## MIXING MODELS, COLORS AND THERMAL EMISSIONS

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Abstract. Color and spectral observations are important for understanding the taxonomy, composition, formation conditions and evolutionary history of Trans-Neptunian Objects (TNOs). Colors or spectra do not, however, uniquely constrain the compositions of these objects, as we demonstrate using simple spectral mixing models. We show that thermal flux observations at 70  $\mu$ m (or any wavelength near or beyond the peak of the object's emission) are much less sensitive to details of how the materials are combined. Coupling reflectance and thermal observations of TNOs will help remove much of the ambiguity intrinsic to color and spectral data.

## 1. Introduction

Trans-Neptunian Objects (TNOs) are thought to be composed of relatively pristine material from the outer reaches of the proto-solar nebula and to be the source population for Jupiter family comets and Centaur objects (e.g., Levison and Duncan, 1997). Nearly 1000 have been discovered since 1992 and more are found each year, revealing patterns in the distribution of TNO orbits (e.g., Jewitt et al., 1998; Millis et al., 2002; Buie et al., these proceedings). Despite their importance for understanding solar system formation, we still know relatively little about TNO physical and chemical characteristics. To advance our understanding, data are needed which constrain their compositions, sizes, shapes, and albedos.

A few of the brightest TNOs have been observed spectroscopically, giving some compositional information (e.g., Jewitt and Luu, 2001; Cruikshank and Dalle Ore, these proceedings). Even fewer have been observed in the thermal infrared (e.g., Thomas et al., 2000; Jewitt et al., 2001), yielding estimates of sizes, albedos, and surface temperatures. Observations of broad-band colors are more abundant, and point to a remarkably diverse population (e.g., Tegler and Romanishin, 1998, 2000, 2003; Delsanti et al., 2001; Doressoundiram et al., 2001; Gil-Hutton and Licandro, 2001; numerous papers in these proceedings) but these colors do not provide unambiguous information about compositions.

The Spitzer Space Telescope (SST) will be capable of observing many TNOs in the thermal infrared, primarily at 70  $\mu$ m. The simple models we present here indicate that the SST data will not only be important for determining albedos and



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### 2. Models

Spectral absorbers (species having strongly wavelength-dependent absorption characteristics, such as ices and tholins) are thought to be constituents of TNOs (e.g., Jewitt and Luu, 2001). They may be accompanied by continuum absorbers (or scatterers) which can darken (or brighten) a surface, and can reduce (or sometimes enhance) the spectral contrast of spectral absorbers. Devolatilized carbonaceous materials are possible darkening continuum absorbers, while ices can act as bright continuum scatterers at wavelengths lacking absorption bands. Remarkably small weight percents of dark continuum absorbers can efficiently mask the spectral signatures of brighter materials, if they are intimately mixed (e.g., Clark, 1981; Clark and Lucey, 1984), while small quantities of bright, scatterers are geographically isolated. A bright species, such as an ice, could conceivably dominate an object's reflectance spectrum while only covering a minute fraction of its surface.

We used Hapke theory (e.g., Hapke, 1981, 1993) to calculate albedos and spectral contrasts for various mixtures between representative spectral absorbers (water ice, Titan tholin), a continuum absorber (charcoal), and a continuum scatterer (water ice). We define spectral contrast as the color difference in magnitudes between two wavelengths, relative to solar colors. Figure 1 illustrates the dependence of spectral contrast on composition for intimate and patchwork mixtures of water ice/charcoal (Figure 1a), tholin/charcoal (Figure 1b), and tholin/ice (Figure 1c). In these models "charcoal" is a proxy for any dark continuum absorber on TNOs, such as devolatilized carbonaceous material, and is simulated with isotropically scattering 10  $\mu$ m grains having 9% single scattering albedo (the geometric albedo is much lower than the single-scattering albedo due to multiple scattering). Ice and tholin grains are also assumed to have 10  $\mu$ m mean diameters, with optical constants for H<sub>2</sub>O ice I<sub>h</sub> taken from Grundy and Schmitt (1998), and Titan tholin optical constants taken from Khare et al. (1994). The exact shapes of these curves depend on our assumptions about grain size and shape, as well on as numerous additional assumptions inherent in the Hapke theory. While the curves in Figure 1 are not unique, they are useful for exploring typical optical behavior of mixtures of possible TNO constituents.

