

Absolute gravity change associated with the March 1997 earthquake swarm in the Izu Peninsula, Japan

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We carried out both absolute and relative gravity measurements in the Izu Peninsula just before and after the March 1997 earthquake swarm occurred. The measurements revealed significant absolute gravity changes, which we find to be made of three spatial components. The first one is located near Cape Kawana, and would be associated with the volcanic activity that caused the earthquake swarm. The second one would be associated with shallow and localized magma intrusion just beneath Ito. The third one may be due to a change in the deep region beneath the Kita-Izu fault system, which is considered to be a major tectonic line of this region. The gravity changes can be used to detect underground mass movement. For this purpose, we first use crustal movement observations to construct an elastic dislocation model with two tensile faults and a left lateral fault. Then we use the gravity changes to constrain the density of the material which filled the tensile faults. We find that the density is likely to be small, and that the gravity changes of the first component are reproduced well by the fault model. The smallness of the density implies that highly vesiculated magma or water would have injected into the faults.

1. Introduction

Earthquake swarms occur once or twice a year around the eastern coast of the Izu Peninsula (Fig. 1) (e.g. Hori, 1989; Mogi, 1991) since 1978, when the Izu-Oshima-Kinkai-Oki earthquake (M7.0) broke out. They are often accompanied by crustal deformation around Ito (Fig. 1(a)).

The relation between earthquake swarms and monogenetic volcanos in this area has long been discussed because of the near similar location of the epicenters and the Higashi-Izu monogenetic volcano group. The connection was evidenced by the sequence of 1989 earthquake swarm and the Teishi-Knoll submarine eruption (see the special issue of *J. Phys. Earth*, Vol. 39, No. 1 for details), which may be regarded as Higashi-Izu volcano group activity. Discussion on the relation between swarms and volcanism dates back to 1930, when an earthquake swarm occurred around eastern Izu. Imamura Laboratory and Seismological Laboratory (1930), Ishimoto and Takahashi (1930), and Nasu (1930, 1935) suggested a magmatic origin of the swarm and the associated crustal uplift. Tsuya (1930) and Kuno (1954) discussed its relation to geological structure. Aramaki and Hamuro (1977) reported the connection between volcanism around Izu and 1975–1976 swarms. The relation between swarms and volcanism has often been discussed since earthquake swarms became frequent in 1978.

An earthquake swarm started on March 3, 1997, and continued for about three weeks. The source area was off Shiofuki-zaki at the beginning of March 3, and then migrated to off Kawana-zaki at about 0800 LT March 3 (Fig. 1). The activity declined with time and nearly ended in about a week (Earthquake Research Institute, 1997; Japan Meteorological Agency, 1997). The total number of the earthquakes recorded at Kamata, Ito City, from 3 to 26 March was 9334 according to Japan Meteorological Agency (1997). The number of the earthquakes larger than M5.0 was 4. They occurred during the early stage on March 3 and 4. The mechanisms of most of the earthquakes are strike-slip type with the P-axis in the NW-SE direction, which is in accord with the regional stress field (e.g. Tsukahara and Ikeda, 1987). This stress pattern, the P-axis being in the NW-SE direction, is interpreted in terms of the collision of the Philippine Sea plate against the Japan Island (e.g. Ukawa, 1991; Koyama and Umino, 1991) or the bending of the slab (Nakamura *et al.*, 1984).

Many kinds of measurements of crustal movement have been carried out around Ito, including GPS, leveling (Geographical Survey Institute, 1997), and EDM (Tsuneishi, 1997), in order to clarify the mechanism of earthquake swarms. During the March swarm, movements showed complex spatial variation, which we examine in Section 3.

Gravity measurements are complementary to these geometrical observations of crustal movement in that they can detect underground mass movement. For revealing magma or fluid motion which would cause earthquake swarms, our group has carried out gravity measurements once or twice a year since 1974, when the Izu-Hanto-Oki earthquake (M6.9) occurred. The results from 1974 to 1985 were summarized by Hagiwara *et al.* (1985). They reported significant gravity

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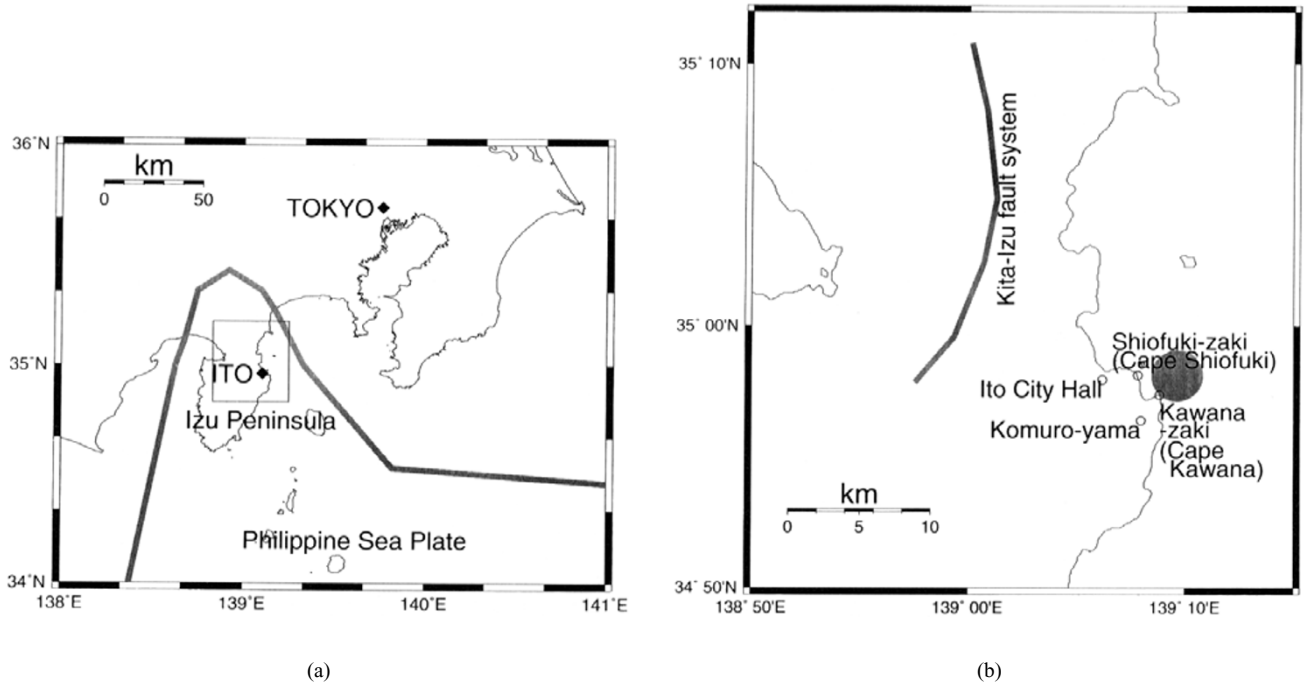


