum_{m=0}^{l}\left(\delta \bar{C}_{l n}\right)^{2}+\left(\delta \bar{S}_{l n}\right)^{2}, \tag{9}
\end{equation*}
$$

where $\delta \bar{l}_{l m}, \delta \bar{S}_{l m}$ are the geopotential coefficient errors.
The cumulative geoid height error is defined as

$$
\begin{equation*}
\sigma_{N}=R_{\mathrm{e}} \sqrt{\sum_{l=2}^{L} \sum_{m=0}^{l}\left(\delta \bar{C}_{l m}\right)^{2}+\left(\delta \bar{S}_{l m}\right)^{2}} . \tag{10}
\end{equation*}
$$

By a combination of Eqs. 9 with 10, the cumulative geoid height error of the satellite gravity gradient tensors is represented as

$$
\begin{equation*}
\sigma_{N}\left(\delta V_{i j}\right)=\frac{R_{\mathrm{e}}^{4}}{G M(\sqrt{T / \Delta t})} \sqrt{\sum_{l=2}^{L} \frac{2 l+1}{A_{i j}^{2}}\left(\frac{R_{\mathrm{e}}+H}{R_{\mathrm{e}}}\right)^{2 l+6} \sigma^{2}\left(\delta V_{i j}\right)}, \tag{11}
\end{equation*}
$$

where $\sigma^{2}\left(\delta V_{i j}\right)$ is the error variance of the gravity gradient tensors. $T$ is the observation time of the gravity gradient data, $\Delta t$ is a sampling interval of measurements, and $T / \Delta t$ is the amount of the gravity gradients. According to the fundamental principles of statistics, if the number of the gravity gradient tensors is increased by $T / \Delta t$ times, the recovery accuracy of the Earth's gravitational field should be improved by about $\sqrt{T / \Delta t}$ times.

From Eq. 11 and Table 2, the cumulative geoid height errors of the satellite gravity gradient tensor are shown as

$$
\begin{align*}
& \sigma_{N}\left(\delta V_{x x}\right)=\frac{R_{\mathrm{e}}^{4}}{G M(\sqrt{T / \Delta t})} \sqrt{\sum_{l=2}^{L} \frac{2 l+1}{\frac{(l+1)^{3}(l+2)(2 l+3)}{3(2 l+1)}\left(\frac{R_{\mathrm{e}}+H}{R_{\mathrm{e}}}\right)^{2 l+6} \sigma^{2}\left(\delta V_{x x}\right)},}  \tag{12}\\
& \sigma_{N}\left(\delta V_{y y}\right)=\frac{R_{\mathrm{e}}^{4}}{G M(\sqrt{T / \Delta t})} \sqrt{\left.\sum_{l=2}^{L} \frac{2 l+1}{\left[(l+1)(l+2)-\sqrt{\left.\frac{(l+1)^{3}(l+2)(2 l+3)}{3(2 l+1)}\right]^{2}}\right.}\right]^{\left(\frac{R_{\mathrm{e}}+H}{R_{\mathrm{e}}}\right)^{2 l+6}} \sigma^{2}\left(\delta V_{y y}\right)}, \tag{13}
\end{align*}
$$

$$
\begin{equation*}
\sigma_{N}\left(\delta V_{z z}\right)=\frac{R_{\mathrm{e}}^{4}}{G M(\sqrt{T / \Delta t})} \sqrt{\sum_{l=2}^{L} \frac{2 l+1}{(l+1)^{2}(l+2)^{2}}\left(\frac{R_{\mathrm{e}}+H}{R_{\mathrm{e}}}\right)^{2 l+6} \sigma^{2}\left(\delta V_{z z}\right)} \tag{14}
\end{equation*}
$$

$$
\sigma_{N}\left(\delta V_{x z}\right)=\frac{R_{\mathrm{e}}^{4}}{G M(\sqrt{T / \Delta t})} \sqrt{\sum_{l=2}^{L} \frac{2 l+1}{[(l+1) / 5]^{2}}\left(\frac{R_{\mathrm{e}}+H}{R_{\mathrm{e}}}\right)^{2 l+6} \sigma^{2}\left(\delta V_{x z}\right)}
$$

