

Research Article

Selecting the best probability distribution for at-site flood frequency analysis; a study of Torne River



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Abstract

At-site flood frequency analysis is a direct method of estimation of flood frequency at a particular site. The appropriate selection of probability distribution and a parameter estimation method are important for at-site flood frequency analysis. Generalized extreme value, three-parameter log-normal, generalized logistic, Pearson type-III and Gumbel distributions have been considered to describe the annual maximum steam flow at five gauging sites of Torne River in Sweden. To estimate the parameters of distributions, maximum likelihood estimation and L-moments methods are used. The performance of these distributions is assessed based on goodness-of-fit tests and accuracy measures. At most sites, the best-fitted distributions are with LM estimation method. Finally, the most suitable distribution at each site is used to predict the maximum flood magnitude for different return periods.

Keywords Flood frequency analysis · L-moments · Maximum likelihood estimation

1 Introduction

Floods are natural hazards and cause extreme damages throughout the world. The main reasons of floods are extreme rainfall, ice and snow melting, dam breakage and the lack of capacity of the river watercourse to convey the excess water. Floods are natural phenomena which cause disasters like destruction of infrastructure, damages in environmental and agricultural lands, mortality and economic losses. Many frequency distribution models have been developed for determination of hydraulic frequency, but none of the distribution models is accepted as a universal distribution to describe the flood frequency at any gauging site. The selection of a suitable distribution usually depends on the characteristics of available data at a particular site. We need to estimate the flood magnitude at a particular site for various purposes including construction of hydraulic structures (barrages, canals, bridges, dams, embankments, reservoirs and spillways), insurance studies, planning of flood management and rescue operations. We come across a number of methods which are available in the literature for flood magnitude estimation, but at-site flood frequency analysis remains the most direct method of estimation of flood frequency at a particular site.

To describe the flood frequency at a particular site, the choice of an appropriate probability distribution and parameter estimation method are of immense importance. The probability distributions used in this study include the generalized extreme value (GEV) distribution, Pearson type-III (P3) distribution, generalized logistic (GLO) distribution, Gumbel (GUM) distribution and three-parameter log-normal (LN3) distribution. These distributions are recommended for at-sight flood frequency analysis in various countries [2, 4, 24]. Furthermore, these distributions are most commonly traced in the hydrological literature for at-site and regional flood frequency analysis.

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Cicioni et al. [3] conducted at-site flood frequency analysis in Italy by using 107 stations. They identified LN3 and GEV as best-fitting distributions based on the Kolmogorov–Smirnov (KS), Anderson–Darling (AD) and Cramer–von Mises (CVM) goodness-of-fit tests. Saf [23] found the GLO as the most suitable distribution for Upper-West Mediterranean subregion in Turkey. Mkhandi et al. [17] used annual maximum flood data of 407 stations from 11 countries of Southern Africa to conduct the regional frequency analysis. They identified Pearson type-III (P3) distribution with probability weighted moment (PWM) method and log-Pearson type-III (LP3) with a method of moment (MOM) as the most suitable distributions for the regions. Młyński et al. [18] identified log-normal distribution as the most suitable for the upper Vistula Basin region (Poland). There have been many studies in the past literature on the comparison of various probability distributions with different parameter estimation methods for at-site flood frequency analysis, e.g. see [1, 5, 7, 22] (Fig. 1).

The most commonly used methods for estimation of parameters in flood frequency analysis are the maximum likelihood estimation (MLE) method, the method of moments (MOM), the L-moments (LM) method and the probability weighted moments method (PWM). The MLE method is an efficient and most widely used method for estimation of parameters. Recently, the LM method has gained more attention in the hydrological literature for estimation of parameters of probability distributions. In this research study, we used LM and MLE methods for estimation of parameters of the candidate probability distribution.

The methods usually use for selection of the best distribution are goodness-of-fit (GOF) tests (e.g. Anderson-Darling and Cramér-von Mises), accuracy measures (e.g. root mean square error and root mean squared percentage error), goodness-of-fit indices (e.g. AIC and BIC) and graphical methods (e.g. Q-Q plot and L-moment ratio diagram). In the hydrological literature, researchers have used different methods in order to find the best probability distribution. To identify the best-suited distribution at each site of Torne River, we have used goodness-of-fit (GOF) test and accuracy measures. The GOF tests are used to test that the data come from a specific distribution. The accuracy measures provide a term by term comparison of the deviation between the hypothetical distribution and the empirical distribution of the data. The details about accuracy measures and goodness-of-fit test used in this study are described in Sect. 3.