Fig. 1. Index map of the study area. (a) Map of the South Kanto area. The gray line represents the boundary of the Philippine Sea Plate. The inset rectangle indicates the area of Fig. 1(b). (b) Map of the north-eastern part of the Izu Peninsula. The shaded circle represents the central region of the earthquake swarm activity in March 1997.

Table 1. Details of the absolute gravity measurements.

	Before the swarm	After the swarm
Measurement dates	February 3, 5–7, 1997	March 8–14, 1997
Absolute gravity value (μgal)	979798779.1 ± 0.3	979798775.0 ± 0.3
Number of total effective drops	10,162	20,259
Standard deviation of a drop (μgal)	34	37
Pole tide (μgal)	0.2	0.4

changes due to large earthquakes. Okubo *et al.* (1991) detected gravity changes caused by the 1989 earthquake swarm and the associated submarine eruption of Teishi Knoll. They interpreted the changes in terms of an elastic deformation due to a tensile fault, a strike-slip fault and an inflation source. All these measurements before 1997 were performed with only LaCoste-Romberg gravimeters, which allowed them to detect only relative gravity changes.

Absolute gravimetry is required to measure temporal gravity changes unequivocally. For this purpose, our group introduced an absolute gravimeter FG5 (#109) (Niebauer *et al.*, 1995) in 1994 (Okubo *et al.*, 1997). We measured the absolute gravity in Ito in February 1997, only a month before the swarm in March. The measurement along with that in March 1997 enabled us to detect absolute gravity changes associated with the earthquake swarm activity. This is the first attempt to measure *absolute* gravity changes of seismic and volcanic origins. We present the results together with their interpretation in the following.

2. Gravity Measurements

2.1 Absolute gravity measurements

We installed the FG5 absolute gravimeter in the Ito City Hall to establish the reference point for our precise gravity network. A gravity decrease of $4 \mu\text{gal}$ was detected there after the earthquake swarm compared with the value of a month before. This value is used as the reference of gravity changes measured by a LaCoste-Romberg gravimeter. The results of the relative measurements are described in Subsection 2.2.

The results of the absolute gravity observations and their details are summarized in Table 1. The absolute measurements are corrected for the instrumental height, the solid earth tide, the pole tide, and the atmospheric pressure. The gravity values are reduced to those at the height of 130 cm above the floor with a gravity gradient of $-2.85 \mu\text{gal}/\text{cm}$. The solid earth tide is corrected with the tidal gravity factor $\delta = 1.164$, except for the permanent tide, for which $\delta = 1.0$ is used. The barometric correction was made with a standard air pressure of 1007.3 hPa and an admittance of $0.3 \mu\text{gal}/\text{hPa}$. The standard error of each absolute gravity

Table 2. Gravity data before and after the earthquake swarm and their difference Δg . g is the gravity values relative to that of the Ito City Hall in February in the unit of μgal . Add 979799149.6 μgal (absolute gravity at the floor in the Ito City Hall in February) to g to obtain the absolute gravity values in February and March. ϵ denotes the estimated observational error in the unit of μgal . The method of the error estimation is described in the text. N is the number of relative gravity measurements carried out at each site. At least two measurements were performed at each site because we used two-way method. The multiple measurements at each point are used to correct the instrumental drift and to improve the error. The location of the sites are mapped on Fig. 2(b). The data of the sites indicated by asterisks are not used in estimating the densities of the materials intruding into tensile faults in Section 3.

