**Research Article** 

# Groundwater quality assessment in Jirania Block, west district of Tripura, India, using hydrogeochemical fingerprints

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#### Abstract

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A hydrogeochemical assessment was carried out in Jirania Block of west district of Tripura, India, to evaluate the geochemical processes and groundwater suitability for drinking and irrigation purposes. Twenty-two groundwater samples were collected at different lithologies in the study area during post-monsoon period. The collected samples were analyzed for major ions and heavy metals using the standard protocols. The analyzed parameters were used for various geochemical plots and indices to classify the groundwater quality. The order of dominance of cations is  $Ca^{2+} > Mg^{2+} > Na^+ > K^+$  and  $HCO_3^- > CI^- > SO_4^{2-} > NO_3^-$  for anions. Higher concentration of Fe and Mn was observed in most of the samples, which are above the acceptable limit of BIS standard. Correlation analysis and factor analysis were applied to unravel the relationship between ions and to identify the possible geochemical processes in groundwater. The interpretation of analytical data, plots and indices reveals that the weathering, mineral dissolution and ion exchange reactions are the major controlling factors for the groundwater geochemistry. The quality of the groundwater is suitable for drinking and irrigation purposes to the local community in this assessment period.

Keywords Groundwater quality · Bicarbonate · Weathering · Factor analysis · Heavy metals

# **1** Introduction

Groundwater is an important source of water supply for domestic, industrial, agricultural and recreational activities in both urban and rural places, and one-third of the world's population is using groundwater for drinking purpose [1–4]. Less bacteriological contamination and occurrence of diverse health necessity nutrients makes the groundwater more reliable in contrast to surface water. The suitability of groundwater is depending upon its steady temperature, natural quality, availability and low susceptibility. The quality of groundwater is also related with various geological and chemical influences that include precipitation, recharge quality, water–rock relation, dissolution, mineralization, ion exchange, etc [3, 5]. Groundwater is one of the prime water supply sources for drinking and agricultural purposes found in most parts of India. A survey of World Bank together with Central Groundwater Board has reflected that India is the world's largest user of groundwater and major fresh groundwater resources of the country is to be utilizable [6–8]. Groundwater pollution is a serious issue all over the world in recent period. Intensive use of groundwater with respect to urbanization, industrialization and agricultural activity has affected both quality and quantity of groundwater imposing greater risk to the groundwater resources. When pollution comes into the subsurface environment, it is distributed over the broad area of groundwater aquifer and makes it unsuitable for consumption and other various utilizations.

In India, it has been found that about 21% of transmissible diseases are water related [9–11]. Cardiovascular

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SN Applied Sciences (2019) 1:1055 | https://doi.org/10.1007/s42452-019-1092-1

Received: 7 May 2019 / Accepted: 13 August 2019 / Published online: 19 August 2019

diseases, problems in nervous system and multiple organ troubles are greatly connected to the consumption of polluted drinking water. It has been found that higher intake of calcium, magnesium, potassium and sodium give serious impact on hypertension and cardiovascular diseases in humans [12]. Higher concentration of fluoride through drinking water creates disorders in skeletal structure, bladder and uterine cancer. Elevated concentration of different types of toxic heavy metals in food and water is unsafe to human health [13].

In adults, arsenic is the most worried issue that cause of acute heavy metal poisoning. Human exposure to arsenic can cause cancer, hyperkeratosis, cardiovascular and nervous system problems. Consumption of chromium through food and water is also very unsafe and harmful, critically damaging lungs and kidneys. High concentrations of lead in drinking water make carcinogenic effects for humans. Ingestion caused by iron in drinking water is another metal pollution alarm especially for young kids as iron rapidly absorbed in gastrointestinal tract. High level of iron exposure can also cause bad effects to liver and kidneys [14].

Asghari et al. [15] have used the physical and chemical elements of groundwater in Zanjan area, Iran, to identify the suitability for drinking purpose, and they found that the groundwater is within the WHO standards and appropriate for drinking. Similar study was carried out by Soleimani et al. [16] in Sarpol-e Zahab area, Iran, and found that higher water hardness was observed in most of the samples.