The estimation of flood frequency of the high return period is of great interest in flood frequency analysis. The flood frequency estimation of return periods is always associated with uncertainties. Uncertainty in flood frequency analysis arises from many sources. Uncertainties included in water resources management can be distinguished in data uncertainties, structural uncertainties and model/parameters uncertainties, see e.g. [13, 14]. Furthermore, there is uncertainty in the estimation of flood frequency of return periods much larger than the actual records, particularly in the type of probability density function (PDF) and its parameters. This is particularly true on the right tail of the PDF, the region of interest for flooding. In addition, there is uncertainty in the measurements. For example, see [15] for an in-depth discussion on epistemic uncertainty (reducible uncertainty) and natural uncertainty (irreducible uncertainty). The flood estimation on high return periods are always associated with high uncertainties. In this study, we quantify the uncertainty of a given quantile estimate for a specific fitted distribution by using the parametric bootstrap method.

In this research paper, the flood frequency calculation, using statistical distribution, is addressed for gauged catchments, for which we dispose a respectively longterm hydrological time series. The choice of an appropriate probability distribution and associated parameter estimation method is vital for at-site flood frequency analysis. The core objective of this study is to find the best-fit distribution among the candidate probability distributions with a particular method of estimation (MLE or LM) for annual maximum peak flow data at each site of the Torne River by using goodness-of-fit (GOF) tests and accuracy measures. We are also interested to look that, is there any best overall distribution and fitting method for these five sites of Torne River?. Finally, to estimate the quantiles of flood magnitude for the return period of 5, 10, 25, 50, 100, 200 and 500 years with non-exceedance probability at each site of the river using the best-fit probability distribution. To address the uncertainty of flood estimations, we estimate standard error of estimated quantiles and construct 95% confidence interval of flood quantile for the return period using the parametric bootstrap method. This is a first study for at-site flood frequency analysis of Torne River.

This paper is organized as follows: Sect. 2 describes the study area and available data for the analysis. Section 3 deals with the model description, parameter estimation method and model comparison methods. Section 4 provides the results and discussion of the application of L-moments and maximum likelihood estimation method based flood frequency analysis of five gauging sites on the Torne River. Finally, Sect. 5 concludes the article.

2 Study area and data

The Torne River works as a border between northern Sweden and Finland, with total catchment area 40157 km² of which 60% is within Swedish border and the remaining

Table 1 Summary of Torne River gauging sites

Station name	Station no.	Latitude	Longitude	Catchment area (km²)	Period for time series
Kukkolankoski Övre	16722	65.98	24.06	33,929.60	1911–2018
Pajala Pumphus	2012	67.21	23.40	11,038.10	1969–2018
Abisko	2357	68.19	19.99	3345.50	1984–2018
Junosuando	04	67.43	22.55	4348.00	1967–2018
Övre Abiskojokk	957	68.36	18.78	566.30	1985–2018

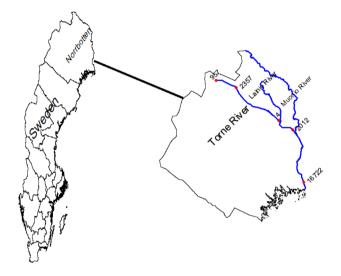


Fig. 1 Locations of gauging stations used in this study

area is in Finland. The Muonio River, which is the biggest contributor of the Torne River, joins shortly after Pajala Pumphus. Another contributor river Lainio (259.74 km long) joins the Torne river shortly after Junosuando. In springtime, water flow is above average level, which converts into flood and this flood causes the damages to the waterfront constructions and buildings [6]. Therefore, Torne River is frequently affected by flooding problem

[Swedish meteorological and hydrological institute (SMHI)]. The data of annual maximum flow of five gauging sites of Torne River (Swedish: Torneälven) are considered in this study. The data have been collected from SMHI (www. smhi.se). The length of the data series varies from 34 to 108 years. The summary of Torne River gauging sites characteristics is presented in Table 1.