Site name	Latitude (deg.)	Longitude (deg.)	February 6, 17–21			March 8–14			Change	
			g	ϵ	N	g	ϵ	N	Δg	ϵ
Ito-City-Hall (Reference)	34.9623	139.1047	0		5	−4		6	−4	
Ito-CityBM	34.9625	139.1045	−316	4	11	−326	3	32	−10	5
BM-9402	35.0417	138.9057	−2,553	9	4	−2,561	7	2	−8	11
BM-9328	35.0588	139.0742	7,951	8	2	7,942	8	2	−9	11
BM-9330	35.0457	139.0978	10,665	8	2	10,660	8	2	−5	11
BM-9332	35.0075	139.0918	3,671	8	2	3,669	8	2	−2	11
BM-9334	34.9765	139.0965	5,476	5	11	5,475	5	4	−1	7
BM-9335	34.9647	139.1037	6,107	8	2	6,095	7	2	−12	11
BM-9336*	34.9517	139.1123	−15,920	8	2	−15,903	6	4	17	10
BM-9337	34.9400	139.1212	−20,132	5	4	−20,146	8	2	−14	9
BM-9340	34.8933	139.1233	−32,326	7	2	−32,337	8	2	−11	11
BM-9341	34.8753	139.1068	−14,930	7	2	−14,936	8	2	−6	11
BM-9344	34.8402	139.0790	−367	7	2	−370	8	2	−3	11
BM-10005	34.9020	139.0437	−129,590	10	2	−129,591	5	4	−1	11
BM-003-002	34.9538	139.0863	−10,154	10	2	−10,158	7	2	−4	12
BM-003-005	34.9470	139.0785	−45,667	10	2	−45,671	7	2	−4	12
BM-003-006	34.9460	139.0703	−60,795	10	2	−60,798	6	4	−3	12
BM-003-007	34.9487	139.0650	−60,700	10	2	−60,710	7	2	−10	12
BM-003-012	34.9438	139.0270	−40,047	8	7	−40,055	6	4	−8	10
BM-003-017*	34.9623	138.9822	−26,215	9	2	−26,228	7	2	−13	11
Hatsushima-Port	35.0392	139.1723	−2,274	9	2	−2,278	8	2	−5	12
Hatsuki-Shrine	35.0393	139.1725	−3,851	9	2	−3,858	8	2	−7	12
Fishery Radio Station	34.9613	139.1217	−40,938	6	2	−40,839	5	6	−1	8
Kawana-Suwa	34.9492	139.1407	10,119	6	2	10,102	6	4	−17	8
Nakaizu*	34.9870	139.0385	−113,195	9	2	−113,225	6	4	−30	11
Kameishi*	35.0203	139.0462	−92,158	9	2	−92,175	6	4	−17	11
Atami-Pass*	35.1168	139.0422	−124,403	9	2	−124,425	7	2	−22	11
10KM-PostBM*	35.0610	139.0377	−123,068	9	2	−123,085	7	2	−17	11

value is about 0.3 μgal , the standard deviation of a single drop divided by the square root of the number of drops. The 1–2 μgal precision of the FG5 gravimeter was confirmed in advance by Okubo *et al.* (1997) by detecting gravity signals due to the ocean tide.

2.2 Relative gravity measurements

We used a LaCoste-Romberg gravimeter (G-581) to measure precise relative gravity changes around Ito. The measurements were carried out on February 6, 17–21, 1997,

before the earthquake swarm and on March 8–14 after the swarm. Relative gravity is measured as the differences from the value at the Ito City Hall, so that the absolute values for all sites can be obtained by adding the differences to the absolute value measured with FG5 (Table 2). The two-way method was taken to correct the instrumental drift and to ensure the precision. We estimated the observational errors as 5–10 μgal from the two-way repeatability as tabulated in Table 2. We removed the systematic cyclic error (Becker *et*

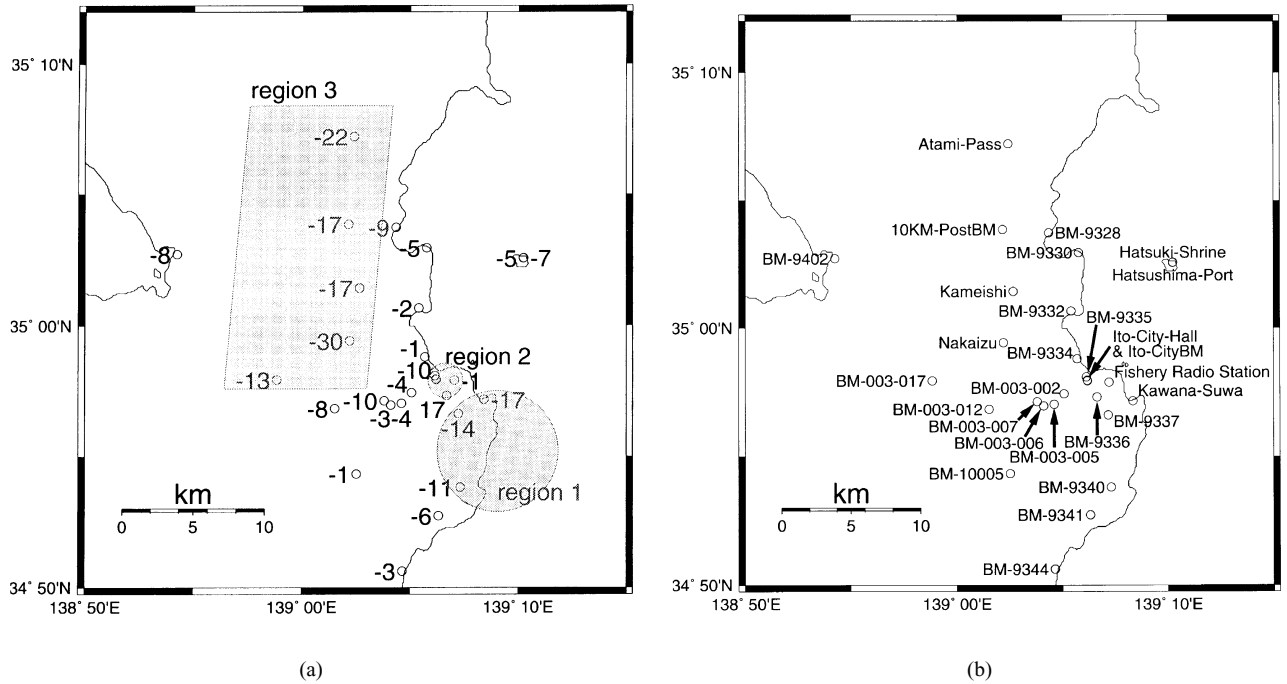


Fig. 2. (a) Absolute gravity changes between February and March 1997 measured with an FG5 absolute gravimeter and a LaCoste-Romberg gravimeter. The unit is μgal . The details are given in Table 2. The precisions are 5–10 μgal . The shaded areas indicate the regions where significant gravity changes were observed. (b) Site map of relative gravity measurements.