$$
\sigma_{N}\left(\delta V_{y z}\right)=\frac{R_{\mathrm{e}}^{4}}{G M(\sqrt{T / \Delta t}} \times
$$

$$
\begin{equation*}
\times \sqrt{\sum_{l=2}^{L} \frac{2 l+1}{\frac{(l+1)^{3}(l+2)(l+3)}{3(2 l+1)}+\left[(l+1)(l+2)-\sqrt{\frac{(l+1)^{3}(l+2)(2 l+3)}{3(2+1)}}\right]^{2}+(l+1)^{2}(l+2)^{2}+[(l+1) / 5]^{2}}\left(\frac{R_{z}+H}{R_{c}}\right)^{2+6} \sigma^{2}\left(\delta V_{x y}\right)} . \tag{16}
\end{equation*}
$$

### 3.2 Single error model of orbital position

In the Earth-centered inertial frame, based on the power spectral principle, the power spectrum of the total error $\delta V_{x y z}$ of the gravity gradient tensor is displayed as

$$
\begin{equation*}
P^{2}\left(\delta V_{x y z}\right)=\left(-\frac{3 G M}{r^{4}}\right)^{2} \sigma^{2}(\delta r), \tag{17}
\end{equation*}
$$

where $V_{x y z}=\ddot{r} / r$ is the satellite gravity gradient tensor, and $r$ and $\ddot{r}$ denote the orbital position and centripetal acceleration of the satellite; $\sigma^{2}(\delta r)$ is the error variance of the satellite orbital position; and $P^{2}\left(\delta V_{x y z}\right)$ is given by

$$
\begin{equation*}
P^{2}\left(\delta V_{x y z}\right)=\frac{\sigma^{2}\left(\delta V_{x y z}\right)}{L_{\max }} \tag{18}
\end{equation*}
$$

where $\sigma^{2}\left(\delta V_{x y z}\right)$ is the error variance of the satellite gravity gradient, and $L_{\text {max }}$ is the maximum spherical harmonic degree of the gravity recovery in theory. However, since the high-frequency signals will be submerged into the observation errors, the maximum degree in fact is lower than that in theory

$$
\begin{equation*}
L_{\max }=\frac{\pi r}{D} \tag{19}
\end{equation*}
$$

where $D=\dot{r}_{0} \Delta t$ is the spatial resolution (half-wavelength), and $\dot{r}_{0}=\sqrt{G M / r}$ is the mean velocity of the satellite.

Combining Eqs. 17-19, the transformational relation between the gravity gradient error $\delta V_{x y z}$ and the orbital position error $\delta r$ is denoted as

$$
\begin{equation*}
\delta V_{x y z}=\sqrt{\frac{9 G M \pi^{2}}{r^{5} \Delta t^{2}}} \delta r . \tag{20}
\end{equation*}
$$

Based on Eqs. 16 and 20, the error model of the cumulative geoid height impacted by the orbital position error is shown as

$$
\sigma_{N}(\delta r)=\frac{R_{t}^{t}}{G M(\sqrt{T / \Delta t})} \times
$$

$$
\begin{equation*}
\times \sqrt{\sum_{l=2}^{L} \frac{2 l+1}{\frac{(l+1)^{3}(l+2)(2 l+3)}{3(l+1)}+\left[(l+1)(l+2)-\sqrt{\frac{(l+1)^{3}(l+2)(l l+3)}{3(2+1)}}\right]^{-2}+(l+1)^{2}(l+2)^{2}+[(l+1) 5]^{2}}}\left(\frac{R+H}{R}\right)^{2+6} \sigma^{2}\left(\sqrt{\frac{9 G M \hbar^{2}}{r^{5} \Delta^{2}}} \delta r\right) \tag{21}
\end{equation*}
$$