3 Methodology

3.1 Candidate probability distributions

To describe the flood frequency at a particular site, the selection of an appropriate probability distribution is always important. We have considered generalized extreme value (GEV), Pearson type-III (P3) distribution, generalized logistic (GLO) distribution, Gumbel (GUM) distribution and three-parameter log-normal (LN3) distribution for the analysis of flood frequency at five gauging sites of the Torne River. The probability density function (pdf) and quantile function y(F) of these distributions are summarized in Table 2. These distributions are common in the literature and are recommended distributions for flood frequency analysis in many countries (see e.g. [2, 4, 21, 22]). We explain the detail of parameter estimation method (MLE and LM) in the following subsection.

Table 2 Probability density and quantiles functions of the probability distributions

Distributions	Probability density function $f(y)$	Quantile function y(F)
GEV	$\frac{1}{\alpha} \left[1 - \kappa \left(\frac{y - \mu}{\alpha} \right) \right]^{\frac{1}{k} - 1} \exp \left\{ - \left[1 - \kappa \left(\frac{y - \mu}{\alpha} \right) \right]^{\frac{1}{k}} \right\}$	$\mu + \frac{\alpha}{\kappa} \Big[1 - (-\log F)^{\kappa} \Big]$
P3	$\frac{1}{\beta^{\alpha}\Gamma^{\alpha}}(y-\mu)^{\alpha-1}\exp\left\{-\frac{(y-\mu)}{\beta}\right\}$	Explicit analytical form is not available
GLO	$\frac{1}{\alpha} \left[1 - \kappa \left(\frac{y - \mu}{\alpha} \right)^{\left(\frac{1}{k} - 1 \right)} \right] \left[1 + \left\{ 1 - \kappa \left(\frac{y - \mu}{\alpha} \right) \right\}^{1/\kappa} \right]^{-2}$	$\mu + \frac{\alpha}{\kappa} \left[1 - \left\{ (1 - F) / F \right\}^{\kappa} \right]$
LN3	$\frac{1}{\alpha\sqrt{2\pi}}\exp\left[-\log\left\{\frac{1-k(y-\mu)}{\alpha}\right\}-\frac{1}{2}\left[-\frac{1}{k}\log\left\{\frac{1-k(y-\mu)}{\alpha}\right\}\right]^2\right]$	Explicit analytical form is not available
GUM	$\frac{1}{\alpha} \exp\left[-\frac{y-\mu}{\alpha} - \exp\left(-\frac{y-\mu}{\alpha}\right)\right]$	$\mu - \alpha \log \left(-\log F \right)$

where μ , α and κ are the location, scale and shape parameters of the distribution

3.2 Maximum likelihood estimation (MLE) method

The MLE method estimates the parameters by maximizing the log-likelihood function of a probability distribution. Suppose we have *n* independent and identically distributed observations y_1, y_2, \dots, y_n . Each y_i has a pdf given by $f(y_i; \boldsymbol{\mu})$. Here, $\boldsymbol{\mu} = (\mu_1, \mu_2, \dots, \mu_k)$ is a vector of unknown parameters to be estimated. Then, the log-likelihood function is defined as $I(\mu) = \sum_{i=1}^{n} \log f(y_i; \mu)$. The maximum likelihood estimate of μ is the value of the parameter vector μ that maximize the $I(\mu)$ for given data Y. We use numerical optimization methods in order to search μ which give the maximum value of $I(\mu)$. Many numerical optimization methods, e.g. Newton-Raphson method, Nelder and Mead, differential evolution, etc. are found in the literature. We have used Nelder and Mead method for numerical optimization proposed by Nelder and Mead [19].

3.3 Theory of L-moments (LM)

L-moments are introduced by Hosking [9, 10], which are linear functions of probability weighted moments (PWM's). L-moments are alternative to the conventional moments, but computed from linear combinations of order statistics. L-moments can be defined for any random variable Y whose mean exists [10]. The rth-order PWM (β_r) is defined as

$$\beta_r = \int_{0}^{1} y(F)F(y)^r dFr = 0, 1, 2, ...$$

where F(y) is a cumulative probability distribution and y(F) is a quantile function of distribution. The first four L-moments in terms of linear combination of PWM are defined as

$$\lambda_{1} = \beta_{0}$$

$$\lambda_{2} = 2\beta_{1} - \beta_{0}$$

$$\lambda_{3} = 6\beta_{2} - 6\beta_{1} + \beta_{0}$$

$$\lambda_{4} = 20\beta_{3} - 30\beta_{2} + 12\beta_{1} - \beta_{0}$$

The first L-moment (λ_1) is a measure of location (mean), while the second L-moment represents the dispersion. Finally, the L-moment ratios defined by Hosking [10] are given below