al., 1987) and the magnetic error (e.g. Torge, 1989) inherent in any LaCoste-Romberg gravimeter to attain the precision.

The error at each point is estimated as follows. We define the two-way repeatability as a half of the maximum value of the deviations of the two-way differences from the linear drift in one set of two-way measurements. The error at each site is first estimated as the two-way repeatability defined above if it is larger than 5 μgal . The error is set to be 5 μgal if the two-way repeatability is less than 5 μgal . Then we apply network adjustment to the whole sets of two-way measurements to estimate internally consistent gravity values and their error covariances. The errors shown in Table 2 are derived from these covariances.

Figure 2 shows the absolute gravity changes between February and March 1997. We find three areas where the gravity changes amount to more than about 10 μgal . (regions 1, 2, and 3 in Fig. 2(a)). In the region 1 near Kawana-zaki, we observed gravity decreases up to 17 μgal . The decreases are explained well by elastic deformation due to strike-slip and tensile faulting. The elastic model is elaborated in the next section. The region 2 is located in the Ito City to the northwest of the first region. It is characterized by rather localized and complicated changes up to about 10 μgal . The changes would be caused by small dikes intruding into the very shallow part of the crust. It could otherwise be due to observational errors, because of the rather noisy environment near the observation sites BM-9336 and BM-9337. The region 3 is located in the mountainous area around the neck of the Izu Peninsula. It is characterized by large gravity decreases up to 30 μgal . The gravity changes would be related to the Kita-Izu fault system. The tectonic implication will be discussed in Section 4.

3. Fault Model

A quantitative approach to model gravity changes due to seismic activity is given by Okubo (1991, 1992). He derived analytical expressions for gravity changes caused by faulting and inflation sources. In his theory, gravity changes arise from three sources, namely, density changes due to elastic deformation, surface uplift or subsidence, and the addition of mass. To reveal mass movement, we should first find fault parameters to calculate the other two sources. Then we subtract their contributions from observed gravity changes to find masses of added materials.

3.1 Determination of fault parameters

First, we apply Okada (1985)'s elastic dislocation theory to the ground displacements around Ito associated with the March 1997 activity to determine the fault geometry. Similar methods were used by Okada and Yamamoto (1991a,b), Tada and Hashimoto (1991), and Okubo *et al.* (1991) to simulate the crustal displacements associated with the 1989 activity. We use GPS observations, leveling, and EDM data (Geographical Survey Institute, 1997; Tsuneishi, 1997) for the modeling.

The positions of about 10 points near Ito have been continuously monitored by the nationwide GSI's GPS network (Fig. 3). The measurements pose strong constraints on horizontal movement. Figure 3 shows horizontal displacements associated with the earthquake swarm. The stations to the south of the swarm epicenters moved to the SW direction and those to the north moved to the SE and NE directions. In particular, the Komuro-yama station, which is the nearest to the epicenters, moved to the SW direction by as much as 10 cm.

The leveling survey was carried out in October–November

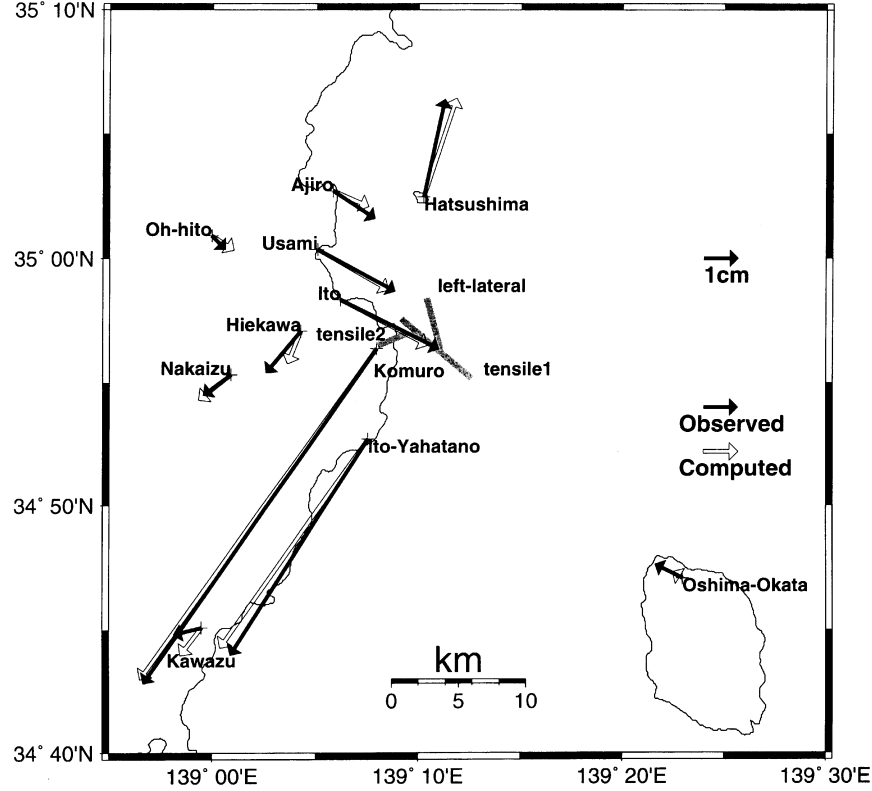


Fig. 3. Horizontal displacements measured by the nationwide GPS network of Geographical Survey Institute (1997). The solid arrows represent the observational results, and the open arrows represent the computed displacements of our elastic deformation model. The thick gray lines indicate the position of the faults determined in our modeling.