### 3.3 Combined error model between gravity gradient and orbital position

According to Eqs. 16 and 21, the combined error model of the cumulative geoid height affected by the gravity gradient error of the satellite gravity gradiometer and the orbital position error of the GPS/GLONASS receiver is expressed as

$$
\begin{align*}
& \sigma_{N}\left(\delta V_{n y 2}, \delta r\right)=\frac{R_{\varepsilon}^{4}}{G M(\sqrt{T \mid \Delta t}} \times \\
& \times \sqrt{\sum_{l=2}^{\frac{L}{l} \frac{(l+1)^{3}(l+2)(2 l+3)}{3(2 l+1)}+\left[(l+1)(l+2)-\sqrt{\frac{(l+1)^{3}(l+2)(2 l+3)}{3(2 l+1)}}\right]^{2}+(l+1)^{2}(l+2)^{2}+[(l+1) / 5]^{2}}\left(\frac{R_{\varepsilon}+H}{R}\right)^{2 l+6} \sigma^{2}(\delta \eta)} \tag{22}
\end{align*},
$$

where $\delta \eta=\sqrt{\sigma^{2}\left(\delta V_{x y z}\right)+\sigma^{2}\left(\sqrt{\frac{9 G M \pi^{2}}{r^{5} \Delta t^{2}}} \delta r\right)}$ is the total error of key instruments from the GOCE satellite, $\sigma^{2}\left(\delta V_{x y z}\right)$ is the error variance of the gravity gradient tensors, and $\sigma^{2}\left(\sqrt{\frac{9 G M \pi^{2}}{r^{5} \Delta t^{2}}} \delta r\right)$ is the error variance of the satellite orbital position.

## 4. RESULTS

### 4.1 Verification of single and combined error models

As displayed in Fig. 4, the solid, dashed and asterisk lines represent the cumulative geoid height errors impacted by the single gravity gradient error $\delta V_{x y z}$ of the GOCE gravity gradiometer, the single orbital position error $\delta r$ of the GPS/GLONASS receiver, and the combined error $\delta\left(V_{x y z}+r\right)$, respectively. The statistical results are listed in Table 5, the accuracy indexes of the GOCE key payloads are shown in Table 6, and the related parameters of the error model are represented in Table 7. In terms of the conformity between solid and dashed lines in Fig. 4, we should know that the accuracy indexes of the GOCE key sensors provided by this paper in Table 6 are matched with each other. Additionally, the correctness of the single error model of the satellite gravity gradient (see Eq. 16) and the dependability of the single error model of the satellite orbital position (see Eq. 21) are verified by the conformance of the accuracy indexes from the GOCE key instruments provided by this paper and the ESA. At degree 250, the cumulative geoid height error is $1.769 \times 10^{-1} \mathrm{~m}$ based on the combined error model (see Eq. 22) derived


Fig. 4. Cumulative geoid height errors based on the matching measurement accuracies of GOCE spaceborne instrumentations.

Table 5
Statistics of cumulative geoid height error based on the matching accuracy indexes of key payloads

| Observation <br> error | Cumulative geoid height error [m] |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Degree 50 | Degree 100 | Degree 150 | Degree 200 | Degree 250 | Degree 300 |
| Gravity <br> gradient <br> $\delta V_{x y z}$ | $2.083 \times 10^{-3}$ | $2.660 \times 10^{-3}$ | $6.228 \times 10^{-3}$ | $2.492 \times 10^{-2}$ | $1.185 \times 10^{-1}$ | $6.068 \times 10^{-1}$ |
| Orbital <br> position <br> $\delta r$ | $2.310 \times 10^{-3}$ | $2.950 \times 10^{-3}$ | $2.906 \times 10^{-3}$ | $2.765 \times 10^{-2}$ | $1.313 \times 10^{-1}$ | $6.730 \times 10^{-1}$ |
| Combined <br> model <br> $\delta\left(V_{x y z}+r\right)$ | $3.110 \times 10^{-3}$ | $3.972 \times 10^{-3}$ | $9.299 \times 10^{-3}$ | $3.722 \times 10^{-2}$ | $1.769 \times 10^{-1}$ | $9.062 \times 10^{-1}$ |

from the gravity gradient error and the orbital position error, which preferably accords with the measurement accuracy of $1.760 \times 10^{-1} \mathrm{~m}$ from the Earth gravity field model GO_CONS_GCF_2_TIM_R2. Therefore, the combined error model established by this study is correct.