L-Coefficient of variation
$$(au_2)=rac{\lambda_2}{\lambda_1}$$

L-Skewness $(au_3)=rac{\lambda_3}{\lambda_2}$
L-Kurtosis $(au_4)=rac{\lambda_4}{\lambda_2}$

The unbiased sample estimators of β_i of the first four PWMs for any distribution can be computed as follows

$$b_0 = n^{-1} \sum_{j=1}^{n} y_{j:n}$$

$$b_1 = n^{-1} \sum_{j=2}^{n} \frac{(j-1)}{(n-1)} y_{j:n}$$

$$b_2 = n^{-1} \sum_{j=3}^{n} \frac{(j-1)(j-2)}{(n-1)(n-2)} y_{j:n}$$

$$b_3 = n^{-1} \sum_{i=4}^{n} \frac{(j-1)(j-2)(j-3)}{(n-1)(n-2)(n-3)} y_{j:n}$$

where the data $(y_{1:n})$ are an ordered sample in ascending order from 1 to n. The parameters with L-moments estimation method are obtained by equating the sample L-moments with distribution L-moments.

3.4 Standard error of estimated parameters

The standard errors (SE) of estimated parameters indicate a measure of reliability of estimates and performance of estimation technique. In this study, we have obtained SE of estimated parameters by Monte Carlo simulation method. The description of this method is given as

- We use estimated parameters with MLE and LM method at each gauging site and draw 1000 samples of size equal to the length of data from each probability distribution.
- For each simulated sample, we obtain the MLE and LM estimates for the parameters of the distribution.
- For each gauging site, the standard errors are obtained by taking the standard deviation of these 1000 MLE and LM estimates of the parameters of each distribution.

3.5 Goodness-of-fit (GOF) tests

The goodness-of-fit tests are used to test that the observed data follow a particular distribution. We consider the Anderson–Darling (AD) test for the study. This test is often used in flood frequency analysis and has shown good performance in case of small sample size and heavy-tailed distributions [12, 20]. The test statistic for AD test is defined as

AD =
$$-n - S$$

where $\sum_{i=1}^{n} \left[\frac{2i-1}{n} \left(\log(1 - F(y_{n-i+1})) + \log(F(y_i)) \right) \right]$
where $F(y_i)$ represents the cumulative distribution function (CDF) of the specified distribution.

3.6 Accuracy measure method

In accuracy measure (AM) methods, we have used the mean absolute error (MAE), mean absolute percentage error (MAPE), root mean square error (RMSE), root mean squared percentage error (RMSPE) and correlation coefficient (R^2) to evaluate how adequately a given distribution fits the observed data. These measures are defined as

$$\begin{aligned} \text{MAE} &= \frac{1}{n} \sum_{i=1}^{n} |F(y_i) - F(\hat{y}_i)| \\ \text{MAPE} &= \frac{100}{n} \sum_{i=1}^{n} \left| \frac{(F(y_i) - F(\hat{y}_i))}{F(y_i)} \right| \\ \text{RMSE} &= \sqrt{\frac{\sum_{i=1}^{n} (F(y_i) - F(\hat{y}_i))^2}{n}} \\ \text{RMSPE} &= \sqrt{\frac{1}{n} \sum_{i=1}^{n} \left(\frac{F(y_i) - F(\hat{y}_i)}{F(y_i)} \right)^2} \times 100 \\ R^2 &= \frac{\sum_{i=1}^{n} (F(\hat{y}_i) - \bar{F}(y_i))^2}{\sum_{i=1}^{n} (F(\hat{y}_i) - \bar{F}(y_i))^2 + \sum_{i=1}^{n} (F(y_i) - F(\hat{y}_i))^2} \end{aligned}$$

where $\bar{F}(y_i) = \frac{\sum_{i=1}^n F(\hat{y}_i)}{n}$ and n represents the size of the data series. In all above accuracy measures, $F(y_i)$ is the empirical cumulative distribution function (CDF) of the data (observed ordered values) and $F(\hat{y}_i)$ indicates the theoretical CDF of the distribution (ordered estimated values from the distribution).