1996, after the October 1996 earthquake swarm, and in March 1997, after the March 1997 activity (Fig. 4). Prominent uplifts up to 3 cm were found near the center of the Ito City.

The EDM network has been surveyed every day at several sites around Ito since 1989 (Tsuneishi, 1991, 1997). During the March swarm, it revealed a fine structure of the crustal movement (Fig. 5).

We use the data described above to model the ground deformation caused by three faults buried in an elastic half-space. The geometrical fault parameters (Table 3) are determined in two steps. First we use the GPS and leveling data to model large scale (~ 10 km) ground deformation, which we attribute to a strike-slip fault and a tensile fault. Second we use the EDM data to model small scale (~ 1 km) ground deformation near the Ito City, which we ascribe to another tensile fault.

In the first step, we model the large-scale ground deformation with two faults. We put the first strike-slip fault near the center of the region of swarm epicenters because most of the earthquakes have strike-slip type mechanisms. We further put the second tensile fault near the swarm region. The fault is required to explain large horizontal movements away from the swarm region, which are evident at the Komuro-yama GPS station, for example. To obtain the fault parameters, we minimize the variance given by

$$V_1 = \sum_{\text{GPS}} \frac{(u_x^{\text{obs}} - u_x^{\text{calc}})^2 + (u_y^{\text{obs}} - u_y^{\text{calc}})^2}{\sigma_{\text{GPS}}^2}$$

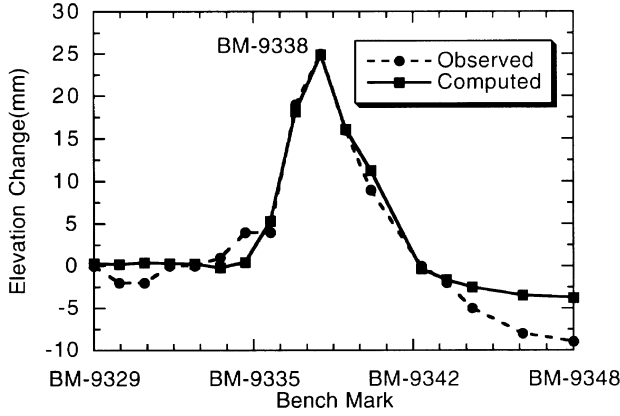
$$+ \sum_{\text{leveling}} \frac{(u_z^{\text{obs}} - u_z^{\text{calc}})^2}{\sigma_{\text{leveling}}^2}, \quad (1)$$

where (u_x, u_y, u_z) are the displacements in the east, north, vertical direction, respectively, and σ is their measurement errors. The errors σ_{GPS} and σ_{leveling} are set to 6 mm and 5 mm, respectively. The superscripts “obs” and “calc” signify the observed and calculated values, respectively. The Poisson ratio of 1/4 is used throughout the calculation. We search the optimum fault parameters by calculating the variance for all the grid points in the range shown in Table 3 so that the global minimum should be attained. The optimum parameters will serve as an initial guess in the following step.

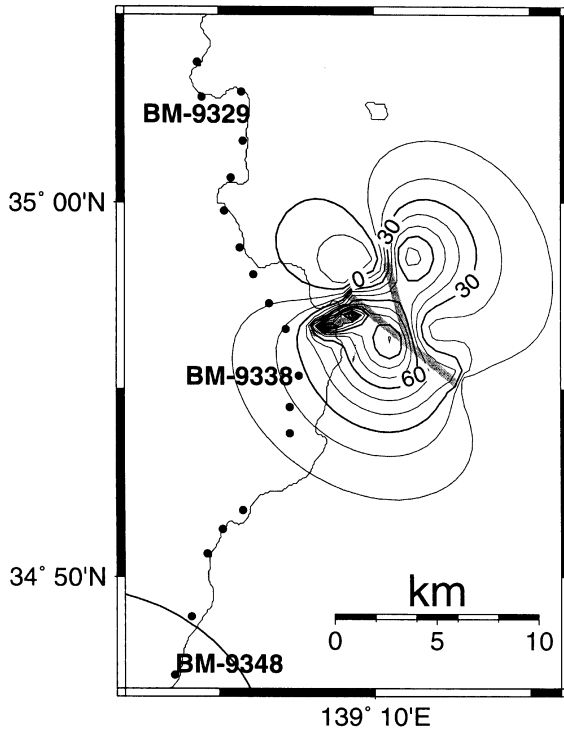
In the second step, we put a small tensile fault near the western edge of the tensile fault introduced in the first step. The fault is required to explain the EDM results. We minimize the variance given by

$$V_2 = \sum_{\text{GPS}} \frac{(u_x^{\text{obs}} - u_x^{\text{calc}})^2 + (u_y^{\text{obs}} - u_y^{\text{calc}})^2}{\sigma_{\text{GPS}}^2} + \sum_{\text{leveling}} \frac{(u_z^{\text{obs}} - u_z^{\text{calc}})^2}{\sigma_{\text{leveling}}^2} + \sum_{\text{EDM}} \frac{(\Delta l^{\text{obs}} - \Delta l^{\text{calc}})^2}{\sigma_{\text{EDM}}^2}, \quad (2)$$

where Δl is the distance change between the EDM sites. We assign 1 cm to σ_{EDM} . The parameters of the first two faults are also slightly adjusted in this step.