Table 6
Matching accuracy indexes of GOCE spaceborne instruments

| Observations | Accuracy indexes |
| :---: | :---: |
| Gravity gradient | $3 \times 10^{-12} / \mathrm{s}^{2}$ |
| Orbital position | $1 \times 10^{-2} \mathrm{~m}$ |

Table 7
Related parameters of GOCE error models

| Parameters | Indexes |
| :--- | :---: |
| Orbital altitude $H$ | 250 km |
| Earth's radius $R_{\mathrm{e}}$ | 6370 km |
| Observation time $T$ | 8 months |
| Sampling interval $\Delta t$ | 5 s |
| Gravitational constant $G M$ | $3.986004415 \times 10^{14} \mathrm{~m}^{3} / \mathrm{s}^{2}$ |

### 4.2 Requirements analysis for GOCE Follow-On mission

### 4.2.1 Impact of observation error in gravity gradient

As illustrated in Fig. 5, the cumulative geoid height errors from GOCE Fol-low-On up to degree 300 are estimated based on different measurement accuracies of $3 \times 10^{-12} / \mathrm{s}^{2}, 3 \times 10^{-13} / \mathrm{s}^{2}, 3 \times 10^{-14} / \mathrm{s}^{2}$, and $3 \times 10^{-15} / \mathrm{s}^{2}$ from the satellite gravity gradients using the single error model of the gravity gradient (see Eq. 16, Tables 6 and 7), respectively. At degree 300, the cumulative geoid height error is $6.068 \times 10^{-1} \mathrm{~m}$ using the gravity gradient error of $3 \times 10^{-12} / \mathrm{s}^{2}$. If the measurement accuracies of the gravity gradient are designed as $3 \times 10^{-13} / \mathrm{s}^{2}, 3 \times 10^{-14} / \mathrm{s}^{2}$, and $3 \times 10^{-15} / \mathrm{s}^{2}$, the cumulative geoid height errors will be improved by 10 times, 100 times, and 1000 times, respectively. The geophysical analysis is described as follows. The measurement precision of the space-borne gravity gradiometer plays a very significant part in improving the recovery accuracy of the Earth's gravitational field. Therefore, if the cold-atom interferometric gradiometer is used in the next-generation satellite gravity mission from GOCE Follow-On with a measurement accuracy of $10^{-13}-10^{-15} / \mathrm{s}^{2}$ (Yu et al. 2006, Johnson 2011), the determination accuracy of the Earth's gravitational field from the future GOCE Follow-On satellite is at least 10 times higher than that from the current GOCE satellite.


Fig. 5. Cumulative geoid height errors based on different measurement accuracies of satellite gravity gradients.

### 4.2.2 Influence of measurement accuracy of orbital position

Figure 6 displays the GOCE Follow-On cumulative geoid height error up to degree 300 using the measurement accuracies of $10^{-2}, 10^{-3}, 10^{-4}$, and $10^{-5} \mathrm{~m}$


Fig. 6. Cumulative geoid height errors based on different measuring accuracies of satellite orbital position.
for the orbital position of the satellite by the single error model of the orbital position (see Eq. 21, Tables 6 and 7). At degree 300, the cumulative geoid height error is $6.730 \times 10^{-1} \mathrm{~m}$ using an orbital accuracy of $10^{-2} \mathrm{~m}$, and the cumulative geoid height errors are linearly improved while the measurement accuracies of the satellite orbital position are enhanced by 10 times, 100 times, and 1000 times, respectively. The geophysical analysis is depicted as follows. The satellite gravity gradiometry is less sensitive to the precise orbit determination of the satellite and the orbital accuracy of the Global Navigation Satellite System (GNSS) (e.g., GPS, GLONASS, Compass, Galileo, etc.) is about cm-level. Therefore, the measurement accuracy of the satellite orbital position from the future GOCE Follow-On mission is preferable for $1-0.1 \mathrm{~cm}$.

