ORIGINAL ARTICLE



Groundwater quality around Tummalapalle area, Cuddapah District, Andhra Pradesh, India

Y. Sreedhar¹ · A. Nagaraju¹

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Abstract The suitability of groundwater for drinking and irrigation was assessed in Tummalapalle area. Forty groundwater samples were analysed for major cations, anions and other parameters such as pH, electrical conductivity, total dissolved solids (TDS), total alkalinity and total hardness (TH). The parameters such as sodium adsorption ratio, adjusted sodium adsorption ratio (adj.SAR), per cent sodium, potential salinity, residual sodium carbonate, non-carbonate hardness, Kelly's ratio and permeability index were calculated for the evaluation of irrigation water quality. Groundwater chemistry was also analysed by statistical analysis, USSL, Wilcox, Doneen, Piper and Chadhas diagrams, to find out their suitability for irrigation. TDS and TH were used as main parameters to interpret the suitability of groundwater for drinking purpose. The correlation coefficient matrix between the hydrochemical parameters was carried out using Pearson's correlation to infer the possible water–rock interactions responsible for the variation of groundwater chemistry and this has been supported by Gibbs diagram. The results indicate that the groundwater in Tummalapalle area is alkaline in nature. Ca-Mg-HCO3 is the dominant hydrogeochemical facies. Water chemistry of the study area strongly reflects the dominance of weathering of rockforming minerals such as bicarbonates and silicates. All parameters and diagrams suggest that the water samples of the study are good for irrigation, and the plots of TDS and TH suggest that 12.5% of the samples are good for human consumption.

A. Nagaraju arveti@yahoo.com

Keywords Hydrogeochemistry · Water quality · Tummalapalle · Andhra Pradesh

Introduction

Water is an indispensable natural resource on earth. Safe drinking water is the primary need of every human and also their basic fundamental right (Nagaraju et al. 2015, 2016c; Li et al. 2015). Fresh water has become a scarce commodity due to over exploitation and pollution of water. Groundwater is the most important source of water supply for drinking, agriculture, irrigation and industrial purposes (Gupta and Sunita 2009; Bhattacharya et al. 2012; Li et al. 2013a, 2014a; Nagaraju et al. 2013, 2014a, 2016a, b; Wu and Sun 2016). There is a growing awareness of the environmental legacy of mining activities that have been undertaken with little concern for the environment (Mahesh et al. 2001; Brindha et al. 2010; Tripathi et al. 2011). Mine water can vary greatly in the concentration of contaminants present, and some mine water discharges can be a potential water resource, where the local water demands for industrial, irrigation and even drinking and domestic uses can be fulfilled by effective utilization (Cidu et al. 2007; Manish Kumar et al. 2007; Suresh et al. 2007; Li et al. 2013b; Nagaraju et al. 2006, 2014b, 2015; Wu et al. 2014).

Metal pollution by mining and associated industrial activities is somewhat mitigated today by strict implementation of clean technology and environmental measures. The drastic increase in population, urbanization and modern land-use applications and demands for water supply has limited the globally essential groundwater resources in terms of both its quality and quantity. Further, the



Department of Geology, S V University, Tirupati, Andhra Pradesh 517 502, India

quality is a function of the physical, chemical and biological parameters, and can be subjective, since it depends on a specific intended use (Ravikumar and Somashekar 2013). Tummalapalle is a famous uranium mining area in India, and the mining activities generally cause groundwater pollution. Therefore, the purpose of this paper is to summarize groundwater quality in and around Tummalapalle uranium mining area and focus its impact on water quality for irrigation and drinking purpose.

Geology of the study area

Geologically, Tummalapalle area is in the south-western part of the Cuddapah basin. Tummalapalle (14°18′ to 14°24′N; 78°13′30″ to 78°22′30″E) is located in Pulivendla Taluk of Cuddapah District of Andhra Pradesh (Fig. 1). It is included in the Survey of India topo sheet No. 57 J/7. Lithologically, the Vempalle Formation consists of purple shale, massive limestone, intraformational conglomerate, dolostone (uraniferous), shale and cherty limestone. The radioactive minerals identified in the ore zone are pitchblende, coffinite and U–Ti complex. The

impersistent conglomerate and purple shale band occurring immediately below and above the mineralized rock, respectively, serve as the marker horizons (Geological Survey of India 2001; Basu 2007).