3.7 Quantile estimation

After selection of the best probability distribution, the main goal of flood frequency analysis is to estimate the quantile y_T for a return period (T) of scientific relevance. $P(Y \ge y_T) = \frac{1}{T}$ indicates the probability of exceedance from flood level y_T once in T years. The cumulative probability of non-exceedance is defined as

$$F = F(y_T) = P(Y \le y_T) = 1 - P(Y \ge y_T) = 1 - \frac{1}{T}$$

The distribution function $F(y_T)$ can be expressed in inverse form as $y_T = y(F)$, and we can directly evaluate estimated quantile y_T by replacing F. Sometimes, inverse of $F(y_T)$ does not exist analytically, and then, the numerical method is used to evaluate y_T for the given value of F. The expressions of quantile function of candidate distributions are summarized in Table 2. The quantile estimate for T years is calculated by substituting the value of $F = (1 - \frac{1}{T})$ in the expressions of quantile in Table 2. The standard error of estimated quantiles represents the uncertainty in the

estimation of flood frequency of return periods. The confidence interval of flood quantile gives an estimated range of values which is likely to include the flood frequency of return periods. We use a parametric bootstrap method for estimation of standard error of estimated quantiles and confidence interval of flood quantiles of return periods. This method is more precise than an asymptotic computation when *n* is small [16, p. 133]. The detail of procedure involves in parametric bootstrap method is given in [16, p. 133].

4 Result and discussion

We summarized the basic statistics of five gauging sites in Table 5. All data on gauging sites in the table are in cubic metre per second. It is observed that all data at these sites are skewed. This is a enough evidence to model the data with non-normal distribution. In flood frequency analysis, the basic statistical assumptions are independence, randomness and stationarity of the data series (see e.g. [8, 11]). The independence and randomness of the data series at given site are tested by using correlation coefficient (r) at lag-1 and Wald-Wolfowitz (WW) test, respectively. To check the stationarity of the data series, Mann-Kendall (MK) test has been applied. The assumptions verification results are summarized in Table 3. The results in Table 3 indicate that the data series at each gauging site of Torne River are suitable for flood frequency analysis and probability density estimation.

The estimated parameters for each distribution at each gauging site by using MLE and LM method of estimation along with standard error (SE) are reported in Table 4. To identify the best distribution at each site, we use GOF tests and accuracy measures. Each distribution with parameter estimation method is ranked in each GOF test and accuracy measure in Table 6. The distribution is assigned a rank score between 1 and 10 in GOF test and accuracy measures, rank score 10 for the best-fitted and 1 for the worse fitted distribution. The rank score scheme is based on the relative magnitude of accuracy measures and AD test P value. The distribution with the lowest RMSE, lowest MAE, lowest RMSEP, lowest MAEP or the highest R² has the highest rank score value 10. In AD test, the distribution with the highest P value has the highest rank score value 10. The best distribution with estimation method at each site is identified based on the total rank score in GOF tests and accuracy measures methods. The total rank score in Table 6 indicates that GLO with MLE estimation method is best for Junosuando. For site Pajala Pumphus and Abisko, the LN3 distribution is performed better than other distributions with the LM method. The GEV and PE3 distribution with

Table 3 Assumptions results of five gauging sites of Torne River

Station name	n r		P value	Mann–Kendall t	Wald-Wol- fowitz test	
				Test statistics	P value	P value
Kukkolankoski Övre	108	- 0.08	0.23	1.85	0.06	0.86
Pajala Pumphus	50	- 0.11	0.25	0.32	0.75	0.94
Abisko	35	0.03	0.80	0.21	0.83	0.17
Junosuando	52	- 0.09	0.34	1.41	0.16	0.95
Övre Abiskojokk	34	- 0.06	0.66	- 0.62	0.53	0.33

where n represents the sample size of time series and r indicates the Pearson correlation coefficient