### 4.2.3 Effect of orbital altitude of satellite

As illustrated in Fig. 7, the bold solid, bold dashed, slim dashed, and slim solid lines, respectively, represent the cumulative geoid height errors up to degree 300 by different satellite orbital altitudes of $200,250,300$, and 350 km based on the GOCE Follow-On combined error model (see Eq. 22, Tables 6 and 7), and the statistical results are listed in Table 8. At degree 300, the cumulative geoid height error is $1.049 \times 10^{-1} \mathrm{~m}$ with an orbital altitude of 200 km . If the GOCE Follow-On satellite operates at the orbital altitudes of 250,300 , and 350 km , the cumulative geoid height errors are reduced by 8.639 times, 75.491 times, and 660.819 times, respectively. The


Fig. 7. Cumulative geoid height errors using different satellite orbital altitudes.

Statistical results of impacts of different orbital altitudes on cumulative geoid height error

| Orbital <br> altitude <br> $[\mathrm{km}]$ | Cumulative geoid height error [m] |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Degree 50 | Degree 100 | Degree 150 | Degree 200 | Degree 250 | Degree 300 |
| $H_{1}=200$ | $2.885 \times 10^{-3}$ | $3.170 \times 10^{-3}$ | $4.366 \times 10^{-3}$ | $9.776 \times 10^{-3}$ | $3.025 \times 10^{-2}$ | $1.049 \times 10^{-1}$ |
| $H_{2}=250$ | $3.110 \times 10^{-3}$ | $3.972 \times 10^{-3}$ | $9.299 \times 10^{-3}$ | $3.722 \times 10^{-2}$ | $1.769 \times 10^{-1}$ | $9.062 \times 10^{-1}$ |
| $H_{3}=300$ | $3.403 \times 10^{-3}$ | $5.816 \times 10^{-3}$ | $2.448 \times 10^{-2}$ | $1.512 \times 10^{-1}$ | $1.058 \times 10^{0}$ | $7.919 \times 10^{0}$ |
| $H_{4}=350$ | $3.800 \times 10^{-3}$ | $9.882 \times 10^{-3}$ | $6.836 \times 10^{-2}$ | $6.238 \times 10^{-1}$ | $6.360 \times 10^{0}$ | $6.932 \times 10^{1}$ |

geophysical analysis is recounted as follows. Because the Earth's gravitational potential presents a tendency of the exponential attenuation with increasing the orbital altitude of the satellite, the lower orbital altitude makes a great contribution to substantially improve the accuracy of the Earth's gravitational field determination. However, the atmospheric drag acting on the satellite will be enhanced by one order of magnitude with the decrease in the orbital altitude by per 100 km . Therefore, although the GOCE Follow-On satellite carries the drag-free control system, the optimal design of the satellite orbital altitude is very crucial all the time. To sum up, we suggest that it is suitable for selecting the orbital altitude range of $200-250 \mathrm{~km}$.

### 4.2.4 Gravity recovery from GOCE Follow-On

As illustrated in Fig. 8, the dashed line represents the GOCE-only Earth gravity field model GO_CONS_GCF_2_TIM_R2 complete up to degree and order 250 released by the TUM, Germany, and the cumulative geoid height error is $1.760 \times 10^{-1} \mathrm{~m}$ at degree 250 . The solid line denotes the cumulative geoid height error up to degree 400 from the future GOCE Follow-On mission by the combined error model (see Eq. 22), based on the gravity gradient error of $3 \times 10^{-13} / \mathrm{s}^{2}$, orbital position error of 0.1 cm and orbital altitude of 200 km , and adopting an observation period of 8 months and a 5-s sampling interval. The statistical results of the cumulative geoid height errors are displayed in Table 9. The reasons why the accuracy of the Earth's gravitational field based on the future GOCE Follow-On mission is improved by a factor of more than 10 as compared with that based on the current GOCE mission are expatiated as follows.