Climate and rainfall

In this area, the climate is tropical with seasonal rainfall. This area has rather a hot summer with temperature as high as 42 °C and minimum temperature being 15 °C. This region is known for its wide variation in contour, heavy vegetation, low rainfall and significant variation in meteorological parameters. The average annual rainfall of the Cuddapah District is about 710 mm and it ranges from nil rainfall in January to 137 mm in October and this is the wettest month of the year. The mean seasonal rainfall distribution is 402.4 mm in south-west monsoon (June–September) and in north-east monsoon (October–December) this is about 239.1 mm. The percentage distribution of rainfall is about 56.7% in south-west monsoon and 33.7% in north-east monsoon (Central Ground Water Board 2013).

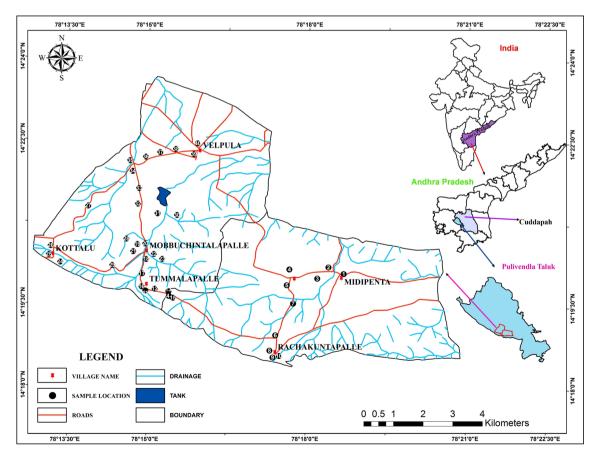


Fig. 1 Map of study area showing sampling locations



Hydrogeology

Groundwater occurs under water table conditions in the weathered zones of Papaghni and Chitravati group of rocks. The water present in the dug wells is mainly due to the numerous joints, fractures and fissures present in these rock types. The quartzites and the basal part of the massive limestone are good aquifers, and the permanent water table in these is generally shallow. Water is alkaline in nature and suits both for irrigation and drinking purposes (Geological Survey of India 2001). It is found that the groundwater in the Tummalapalle village is saline due to unhygienic conditions, since water found away from the village in surrounding area is generally sweet. The depth of water level is observed between 10 and 20 m in Vemula mandal (Central Ground Water Board 2007, 2013).

Methodology

Water samples were collected in and around the Tummalapalle uranium mining area. Sampling was carried out in the month of May 2014. A total of 40 groundwater samples were collected. Sampling locations were recorded using a potable GPS device and they are shown in Fig. 1. Samples were collected in precleaned and well-dried polyethylene bottles. The water samples collected in the field were analysed for electrical conductivity (EC), pH, total dissolved solids (TDS), major cations such as calcium, magnesium, sodium and potassium, and anions such as bicarbonate, carbonate, chloride, fluoride and sulphate, adopting the standard methods (Hem 1985; Raghunath 1987; APHA 2012). The sampling methods are chosen according to the objectives of the investigation and the geology (lithology, hydrogeology, foreign materials) of the area. The chemical quality data are shown in Table 1. Pearson correlation coefficient was calculated for measuring the strength of the linear relationship between these parameters. Various methods and graphs were used to study and interpret the water analyses data.

Results and discussion

Geochemical properties and principles that govern the behaviour of dissolved chemical constituents in ground-water are referred to as hydrogeochemistry. The dissolved constituents occur as ions, molecules or solid particles. The chemical composition of groundwater is related to the solid product of rock weathering and changes with respect to time and space. Hence, the variation in the concentration levels of the different hydrogeochemical constituents

determines its usefulness for domestic, industrial and agricultural purposes. The minimum, maximum, average, SD and SE values of the hydrochemical data of the study area are shown in Table 1 along with the calculated parameters of the hydrogeochemical data.