The hydrochemistry of groundwater provides important facts in order to observe the suitability of water for both domestic and drinking purposes. The facts include physiochemical variation in groundwater of a particular area that is greatly affected by the local geology, geological structures and anthropogenic impacts [6]. Significant studies on hydrogeochemistry of groundwater have been investigated in various parts of India [2, 17–25].

The main aim of the current research is to evaluate the hydrochemistry of groundwater for drinking and irrigation purposes in Jirania Block of West Tripura District, Tripura, India, where groundwater is a prime source for the local people. Through the literature survey, it is found that there is no systematic work on hydrogeochemistry of groundwater in this area. Thus, it is important to estimate the extent to which the water consumed by the residents of rural areas has been polluted. The important significance of the study is that the majority of the population of this area belongs to tribal community where people are less aware about the health issues caused by poor quality of water due to poor literacy rate and general awareness. Based on the above facts, the groundwater quality of this region is prime important. Hence, the objective of this research is to use the hydrogeochemical tool to assess the suitability of groundwater in terms of domestic and irrigation purposes.

This study also aims at (includes) the understanding of the various geochemical processes which alter the groundwater quality in the study area.

# 2 Study area

This study was conducted at Jirania Block under the West Tripura District in the State Tripura, India. The area is geographically bound by latitude between 23°40'0"N and 23°55'0"N and longitude between 91°20'0"E and 91°30'0"E with an average altitude of 45 m above the MSL (Fig. 1). It is located about 18 km away from Agartala, the state capital of Tripura, and it is connected by the national highway NH-44. The climate of study area is tropical monsoon type. The average annual rainfall is around 218 cm, and it receives maximum rainfall from South-west monsoon, which remains in the month of June to September. The total geographic area is 212.32 km<sup>2</sup>, and the total population of Jirania is 55,873. According to the land use/cover in the study area, natural vegetation covers 101.93 km<sup>2</sup> followed by agriculture (86.38 km<sup>2</sup>), settlement (16.48 km<sup>2</sup>), water body (5.18 km<sup>2</sup>) and barren land (2.35 km<sup>2</sup>).

## 2.1 Geology and lithology

The study area consists of three main geological Formations viz, Tipam, Dupitila and Bokabil and also the Alluvium deposit along the banks of Haora River in the west district (Fig. 1). Tipam Formation mainly consists of sandstones, which is soft and fragile and associated with alternating layers of clay/mud [26]. Dupitila Formation overlaid by Tipam Formation, and it consists of shaley sandstone. Bokabil Formation overlaid by Dupitila Formation and mainly consists of shale with sandstone and limestone beds. This succession consists of some primary sedimentary structures. Quartz, feldspars and micas are the major mineral content in those rocks.

In Tripura, all the formations are comprised of sedimentary rocks with good porosity. These rocks are considered as potential aquifers and conduits for groundwater movement, in which the groundwater can be pumped for various purposes [26]. Groundwater in hilly regions of Tripura is the noticeable solution for public water supply. The major lithological units in the study area range from sandstone in Tipam Formation to shale–sandstone in Bokabil Formation as shown in Fig. 1. The sandstones are medium to fine grained. Alluvium deposit along the banks of seasonal rivers is represented by unconsolidated silt, sand, clay and decomposed organic matters.

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Fig. 1 Geology and sample location map of the study area

# 3 Methodology

Twenty-two groundwater samples were collected in Jirania Block (Fig. 1) during post-monsoon period. One liter capacity of polyethylene (HDPE) bottles was used to collect the samples. Before sampling, bottles were washed thoroughly with dilute HNO<sub>3</sub> acid and then with distilled water. Quality control measures were adopted throughout the sampling to meet the Quality Assurance (QA) [27]. pH, electrical conductivity (EC) and total dissolved solids (TDS) were measured in the field using multiprobe water quality meter.