Firstly, the orbital altitude of the future GOCE Follow-On satellite ( 200 km ) is lower than that of the current GOCE satellite ( 250 km ). Hence, the negative effects on the signal attenuation of gravity information with the increase in the orbital altitude are propitious to be mitigated to a great extent by the future GOCE Follow-On satellite.


Fig. 8. A comparison of cumulative geoid height errors between the current GOCE and future GOCE Follow-On satellites.

Table 9
Statistical results of cumulative geoid height errors from GOCE and GOCE Follow-On

|  | Cumulative geoid height error [m] |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Gravity satellites | Degree <br> 50 | Degree <br> 100 | Degree <br> 150 | Degree <br> 200 | Degree <br> 250 | Degree <br> 300 | Degree <br> 350 |
|  | 1.493 | 3.225 | Degree <br> 400 |  |  |  |  |  |
|  | $\times 10^{-2}$ | $\times 10^{-2}$ | $\times 10^{-2}$ | 1.198 |  |  |  |  |
|  | 1.760 | $-10^{-1}$ | - | - | - |  |  |  |
|  | 1.493 | 3.171 | 4.366 | 9.776 | 3.025 | 1.049 | 3.843 | 1.458 |
|  | $\times 10^{-4}$ | $\times 10^{-4}$ | $\times 10^{-4}$ | $\times 10^{-4}$ | $\times 10^{-3}$ | $\times 10^{-2}$ | $\times 10^{-2}$ | $\times 10^{-1}$ |

Secondly, the measurement accuracy of the space-borne instruments (e.g., cold-atom interferometric gravity gradiometer) from GOCE Follow-On is at least one order of magnitude higher than that (e.g., electrostatic suspension gravity gradiometer) from GOCE. Herewith, the accuracy of the Earth's gravitational field from GOCE Follow-On is conducive to be substantially improved due to an optimum Signal-to-Noise Ratio (SNR) of the satellite observations.

## 5. CONCLUSION

The core objectives of this investigation are to establish the recovery error model of the satellite gravity gradiometry through the power spectral principle and perform the requirements analysis for the future GOCE Follow-On mission. The summary of the concrete results is stated as follows.

Firstly, we developed the single and combined error models aiming at availably and rapidly estimating the accuracy of the gravity field by adopting the errors in gravity gradient and orbital position.

- We demonstrated the matching relationship of the measurement accuracies from the GOCE space-borne sensors making use of the single error model, and proved the reliability of the single error model by the measurement accuracy of this study in accordance with that of ESA.
- We validated the dependability of the combined error model. At degree 250, the cumulative geoid height error of this study basically tallies with that of the existing GO_CONS_GCF_2_TIM_R2 model.
Secondly, we implemented the sensitivity analysis for the nextgeneration GOCE Follow-On system regarding different measurement accuracies of the gravity gradient, different observation errors in the orbital position, and different orbital altitudes of the satellite.
- The recovery accuracy of the Earth's gravitational field from GOCE Fol-low-On is at lowest 10 times superior to that from GOCE if the cold-atom interferometric gradiometer is applied in the future GOCE Follow-On satellite. Accordingly, it is a preferred choice for the measurement accuracy of $10^{-13}-10^{-15} / \mathrm{s}^{2}$ of the gravity gradient in the next-generation GOCE Follow-On mission.
- Because the satellite gravity gradiometry measurement has a low sensitivity to the observation accuracy of the orbital position and the topmost precision of the GPS, GLONASS, Compass, and Galileo orbit determination is just cm-level at present, an orbital accuracy of $1-0.1 \mathrm{~cm}$ is an optimal design for the future GOCE Follow-On system.
- The signal strength of the Earth's gravitational field is apt to be weakened with the increase in the orbital altitude of the satellite, and the adverse influences of the non-conservation force will be sharply added with reducing the orbital altitude at the same time. Consequently, the future GOCE Follow-On satellite had better operate at an orbital altitude of 200250 km .
- The accuracy of the gravity recovery from the future GOCE Follow-On satellite sufficiently surpasses the current GOCE satellite on account of the lower orbital altitude of the satellite and the higher observation accuracy of the key payloads.

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