Table 4 Estimated parameters with MLE and LM methods

Station	Probabili	ty distrib	utions												
	GEV			P3			GUM		GLO			LN3			
	MLE			MLE	MLE			MLE		MLE			MLE		
	μ̂	â	κ̂	μ̂	â	ĥ	$\hat{\mu}$	â	$\hat{\mu}$	â	ĥ	$\hat{\mu}$	â	ĥ	
16722	1993.73	450.20	0.16	297.25	128.76	14.72	1955.47	433.50	2149.33	280.04	- 0.11	2152.08	483.32	- 0.17	
	(48.38)	(34.34)	(0.07)	(763.60)	(58.50)	(12.50)	(43.91)	(33.52)	(49.03)	(22.95)	(0.07)	(51.38)	(33.78)	(0.09)	
2012	768.35	223.53	0.43	1663.77	- 52.30	15.99	718.05	240.19	839.62	117.07	0.08	844.26	204.85	0.16	
	(34.29)	(24.67)	(0.08)	(516.20)	(36.45)	(20.76)	(36.57)	(25.24)	(29.98)	(13.92)	(0.10)	(31.44)	(21.03)	(0.12)	
2357	207.26	52.41	0.24	- 285.00	5.69	90.04	200.73	50.87	225.42	30.92	- 0.04	225.60	53.77	- 0.06	
	(9.80)	(6.86)	(0.11)	(869.20)	(9.86)	(308.12)	(9.01)	(6.84)	(9.49)	(4.35)	(0.12)	(9.93)	(6.45)	(0.15)	
4	281.69	74.13	0.02	113.75	40.43	5.18	280.81	73.61	307.58	48.63	- 0.20	308.92	85.91	- 0.32	
	(11.84)	(8.10)	(0.11)	(51.00)	(15.39)	(3.10)	(10.49)	(8.00)	(11.97)	(6.16)	(0.08)	(12.80)	(9.24)	(0.12)	
957	108.80	26.47	0.08	61.85	17.66	3.42	107.71	25.83	119.34	17.35	- 0.13	118.11	29.65	- 0.28	
	(5.18)	(3.77)	(0.14)	(27.44)	(13.39)	(4.08)	(4.56)	(3.47)	(5.56)	(2.54)	(0.14)	(5.76)	(3.86)	(0.19)	
Station	Probabili	ty distrib	utions												
	GEV			Р3			GUM			GLO			LN3		
	LM			LM			LM		LM			LM			
	μ̂	â	ĥ	$\hat{\mu}$	â	ĥ	$\hat{\mu}$	â	$\hat{\mu}$	â	ĥ	$\hat{\mu}$	â	ĥ	
16722	1990.07	456.59	0.15	11.28	114.04	19.12	1959.44	403.29	2157.98	276.98	- 0.07	2154.46	490.66	- 0.15	
	(46.62)	(36.53)	(0.07)	(994.14)	(57.03)	(19.50)	(40.39)	(36.53)	(46.90)	(22.79)	(0.05)	(48.38)	(35.96)	(0.09)	
2012	764.68	220.31	0.40	1964.77	- 39.13	29.07	728.59	170.98	839.06	117.80	0.06	840.27	208.73	0.12	
	(33.65)	(23.58)	(0.11)	(781.80)	(42.25)	(22.81)	(26.28)	(22.34)	(28.85)	(13.84)	(0.08)	(32.31)	(22.61)	(0.13)	
2357	206.33	53.49	0.22	- 305.49	5.80	91.88	201.21	45.18	225.54	31.26	- 0.03	225.36	55.39	- 0.07	
	(10.49)	(6.96)	(0.13)	(270.70)	(15.19)	(23.56)	(8.01)	(7.06)	(9.51)	(4.46)	(0.09)	(10.37)	(6.62)	(0.16)	
4	279.83	73.16	- 0.01	148.25	50.90	3.43	280.25	74.02	308.19	48.68	- 0.18	306.66	85.97	- 0.37	
	(10.94)	(9.12)	(0.10)	(89.07)	(21.98)	(8.62)	(10.75)	(9.07)	(12.16)	(6.18)	(0.08)	(13.10)	(9.70)	(0.14)	
957	109.40	28.50	0.14	- 3.65	7.84	16.06	107.62	25.38	119.92	17.40	- 0.08	119.68	30.82	- 0.17	
	(5.50)	(4.18)	(0.13)	(103.66)	(7.83)	(20.99)	(4.55)	(4.10)	(5.26)	(2.54)	(0.10)	(5.90)	(3.93)	(0.17)	

Here, $\hat{\mu}$, $\hat{\alpha}$ and $\hat{\kappa}$ represent the estimated location, scale and shape parameters, respectively. The value in parenthesis is a standard error of estimated parameter

the LM method are best-fit distributions for gauging site Kukkolankoski Övre and Övre Abiskojokk, respectively.