 $Ca^{2+}$ ,  $Mg^{2+}$ ,  $HCO_3^{-}$  and  $Cl^{-}$ were analyzed using titrimetric method [27]. Na<sup>+</sup> and K<sup>+</sup> were measured by atomic absorption spectrophotometer (PerkinElmer AA-Analyst 700).  $SO_4^{2-}$  was measured by turbidimetric method using  $BaSO_4$ , and  $NO_3^{-}$  was measured using UV visible spectrophotometer.

To confirm the analytical accuracy for samples, ionic balance error was computed using cation and anion concentrations [28]. The error percentage for the samples is within the standard limit of  $\pm$  5%.

For heavy metals analysis, samples were acidified using  $HNO_3$  by bringing the pH to ~2 and preserved at a temperature of 4 °C. Fe, Mn, Pb, Cd, As, Cu and Zn

were analyzed using atomic absorption spectrophotometer (PerkinElmer A Analyst 700). Most of the samples show below detection limit for Pb, Cd and As. Hence, these metals were not considered for further discussion. The minimum detection limit for Fe, Mn, Cu and Zn is 0.001 mg/L.

WATCLAST Program [29] was used for the analytical data to classify the groundwater based on geochemical indices and for graphical representation. Log pCO<sub>2</sub>, ionic strength and the saturation index of carbonate minerals were calculated using WATQ4F geochemical modeling software [30]. Correlation and factor analysis were used for interpreting the hydrochemical data set using Statistical Package for Social Studies (SPSS) version 17. Maps and spatial diagrams were created using ArcGIS version 10.1.

# 4 Results and discussion

## 4.1 General chemistry

Table 1 shows the physical and chemical composition of groundwater in the study area. pH in water is an important indicator for quality and also for geochemical reactions [31]. pH in the groundwater samples ranging from 5.68 to 7.83 with an average of 6.53. It indicates that the

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Sample ID	Location names	pН	EC	TDS	Ca <sup>2+</sup>	$Mg^{2+}$	Na <sup>+</sup>	K <sup>+</sup>	$HCO_3^-$	Cl⁻	SO4 <sup>2-</sup>	NO <sub>3</sub> <sup>-</sup>
1	Champaknagar	6.0	186.4	112.0	18.72	13.61	7.99	3.09	146.16	23.57	10.84	1.26
2	Pas. Debendranagar	5.7	75.1	48.0	5.90	3.74	4.21	2.49	40.32	18.86	10.89	0.58
3	Jiraniakhola	6.2	218.0	142.0	18.72	11.34	10.30	4.43	176.40	14.14	15.41	0.59
4	Belbari	5.8	86.4	52.0	5.90	3.74	4.38	1.68	60.48	14.14	11.16	2.07
5	Madhabbari	6.4	228.0	152.0	18.72	11.34	9.71	2.84	176.40	14.14	10.21	0.51
6	Radhapur	6.8	183.5	118.0	10.20	7.77	2.67	0.89	146.16	18.86	10.84	0.64
7	Champabari	7.8	193.1	96.0	10.20	8.01	2.99	0.32	75.60	14.14	13.83	0.68
8	Uttar Joynagar	6.8	119.9	60.0	10.20	7.77	1.59	0.29	45.36	14.14	15.41	1.64
9	Janmejoynagar	7.8	127.5	64.0	10.20	7.77	3.67	0.34	45.36	18.86	37.11	1.51
10	Kobrakhamar	6.4	176.4	88.0	10.20	7.77	1.26	0.32	85.68	37.72	14.66	0.98
11	Bankimnagar	6.3	151.4	76.0	10.20	7.77	1.43	0.72	40.32	28.29	9.66	1.87
12	Harijoy Chow Para	6.5	260.0	132.0	10.20	7.77	3.69	0.29	115.92	37.72	12.00	2.59
13	Radhamohanpur	6.8	184.9	92.0	16.03	9.08	2.86	0.48	90.72	14.14	23.77	1.84
14	Kaiyachangbari	7.5	162.6	82.0	16.03	11.34	2.87	0.35	75.60	14.14	37.11	1.69
15	Barjala Binapani	6.7	187.5	94.0	10.20	4.54	0.90	0.17	110.88	18.86	7.88	0.52
16	Joynagar	6.2	127.5	86.0	10.20	7.77	1.83	0.78	115.92	42.66	10.89	0.58
17	Pas. Jiraniakhola	6.1	153.9	98.0	5.90	3.74	7.33	1.82	146.16	14.14	15.41	0.74
18	Dak. Radhapur	6.3	139.3	88.0	16.03	9.08	1.39	0.27	95.76	42.66	15.41	1.36
19	Pb. Janmey joynagar	6.9	143.2	94.0	10.20	4.54	2.37	0.88	60.48	28.44	10.89	0.54
20	Pb. Barjala	6.4	151.7	98.0	5.90	3.74	0.66	0.12	110.88	33.18	10.89	1.63
21	East Belbari	5.9	101.3	68.0	22.47	13.61	1.00	0.18	70.56	23.57	15.41	1.28
22	Majlishpur	6.4	183.0	118.0	10.20	4.54	3.36	0.52	115.92	18.86	10.84	1.37