(a)



(b)

Fig. 4. Vertical displacements measured by the leveling of Geographical Survey Institute (1997). (a) Comparison between the observations and the model results. The solid circles show the observed elevation changes of benchmarks. The solid squares show the results of our model calculation. (b) Contour map of the vertical displacements derived from our model. The unit is mm. The thick gray lines indicate the position of the faults determined in our modeling.

Table 3 shows the optimum fault parameters. Notice that the optimum dip angles are 90° (vertical) for all the faults. The verticality signifies that the prevalent differential stress is horizontal in this area.

Figures 3–5 show the calculated ground displacements together with the locations of the faults. The observations are reproduced well by the fault model.

We present in Fig. 6 the variance V_2 as a function of important fault parameters. The variance shows a steep minimum

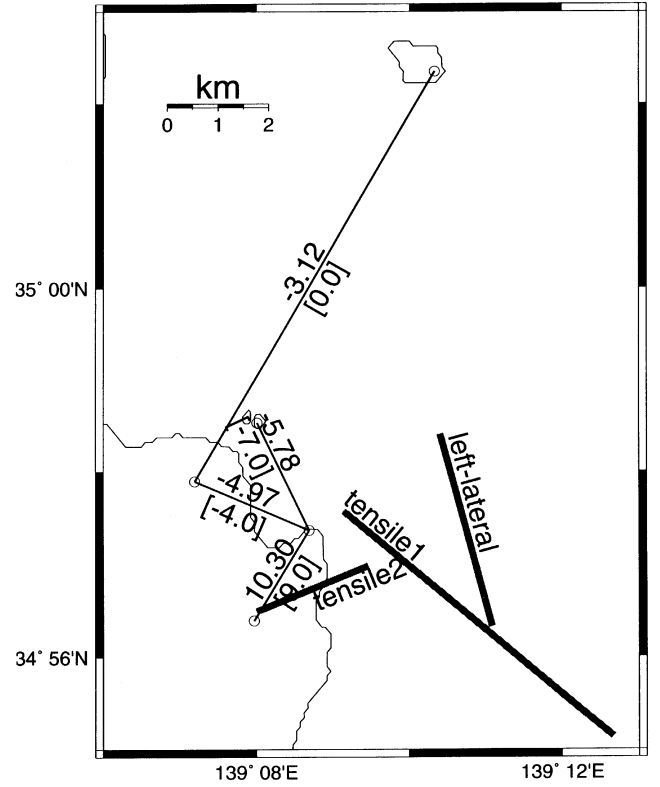


Fig. 5. Distance changes between several points near Ito measured by Tsuneishi (1997). The lines show the measurement lines. The numbers on the line show the results of our model. The numbers in square brackets show the observations. The unit is cm. The thick gray lines indicate the position of the faults in our model.

as a function of the parameters. This signifies that the fault geometry and displacements are tightly constrained. The minimum value of the variance is 22.4, which we expect to be 20 for the χ^2 distribution. The two values agree within the expected standard deviation of $2 \times 20 = 40$. The expected variance 20 equals the degree of freedom. It is given by the number of the data 44 (18 leveling data, 4 EDM data, and 11×2 GPS data) minus the number of the fault parameters 24 (8 parameters for each of the three faults).

3.2 Determination of the density of fault-filling matter

Next, we use the gravity data and the fault parameters obtained above to estimate the density of the material which would have filled the tensile faults. We minimize the variance given by

$$V_3 = \sum_{\text{gravity}} \frac{(\Delta g^{\text{obs}} - \Delta g^{\text{calc}})^2}{\sigma_{\text{gravity}}^2}, \quad (3)$$

where Δg is the gravity change. The gravity measurement error σ_{gravity} is set to $7 \mu\text{gal}$. We did not use the data denoted by asterisks in Table 2 because it is evident that they are not associated with the movement of the faults. The value of 2.67 g/cm^3 is used for the density of the surrounding medium in the calculation.

Figure 7 shows the change of the variance V_3 as a function of the densities which would have filled the two tensile faults. For the larger fault, the density is likely to be below 1.5 g/cm^3 . For the smaller fault, the density cannot be constrained. The

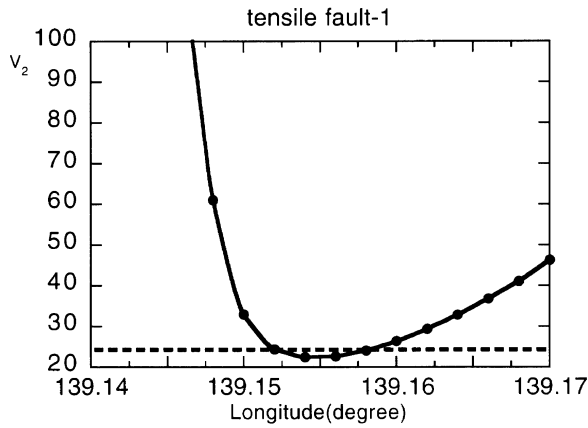
Table 3. Optimum parameters for the three faults.