In this study, a single distribution has not emerged as the best distribution for all gauging sites. This was also the case in [1, 5]. Overall, the LM estimation method performed better for identifying the suitable distribution (also see, [1]). The most suited distribution with MLE estimation method is identified at gauging site which has the highest CV and skewness, see Tables 5 and 6. It seems that the sites having extreme average of annual maxima of flood and catchment area (either very large or very small) are in favour of the LM method of estimation, see Tables 1, 5 and 6. If we look the landscape setting, the gauging sites which are at an extreme position (close and far away) to the Gulf

Table 5 Descriptive statistics (cubic metre per second)

Station	n	Mean	Median	S	CV	Skewness	Kurtosis
16722	108	2192.22	2140.00	493.03	0.22	0.37	- 0.19
2012	50	827.28	837.53	211.16	0.26	- 0.54	0.90
2357	35	227.29	223.53	54.71	0.24	0.17	- 0.07
04	52	322.98	296.50	93.32	0.29	0.93	0.95
957	34	122.27	120.67	31.39	0.26	0.73	1.24

Here, S represents the standard deviation and CV indicates the coefficient of variation

of Bothnia are in favour of the LM estimation method. The sample size of the time series does not seem to be an important factor in favour of particular distribution or estimation method in this study.

One major objective of flood frequency analysis is to estimate the quantiles in the extreme upper tail of the best-fitted distribution at each gauging site. The quantiles estimate for the return periods 5, 10, 25, 50, 100, 200 and 500 years are calculated by using quantile function and parameters value of the best-fitted distributions. Quantile estimate y_T with non-exceedance probability F for the best-fitted distributions is given in Table 7. The estimate of uncertainty (σ_s) in quantile estimates and 95% confidence intervals of quantiles of flood for different return period are also presented in Table 7. The SE indicates that longer return periods have more uncertainty around the flood quantile estimates.

5 Conclusion

In this study, the annual maximum steam flow series of five gauging sites of Torne River are examined. Flood frequency analysis is performed by using GEV, P3, GUM, GLO and LN3 distributions. The MLE and LM parameter estimation techniques are used to estimate the distribution's parameters. The study investigates the selection of best-fit probability distribution and estimation method for at-site flood frequency analysis of Torne river. The best-fit frequency distribution is identified at each gauging site based on the highest total rank score in goodness-of-fit tests and accuracy measures.

The results indicate that the GLO distribution using MLE for gauging site Junosuando and the LN3 distribution with a LM method for Pajala Pumphus and Abisko perform better than other distributions of this study. The GEV and P3 distributions using the LM method are the most suitable distribution at Kukkolankoski Övre and Övre Abiskojokk, respectively. At most gauging sites, the best distributions using LM estimation method are identified as the best-suited distributions.

The results found in this research study for flood frequency analysis of Torne River can be used in flood study, water resource planning and designing of hydraulic structures within the same basin and similar catchments. The best-fitted distributions used in this study could be considered as candidate distributions for regional flood frequency analysis of Torne River basin or at-site flood frequency analysis on other rivers in Sweden as well.