Table 1 Chemical composition of groundwater samples in the study area (all the values are in mg/L except pH and EC in µs/cm)

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Fig. 2 Spatial map of EC in the study area

# **Fig. 3** Piper plot for the groundwater samples







Fig. 4 Gibbs plot for the mechanisms controlling the groundwater chemistry

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Fig. 5 Relationship between cations, anions and the total concentration

groundwater is near neutral condition and few samples in acidic nature. This kind of water may aggravate the dissolution of minerals during weathering process. EC values range from 75.1 to 260  $\mu$ S/cm with an average of 160.94  $\mu$ S/cm. According to Wilcox [32] classification, all the samples fall in excellent category, except one sample. The spatial distribution of EC values shows that the higher values are noted in the western side of the study area (Fig. 2). TDS values range from 48 to 152 mg/L with an average of 93.6 mg/L, and it follows the same spatial trend of EC. Ca<sup>2+</sup> is the dominant cation ranging from 5.9 to 22.47 mg/L with an average of 11.93 mg/L, followed by Mg<sup>2+</sup>, Na<sup>+</sup> and K<sup>+</sup>. HCO<sub>3</sub><sup>-</sup> is the dominant anion ranging from 40.32 to 176.4 mg/L with an average of 97.6 mg/L, followed by Cl<sup>-</sup>, SO<sub>4</sub><sup>2-</sup> and NO<sub>3</sub><sup>-</sup>.

### 4.2 Hydrogeochemical processes

Piper plot [33] was used to categorize the major water types in the study area. In this plot, most of the samples fall in Ca–HCO<sub>3</sub> water type (Fig. 3), indicating the dominance of weathering process by the fresh water recharge [2]. It also shows that the water–rock interaction is the major process in the study area. Some minor representations are

in Ca–Cl water type. Overall, the plot shows that the alkaline earth (Ca<sup>2+</sup> and Mg<sup>2+</sup>) exceeds alkali (Na<sup>+</sup>) and weak acids (HCO<sub>3</sub><sup>-</sup>) exceeds the strong acid (Cl<sup>-</sup>).

Gibbs plot [34] was also used to identify the major hydrogeochemical processes viz weathering, evaporation and precipitation, which may control the groundwater chemistry. This plot consists of two separate diagrams for cations and anions. Majority of the samples fall in weathering zone, indicating the water-rock interaction dominance [35]. Few samples fall in precipitation zone (Fig. 4).

The concentration of ions and their interrelationships is mainly used for the chemical classification. The plot (Ca + Mg) versus HCO<sub>3</sub> shows that most of the samples fall below equiline 1:1 indicating the predominance of bicarbonate over alkali earth (Fig. 5). This is due to the intensive weathering of rocks. The (Ca + Mg) versus TZ<sup>+</sup> plot lies below the equiline 1:1 indicating the dominant of total cations over alkali earth. The plot for (Na + K) versus TZ<sup>+</sup> also indicates the dominance of total cations over alkali earth. Na versus Cl plot shows that the majority of samples fall in Cl zone indicating the leaching of secondary salts [36].