The mixing models show that color and albedo tend to be highly nonlinear functions of composition, and that different styles of mixing lead to very different dependence on composition. Intimate mixtures of a bright ice and a dark continuum absorber exhibit low spectral contrast until the ice is nearly pure, while patchwork mixtures show strong spectral contrast unless there is almost no ice at all (Figure 1a). Mixtures of two dark materials (Figure 1b) tend to be strongly colored if there is any significant fraction of the spectral absorber, for both intimate and patchwork mixtures. Even more remarkable behavior is seen when materials are combined within individual grains, such as when tholin molecules are introduced into the interior of water ice grains, as in Figure 1c. A very small quantity of tholin dramatically reddens the ice. Paradoxically, adding ice to tholin also reddens the mixture, because it reduces the opacity of individual particles, increasing their single scattering albedos, most notably in the red. This trend continues until so much ice has been added that the opacity declines at the more strongly-absorbing wavelength, as well.

For all of these mixtures, the surface composition inferred from observed spectral contrast can be strongly influenced by the details of how the materials are mixed. This sensitivity illustrates the uncertainties inherent in interpreting colors and/or spectra without further information. Knowledge of the albedo of a surface can remove much of this uncertainty, as exemplified by Figure 1a, where simultaneously knowing the K2 albedo and K1-K2 spectral contrast would likely place an object on one or the other of the two mixing lines and constrain the composition nicely. Figure 2 shows predicted thermal fluxes at 24 and 70  $\mu$ m from a 100 km radius TNO covered by mixtures of charcoal and H2O ice, located 40 AU from the Sun. The curves for patchwork and intimate mixtures are nearly coincident, especially at 70  $\mu$ m. Results for both the Standard Thermal Model (STM) and the Isothermal Latitude Model (ILM) are shown, with the STM giving higher fluxes. The flux uncertainty associated with the pole position is about 30% at 70  $\mu$ m with a corresponding albedo uncertainty of about 20%. For objects where the pole position can be constrained (binaries, TNOs with strong lightcurves), it should be possible to measure albedos accurate to better than 10%. If we assume SST produces a catalog of large-TNO albedos accurate to 20% including all uncertainties, then Figure 1a indicates that compositional uncertainties can be determined to a similar accuracy, along with distinguishing intimate from patchwork mixtures.

# 3. Conclusions

The diverse and nonlinear effects of various different mixtures of materials present a daunting challenge for interpreting color and spectral observations of TNOs in terms of their compositions. However, with an albedo constraint, the ambiguities can be significantly reduced. For this reason, SST will play an extremely valuable role in advancing understanding of TNO compositions and physical characteristics, when combined with data from ongoing visible and near-infrared observing campaigns.



*Figure 1.* Mixing curves between possible TNO surface components. End-members are marked with black diamonds and intermediate compositions are indicated by a large dot at 50% composition and smaller dots at 10% intervals. In panel (a), K<sub>1</sub> and K<sub>2</sub> are custom filters: K<sub>1</sub> is centered on the 2  $\mu$ m water ice absorption band and K<sub>2</sub> on the continuum region around 2.25  $\mu$ m. In )c), intimate and patchwork models are not labeled, but are represented as in other panels, while the dotted curve represents a molecular mixture.



*Figure 2.* Thermal emission from a hypothetical 100 km TNO at 40 AU for models with intimate (gray curves) and patchwork (dashed curves) mixtures of charcoal and H<sub>2</sub>O ice, as in Figure 1. Curves are shown for the Standard Thermal Model (STM, pole-on orientation) and Isothermal Latitude Model (ILM, equator-on orientation). The STM produces higher fluxes than the ILM. The SST sensitivities at 24 and 70  $\mu$ m are about 0.15 and 1 mJy, 3 $\sigma$  in 500 seconds.

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