Table 6 Rank score of distribution in both GOF tests and accuracy measures

Station	Method	Distribution	AD	RMSE	MAE	RMSEP	MAEP	R^2	Total rank
Kukkolankoski Övre	MLE	GEV	7	5	7	7	7	5	38
		P3	6	7	6	6	5	7	37
		GLO	4	3	4	4	4	3	22
		LN3	5	6	5	5	6	6	33
		GUM	2	1	2	2	2	1	10
	LM	GEV	10	10	10	10	10	10	60
		P3	9	9	9	9	9	9	54
		GLO	3	4	3	3	3	4	20
		LN3	8	8	8	8	8	8	48
		GUM	1	2	1	1	1	2	08
Pajala Pumphus (2012)	MLE	GEV	3	4	7	4	4	4	26
rajaia rampiras (2012)	*****	P3	6	5	5	5	5	5	31
		GLO	4	8	3	8	8	8	39
		LN3	7	6	6	6	6	6	37
		GUM							
	1.04		1	1	1	2	1	1	7
	LM	GEV	5	3	8	3	3	3	25
		P3	9	9	9	9	9	9	54
		GLO	8	7	4	7	7	7	40
		LN3	10	10	10	10	10	10	60
		GUM	2	2	2	1	2	2	11
Abisko (2357)	MLE	GEV	3	3	3	3	4	3	19
		P3	8	5	6	4	6	5	34
		GLO	4	6	4	9	3	6	32
		LN3	7	4	5	5	5	4	30
		GUM	2	1	2	2	2	1	10
	LM	GEV	6	7	8	6	8	7	42
		P3	9	9	9	7	9	9	52
		GLO	5	8	7	10	7	8	45
		LN3	10	10	10	8	10	10	58
		GUM	1	2	1	1	1	2	8
Junosuando (4)	MLE	GEV	5	1	1	2	2	1	12
		Р3	2	1	1	2	2	1	09
		GLO	10	10	6	9	7	10	52
		LN3	4	3	4	5	5	3	24
		GUM	8	6	9	6	8	6	43
	LM	GEV	6	9	8	7	6	9	45
		P3	1	4	5	1	1	4	16
		GLO	9	5	3	10	9	5	41
		LN3	3	7	7	4	4	7	32
		GUM	7	8	10	8	10	8	51
Övre Abiskojokk (957)	MIE								
Ovie Abiskojokk (957)	MLE	GEV	6	5	7 1	7	7	5	37
		P3	1	7	1	4	1	7	21
		GLO	5	10	4	3	4	10	36
		LN3	4	6	5	6	6	6	33
		GUM	3	9	3	5	3	9	32
	LM	GEV	8	1	8	9	8	1	35
		P3	9	2	10	10	10	2	43
		GLO	7	4	6	1	5	4	27
		LN3	10	3	9	8	9	3	42
		GUM	2	8	2	2	2	8	24

The bold values are rank score for best-fitted distribution



Table 7 Quantile estimates of flood with 95% confidence interval at five gauging sites of Torne River

Station	Distribution	Method	Statistics	Non-exceedance probability and return periods (years)							
				0.80	0.90	0.96	0.98	0.99	1.00	1.00	
				5.00	10.00	25.00	50.00	100.00	200.00	500.00	
16772	GEV	LM	Lower	2475.06	2703.84	2929.24	3056.12	3157.22	3236.94	3324.16	
			Fit	2601.25	2857.74	3142.00	3327.52	3492.74	3640.52	3812.68	
			Upper	2728.22	3011.45	3358.67	3612.61	3867.51	4120.59	4461.08	
			$\sigma_{\scriptscriptstyle S}$	64.33	78.28	109.83	142.87	181.92	225.49	288.13	
2012	LN3	LM	Lower	943.17	1016.74	1080.04	1114.79	1141.18	1164.31	1188.27	
			Fit	1007.08	1087.59	1168.75	1218.72	1262.09	1300.52	1345.53	
			Upper	1067.68	1155.55	1258.97	1332.30	1402.33	1470.83	1558.80	
			$\sigma_{\scriptscriptstyle S}$	31.69	35.55	45.80	55.86	66.96	78.66	94.65	
2357	7 LN3	LM	Lower	250.78	272.75	293.06	304.31	313.38	321.13	330.25	
			Fit	273.37	299.61	328.49	347.65	365.24	381.63	401.88	
			Upper	296.48	327.91	368.54	400.46	433.38	466.35	511.51	
			$\sigma_{\sf s}$	11.66	14.15	19.52	24.73	30.66	37.15	46.49	
04	GLO	MLE	Lower	350.20	392.09	446.76	486.07	529.94	577.23	641.49	
			Fit	385.52	442.47	525.23	596.78	678.28	771.63	916.74	
			Upper	424.56	499.39	629.33	754.07	920.40	1132.33	1491.74	
			$\sigma_{\sf s}$	18.70	27.11	46.15	68.45	99.73	142.66	223.51	
957	57 P3	LM	Lower	133.74	146.98	159.11	166.19	171.67	176.38	181.46	
			Fit	147.67	163.85	182.29	194.88	206.65	217.81	231.83	
			Upper	162.08	183.06	210.76	231.51	251.71	272.35	299.50	
			$\sigma_{\sf s}$	7.19	9.28	13.09	16.44	20.02	23.77	28.89	

The lower limit is the fifth percentile and upper limit is the 95th percentile of the quantile distribution. The standard deviation of a quantile distribution is σ_s which represents the standard error of quantile estimates. The flood quantile estimates for different return periods indicate with bold value

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Data availability The data that support the findings of this study are openly available at SMHI [25].

Compliance with ethical standards

Conflict of interest The authors declare that they have no conflict of interest.

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