Since weathering is the chief controlling process in the groundwater, Chloro-alkaline indices, i.e., CAI1 and CAI2

<u> </u>	G 1	Samples	Category Grade Samples			<u> </u>	Samples
Category	Grade	( <i>n</i> = 22)	Category Grade		( <i>n</i> = 22)	Category	( <i>n</i> = 22)
Na% Wilcox (1955)			USGS hardı	iess	TDS classification (USSL, 1954)		
Excellent	0–20	17	Soft <75 15		< 200	22	
Good	20-40	5	Slightly hard	75-150	7	200–500	0
Permissible	40-60	0	Moderately hard	150- 300	0	500-1500	0
Doubtful	60-80	0	Very hard	> 300	0	1500-3000	0
Unsuitable	> 80	0	IBE Schoeller	(1965)	Cation facio	es	
Na%	Eaton (19	50)	(Na+k)rock->Ca/Mg g.w.		2	Ca–Mg facies	21
Safe	< 60	22	(Na+k)g.w>Ca/Mg rock		20	Ca–Na facies	1
Unsafe	> 60	0				Na–Ca facies	0
S.A.R. I	Richards (1	954)	Schoeller classification (1967)	1		Na facies	0
Excellent	0-10	22	Type I		21	Anion facie	2S
Good	18-Oct	0	Type II		1	HCO <sub>3</sub> facies	0
Fair	18-26	0	Type III		0	HCO <sub>3</sub> -Cl-SO <sub>4</sub> facies	5
Poor	> 26	0	Type IV		0	Cl-SO <sub>4</sub> -HCO <sub>3</sub> facies	17
R.S.C.	Richards(1	954)	Corrosivity Rati	o ( <b>1990</b> )		Cl-facies	0
Good	< 1.25	21	Safe	< 1	2	Hardness classification (Ha	nda, 1964)
Medium	1.25- 2.5	1	Unsafe	> 1	20	Permanent hardness (NCH)	
Bad	> 2.5	0	Chloride classification (Stuyf	zand, 1989	A1	4	
EC Wilcox (1955)			Extremely fresh		0	A2	2
Excellent	< 250	21	Very fresh	17		A3	0
Good	250– 750	1	Fresh	5		Temporary hardness (CH)	
Permissible	750– 2250	0	Fresh Brackish	0		B1	14
Doubtful	2250- 5000	0	Brackish	0		B2	0
Unsuitable	> 5000	0	Brackish-salt	0		B3	0
			Salt		0		
			Hyperhaline	0			

 Table 2
 WATCLAST program output for the groundwater samples





Fig. 6 Log pCO<sub>2</sub> versus ionic strength

Fig. 7 Saturation index of carbonate minerals

Table 3 Summary of the chemical composition of groundwater and their standards (all the values are in mg/L except pH and EC in  $\mu$ s/cm)

Parameters	Max	Min	Avg	BIS standard (acceptable limit)	WHO limit
pН	7.83	5.68	6.53	6.5–8.5	6.5–8.5
EC	260	75.1	160.94	300	1500
TDS	152	48	93.6	500	500
Ca <sup>2+</sup>	22.47	5.9	11.93	75	75
Mg <sup>2+</sup>	13.61	3.74	7.74	30	50
Na <sup>+</sup>	10.3	0.659	3.57	50	200
K <sup>+</sup>	4.43	0.121	1.05	-	12
Cl⁻	42.66	14.14	22.96	200	250
HCO3-	176.4	40.32	97.6	30	500
$NO_3^-$	2.07	0.51	1.2	-	45
SO4 <sup>2-</sup>	37.11	7.88	15.02	200	250

were used to assess the degree of base exchange during weathering [37]. During water–rock interaction, if there is an exchange of alkali elements (Na<sup>+</sup> and K<sup>+</sup>) in water with alkaline earth (Ca<sup>2+</sup> or Mg<sup>2+</sup>) in rock, then both indices are positive and vise versa. In this calculation, all the ionic concentration is expressed in epm values. Based on the calculation, positive values are observed in most of the samples indicating reverse ion exchange process during weathering (Table 2). In Schoeller [38], classification of water types, the majority of samples fall in type I and



Log pCO<sub>2</sub> and ionic strength are used to assess the nature of groundwater in the sense of residence time in the aquifer [39]. In fresh water, the ionic strength is less than 0.005 [31] and all the samples in the study area show lower ionic strength indicating higher inflow of fresh water recharge (Fig. 6). In Log pCO<sub>2</sub>, all the samples are higher than the atmospheric value (-3.5) suggesting longer residence time in the aquifer, which enhance the water–rock interaction [40].