Strike-slip fault (left lateral)		
	Grid search range	Optimum value
Location of the western edge	139.160°E–139.205°E	139.175°E
	34.965°N–34.985°N	34.970°N
Strike angle	N160°E–N180°E	N165°E
Length	3.0–8.0 km	4.0 km
Width	3.0–6.0 km	4.0 km
Dip angle	70°–100°	90°
Depth of the upper edge	1.0–4.0 km	1.0 km
Left-lateral dislocation	0.5–0.9 m	0.8 m
Tensile fault-1 (large)		
	Grid search range	Optimum value
Location of the western edge	139.070°E–139.210°E	139.154°E
	34.920°N–35.000°N	34.956°N
Strike angle	N100°E–N170°E	N130°E
Length	3.0–11.0 km	7.0 km
Width	2.0–7.0 km	6.0 km
Dip angle	60°–110°	90°
Depth of the upper edge	0.0–6.0 km	0.0 km
Tensile opening	0.2–0.7 m	0.3 m
Density of the material which filled the cavity	0.0–3.63 g/cm ³	0.0 g/cm ³
Tensile fault-2 (small)		
	Grid search range	Optimum value
Location of the western edge	139.120°E–139.150°E	139.135°E
	34.935°N–34.960°N	34.938°N
Strike angle	N60°E–N80°E	N68°E
Length	1.0–4.0 km	2.4 km
Width	0.2–4.0 km	0.4 km
Dip angle	60°–120°	90°
Depth of the upper edge	0.0–4.0 km	0.0 km
Tensile opening	0.1–0.8 m	0.35 m
Density of the material which filled the cavity	0.0–3.63 g/cm ³	not constrained

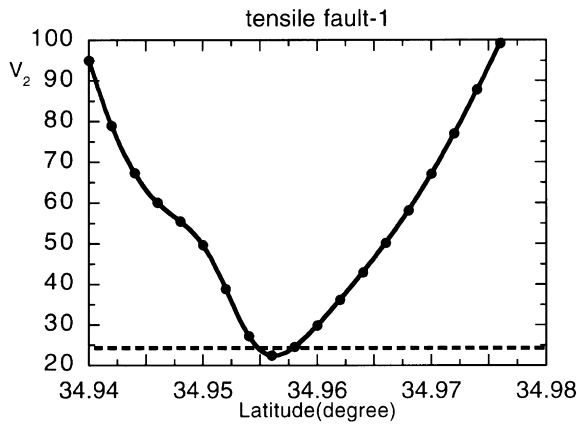
smallness of the density in the larger fault implies that it would have contained large amount of gas or water. The minimum value of the variance is 10.8, while its expected value is 20, which is given by the number of the data 22 minus the number of the density parameters 2. The calculated variance is about half the expected one, but the difference is within the expected standard deviation of 40. The smallness

of the calculated value may signify that the estimated error of 7 μ gal is an overestimate.

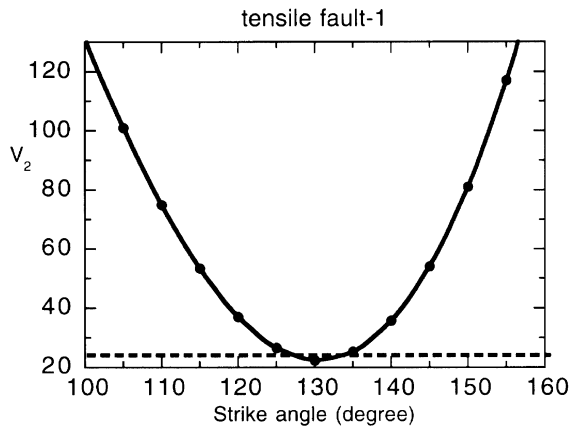
Previous analyses of gravity changes related to fissure eruptions also revealed the intrusion of low-density matters. Okubo and Watanabe (1989) analyzed the gravity changes associated with the 1986 fissure eruption of the 1986 Izu-Oshima Volcano to find that the material which filled the



(a)

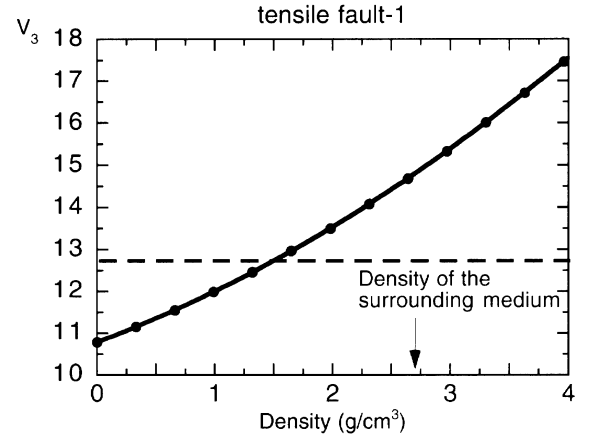


(b)

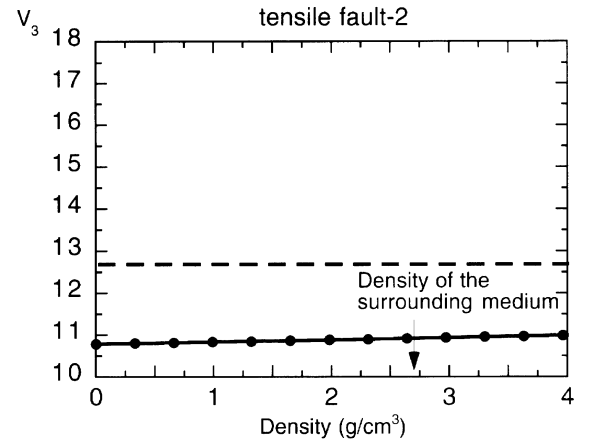


(c)

Fig. 6. Variance V_2 as a function of several important parameter values. The minimum value V_{\min} of the variance is 22.4. The broken line indicates the value $V_{\min} + 2$, for which the likelihood is reduced by e^{-1} . (a) Variance against the longitude of the western edge of the tensile fault 1. (b) Variance against the latitude of the western edge of the tensile fault 1. (c) Variance against the strike angle of the tensile fault 1.