Geochemical model, WATEQ4F [30], was used to calculate the saturation index (SI) of carbonate minerals in the groundwater. The results show that high SI of calcite followed by magnesite, anhydrite and dolomite (Fig. 7). All the samples fall under saturation state indicating the dissolution of minerals during water–rock interaction [41].

#### 4.3 Water quality

The analyzed parameters are compared with BIS [42] and WHO [43] standards to determine the drinking water quality (Table 3). All the ions are within the permissible limits of WHO standard except pH, in which 13 samples are less than 6.5.  $HCO_3$  values for all the samples are shown above the acceptable range of BIS standard.

Based on chloride classification by Stuyfzand [44], the majority of samples fall in very fresh and few samples in



**Fig. 8** USSL plot for the irrigation quality

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# **Fig. 9** Donnen plot for the irrigation water quality





fresh category (Table 2). For corrosivity ratio, most of the samples fall in safer zone, which is less than 1. Carbonate hardness (CH) is higher in most of the samples indicating temporary hardness. According to Sawyer and McCarty [45] hardness, samples vary from soft to slightly hard (Table 2).

Groundwater suitability for irrigation purpose is estimated by Na%, EC, sodium absorption ratio (SAR) and residual sodium carbonate (RSC) [32, 46]. Na concentration is an important indicator to determine the soil structure and crop yield [47]. In Na%, samples range from excellent to good category and in Na% Eaton [48] classification; all the samples are in safe category (Table 2). SAR values in all the samples fall in excellent category (Table 2). In residual sodium carbonate [46] classification, all samples are in good category except one sample in medium category, indicating that the water is fit for irrigation.

USSL plot is used to classify the irrigation water quality (Fig. 8). In the plot, most of the samples fall in C1S1 zone indicating low salinity and low sodium hazard, which is suitable for plants growth. Permeability index is used to determine the quality of irrigation water based on soil texture for development in agriculture. Based on Doneen [49] plot using permeability index values, groundwater is classified into class I, class II and class III to determine the suitability of water for irrigation purpose (Fig. 9). In this plot, samples are equally distributed in class I and II, which indicate good permeability and the water can be used for irrigation.

Table 4 shows the comparison of present study chemical composition of groundwater with few other parts

SN Applied Sciences A Springer Nature journal of world and India regions [50–60]. It is clearly noticed that the chemical concentration of West Tripura (present study) is much lower than the world and India regions. It is also interesting to observe that the chemical concentration of groundwater of Nalbari district of Assam (which is close to the state of Tripura) is lower than the other parts of world and India regions. But, most of the chemical elements of present study are lower than the groundwater of Nalbari district of Assam, except Cl and SO<sub>4</sub>. The higher concentration of Cl and SO<sub>4</sub> in the study area indicates the leaching of secondary salts and weathering of host rocks [17].

#### 4.4 Heavy metals

The heavy metal concentration shows significant variation in each sample (Fig. 10). Fe ranges from 120 to 5140 ppb with an average of 2692 ppb and 90% of samples fall above the BIS maximum acceptable limit (300 ppb). Mn ranges from 11 to 619 ppb with an average of 125 ppb. 40% of the samples falls above the BIS maximum acceptable limit (100 ppb). Cu ranges from 10 to 38 ppb with an average of 15 ppb, whereas Zn ranges from 13 to 50 ppb with an average of 27 ppb. All the samples fall well below the BIS maximum acceptable limit of Cu (50 ppb) and Zn (5000 ppb) [42]. Higher concentration of Fe and Mn is due to the weathering of bedrocks, which release metals into the groundwater [61].