(a)



(b)

Fig. 7. Variance V_3 as a function of the densities of the materials which filled the two tensile faults. The minimum value V_{\min} of the variance is 10.8. The broken line indicates the value $V_{\min} + 2$, for which the likelihood is reduced by e^{-1} . (a) Variance against the density of the material which filled the tensile fault 1. It shows that the density is below 1.5 g/cm³. (b) Variance against the density of the material which filled the tensile fault 2. It shows that the density cannot be constrained.

fissure zone had a density of 0.8 ± 0.5 g/cm³. Okubo *et al.* (1991) showed that the matter which filled the tensile crack related to the 1989 Teishi-Knoll Eruption had almost zero density. Thus the intrusion of gas or water into fissures would be a common phenomenon in volcanic activities.

Figure 8 shows the contour map of calculated gravity change with the densities of the intruding material set to zero. It displays that the gravity changes in the region 1 results only from the motion of the faults obtained from GPS, leveling, and EDM data. The gravity changes in the region 2 and 3 require other explanations.

4. Discussion

We shall discuss the characteristics of the crustal movements and gravity changes associated with the March earthquake swarm in relation to the tectonic environment. Our discussion on the tectonics is based on the tectonic model of

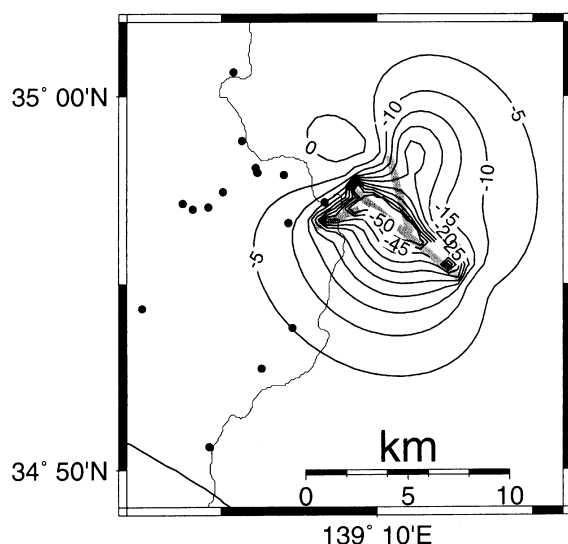


Fig. 8. Contour map of the calculated gravity change associated with the fault motion. The unit is μgal . The densities of the materials which filled the two tensile faults are set to zero. The thick gray lines indicate the position of the faults.

Koyama and Umino (1991) and Koyama (1993).

The crustal displacements around Ito are explained well by the three-fault model: a strike-slip fault and two tensile faults. The principal part is due to the larger tensile fault that extends to the NW-SE direction. This type of fault is often invoked to explain the crustal movement of this area associated with earthquake swarms (Okada and Yamamoto, 1991a,b; Okubo *et al.*, 1991; Tada and Hashimoto, 1991), and would be related to volcanic activities. It is also consistent with the NW-SE vent alignments of the Higashi-Izu monogenetic volcanos (Aramaki and Hamuro, 1977; Koyama and Umino, 1991; Koyama *et al.*, 1995). This alignment of the fissures is interpreted as a result of NW-SE compressive and NE-SW tensile stress field of this region. The stress field itself can be understood as the result of the collision of the Philippine Sea plate with the Japan Island and the release of stress by the West-Sagami-Bay Fracture (Koyama and Umino, 1991).

In Section 2, we have identified three regions where gravities changed significantly. We shall discuss the characteristics of each region in the following.

The three-fault model explains the gravity changes in the region 1 very well, and reveals that the density of the material that filled the larger tensile fault is very small. The smallness implies that it contained large amount of exsolved gas or water. Such a low density is consistent with the shallowness of the fault, because most of the water would exsolve from the magma at depths shallower than 1 km (e.g. Woods, 1995). The strength of the uppermost crust, in this case, prevented the highly vesiculated magma from erupting.

The complex pattern of the gravity changes in the region 2 may indicate the intrusion of magma or fluid to the shallow part beneath Ito. The shallow intrusion is also suggested by the existence of the smaller tensile fault in our three-fault model. The fault is very shallow and extends to the land region.

On the other hand, the gravity decrease in the region 3 remains puzzling. Lack of other observations in this area prohibits detailed modeling. Similar decreases were observed in 1994 and 1995. They can be associated with the earthquake swarms of March 1994 and October 1995, respectively. Hence, the gravity decrease in the mountain area of the Izu Peninsula would commonly accompany earthquake swarms near Ito. An explanation for the gravity decrease may be constructed based on the block rotation model of Koyama and Umino (1991). They proposed that the northeastern region of the Izu Peninsula is composed of slivers which are rotated counterclockwise due to the left-lateral shearing motion. They assigned the western boundary of the slivers to the Kita-Izu Fault system, which extends to the center of the gravity decrease (Fig. 1). As the slivers move, a large deformation in the deeper parts of the Kita-Izu Fault system is expected. This may have given rise to the observed gravity decrease.

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