lable 4 Comparison of chemical composition of th	ie preser	it stuay w	un ouner p		e worig and	) in India	all the va	ues are in r	ng/L exce	рт рп апа	i/ch ui ha	cm)
Study areas	Ηd	EC	TDS	Ca <sup>2+</sup>	Mg <sup>2+</sup>	Na <sup>+</sup>	κ+	HCO <sub>3</sub> <sup>-</sup>	CI-	$\mathrm{SO_4}^{2-}$	$NO_{3}^{-}$	References
Ulsan City, Korea ( <i>n</i> =40)	7.1	770	698	94.3	41.7	19.2	3.2	169	19	351	I	Venkatramanan et al. [50
South-eastern Tunisia ( $n = 47$ )	7.3	6390	3056	317.5	195	513.5	25.6	479.7	729	1193	41.2	Ayed et al. [ <b>51</b> ]
Ardabil area, Iran ( $n = 77$ )	7.4	1424	919	112.6	42.3	131.5	23.7	381.2	161	273.4	26.5	Aghazadeh et al. [ <mark>52</mark> ]
South-western Nigeria ( $n = 21$ )	6.8	1574	1136	78.4	89.9	193.4	5.8	411.2	181	298.2	3.4	Emenike et al. [ <mark>53</mark> ]
Gedaref State, Sudan ( $n = 40$ )	8	888	602	37.4	34.2	214.2	41.7	499.3	35.9	0.7	3.8	Elubid et al. [54]
Lakshimpur district, Bangladesh ( $n = 70$ )	7	1135	ı	55.7	46.1	159.7	10.8	430.1	227	16.1	ı	Bhuiyan et al. [ <mark>55</mark> ]
Central Telangana, India; ( <i>n</i> = 40)	7.1	888	594	61.7	60.8	233	4.2	672.6	154	59	48.9	Roy et al. [ <b>56</b> ]
Malwa Region, Punjab, India; ( <i>n</i> =24)	7.2	2016	1108	153.1	51.2	276.5	15.3	I	187	263.4	5.07	Kaur et al. [ <mark>57</mark> ]
Madurai Region, central Tamilnadu, India; ( <i>n</i> = 54)	7	1324	805	30.6	24.9	67.1	21.9	372.1	27.4	1.5	33.5	Adithya et al. [58]
Prakasam District, Andhra Pradesh, India; ( $n = 30$ )	7.8	2984	1983	65.6	76.8	444.7	54.4	742.6	556	71.4	56.2	Subba Rao [ <mark>59</mark> ]
Nalbari district of Assam, India; ( $n = 50$ )	7.4	367	235	28	9.5	16	2.9	136	5.2	3.8	1.6	Jain and Vaid [60]
West Tripura, India (present study); $(n = 22)$	6.5	161	94	11.9	7.74	3.5	1	97.6	22.9	15	1.2	Present study
n total number of samples												

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#### 4.5 Statistical analysis

#### 4.5.1 Correlation analysis

Correlation analysis was used for the groundwater samples to reveal the inter-relationship between ions [22, 24, 36]. Correlation coefficient > 0.5 considered as good correlation between ions. pH and  $SO_4^{2-}$  show good correlation indicating the dissolution of sulfate minerals during weathering process (Table 5).  $HCO_3^-$  had a good correlation between EC, TDS, Na<sup>+</sup> and K<sup>+</sup> indicating the dominance of  $HCO_3^-$  ion in groundwater due to weathering and dissolution of minerals [2]. Ca<sup>2+</sup> and Mg<sup>2+</sup> showed good correlation indicating the dissolution of carbonate minerals from the host rocks or simple cation exchange reaction.

#### 4.5.2 Factor analysis

Factor analysis was also used for the groundwater samples to determine the various geochemical process indicated by the factor loading of ions [20, 21]. Totally, four factors were extracted with the total variance of 24%, 22%, 19% and 17% (Table 6). Factor 1 is loaded with EC, TDS and  $HCO_3^-$ , indicating that the overall chemistry of groundwater is mainly controlled by  $HCO_3^-$  ion and also showing the indicative of dominant mineral dissolution to increase the  $HCO_3$  concentration. Factor 2 is loaded with Na<sup>+</sup> and K<sup>+</sup> indicating the feldspar mineral weathering from the source rock [62]. Factor 3 is loaded with Ca<sup>2+</sup> and Mg<sup>2+</sup> indicating the reverse ion exchange reaction, where the Ca and Mg are released from mineral weathering [58]. Factor 4 is loaded with pH and  $SO_4^{2-}$  indicating the dissolution of sulfate minerals during weathering.

#### 5 Conclusions

The outcome of this study has given a better insight on the nature of groundwater in terms of quality and the geochemical processes in the study area. Most of the samples show slightly acidic in nature, which enhance the dissolution of secondary minerals during water–rock interaction. Ca–HCO<sub>3</sub> is the major water type in Piper plot indicating the dominance of weathering process through fresh water recharge. Gibbs plot also shows the weathering is the major controlling process in groundwater. Higher Log pCO<sub>2</sub> values support the longer residence time of water in the aquifer matrix, which enhance the dissolution and ion exchange reactions. Saturation index of carbonates minerals shows the under-saturation condition, which again supports the mineral dissolution during weathering. The heavy metals in the groundwater are **Fig. 10** Heavy metals concentration in the groundwater



Table 5Correlation analysis forthe groundwater samples

	рН	EC	TDS	Ca <sup>2+</sup>	Mg <sup>2+</sup>	Na <sup>+</sup>	K <sup>+</sup>	HCO <sub>3</sub> <sup>-</sup>	Cl⁻	SO4 <sup>2-</sup>	NO <sub>3</sub> <sup>-</sup>
pН	1.000					1				0	
EC	.248	1.000									
TDS	.009	.893	1.000								
Ca <sup>2+</sup>	006	.292	.357	1.000							
Mg <sup>2+</sup>	.080	.292	.290	.919	1.000						
$Na^+$	217	.380	.553	.303	.294	1.000					
$K^+$	447	.154	.382	.257	.236	.898	1.000				
$HCO_3^-$	231	.669	.870	.327	.267	.640	.527	1.000			
Cl <sup>-</sup>	229	.002	.010	076	031	447	344	045	1.000		
SO4 <sup>2-</sup>	.620	097	254	.208	.289	024	212	279	258	1.000	
$NO_3^-$	.028	.007	213	052	.044	259	376	351	.134	.283	1.000

Bold values are statistically significant

Table 6 Factor analysis for the groundwater samples

	Factor 1	Factor 2	Factor 3	Factor 4
pН	.170	221	081	.903
EC	.942	.007	.134	.148
TDS	.946	.216	.161	087
Ca <sup>2+</sup>	.194	.093	.927	.032
Mg <sup>2+</sup>	.165	.039	.950	.101
Na <sup>+</sup>	.341	.836	.255	047
K <sup>+</sup>	.140	.861	.232	305
HCO <sub>3</sub> <sup>-</sup>	.784	.404	.169	249
Cl <sup>-</sup>	.129	703	.033	519
SO4 <sup>2-</sup>	241	060	.323	.799
NO <sub>3</sub> <sup>-</sup>	162	503	.179	.154
% of variance	24.804	22.364	19.058	17.684

Bold values are statistically significant

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mostly contributed from the weathering of rock forming minerals. Factor analysis indicates weathering, dissolution and ion exchange are the dominant processes that control the hydrogeochemistry of groundwater. Overall, the groundwater quality is suitable for drinking and irrigation purposes, except for the cautious of high level Fe, Mn and HCO<sub>3</sub>.

Acknowledgements The authors from Tripura University thank the Head of the Department of Chemistry, Tripura University for providing necessary laboratory support to carry out the chemical analysis for the water samples. The authors are also thankful to Ms. Ninu Krishnan Modon Valappil for producing geological and spatial maps.

#### **Compliance with ethical standards**

**Conflict of interest** On behalf of all the authors, the corresponding author states that there is no conflict of interest.

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