Groundwater chemistry

From Table 1, it is evident that in the study area, the pH is in the range of 6.70–7.90 and the EC is ranging from 12 10 to 2240 µmhos/cm. Calcium is varying from 25 to 173 mg/l and magnesium is ranging from 27 to 133 mg/l; sodium is between 6 and 86 mg/l and potassium is varying from 2 to 73 mg/l. The fluoride concentration is ranging from 0.11 to 0.62 mg/l. The total dissolved solids lie between 787 and 1456 mg/l. The concentration of sulphate varies between 8 and 94 mg/l. The bicarbonate and carbonate concentrations in groundwater ranges from 130 to 509 mg/l and 38 to 222 mg/l, respectively. The chloride content varies from 21 to 207 mg/l/. The alkalinity varies from 36 to 144 mg/l and the hardness values vary from 120 to 596 mg/l. The abundance of the major ions in groundwater is in following order:

$$Ca^{2+} > Mg^{2+} > Na^{+} > K^{+} \text{ and } HCO_{3}^{-} > CO_{3}^{-}$$

 $> Cl^{-} > SO_{4}^{2-} > F^{-}$

Classification of groundwater

Piper trilinear diagram

Piper (1944, 1953) based on the concentration of dominant cations and anions have proposed a trilinear diagram to show the percentages at mill equivalents per litre of cations and anions in water samples. The Piper diagram was modified by Davis and Dewiest (1967). This is useful to understand the total chemical character of water samples in terms of cation-anion pairs. The Piper diagram reveals similarities and differences among groundwater samples because those with similar qualities will tend to plot together as groups (Walton 1970). The Piper diagram consists of two triangular and one intervening diamondshaped field (Fig. 2). The percentage reacting values at the three cation groups—Ca, Mg and (Na + K)—are plotted as a single point in the left triangular field and the three anion groups—(HCO₃ + CO₃), SO₄ and Cl—similarly on the right triangular field.

In the present study, it is clear that about 42.5% samples are falling in the area of 5 and this indicates that carbonate hardness (secondary alkalinity) exceeds 50%. About 37.5% samples are falling in the area of 9 and this indicates that



Table 1 Descriptive statistics for groundwater samples of Tummalapalle area

S. No.	Constituents	Min	Max	Average	SD	SE
1	Calcium (Ca) (ppm)	25	173	77	26.06	4.12
2	Magnesium (Mg) (ppm)	27	133	66	23.33	3.69
3	Sodium (Na) (ppm)	6	86	33	16.72	2.64
4	Potassium (K) (ppm)	2	73	14	14.99	2.37
5	Bicarbonate (HCO ₃) (ppm)	130	509	303	79.75	12.61
6	Carbonate (CO ₃) (ppm)	38	222	113	36.01	5.69
7	Sulphate (SO ₄) (ppm)	8	94	17	15.56	2.46
8	Chloride (Cl) (ppm)	21	207	71	42.86	6.78
9	Fluoride (ppm)	0.11	0.62	0.40	0.10	0.02
10	рН	6.70	7.90	7.30	0.23	0.04
11	Specific conductance (µmhoscm ⁻¹)	1210	2240	1884	240.58	38.04
12	Total dissolved solids (ppm)	787	1456	1225	156.38	24.73
13	Hardness as CaCO ₃ (ppm)	120	596	296	103.24	16.32
14	Alkalinity as CaCO ₃ (ppm)	36	144	89	20.44	3.23
15	Sodium adsorption ratio (SAR)	0.15	1.54	0.67	0.33	0.05
16	Adjusted SAR (adj.SAR)	0.33	4.59	1.79	0.89	0.14
17	Sodium percentage	4.73	35.03	15.86	6.79	1.07
18	Potential salinity	0.72	6.13	2.17	1.20	0.19
19	Residual sodium carbonate	-3.92	1.03	-0.60	0.93	0.15
20	Permeability index (PI)	23.97	55.20	34.72	6.58	1.04
21	Kelly's ratio	0.04	0.50	0.16	0.09	0.01
22	Non-carbonate hardness (ppm)	-51.29	195.82	29.90	46.55	7.36

none of the cation and anion pairs exceed 50%. About 20% samples are falling in the area of 6 which indicates that non-carbonate hardness (secondary salinity) exceeds 50%. The Piper diagram confirms that all the groundwaters in the study area are characterized as alkaline earth's (Ca + Mg) exceeds alkalies (Na + K) and weak acids (CO₃ + HCO₃) exceed strong acids (SO₄ + Cl + F). This is due to the dolomitic rocks which are responsible for release of chemical elements into the groundwaters of the Tummalapalle area.

Chadhas diagram

The groundwater of the study area has been classified as per Chadha's diagram (Chadha 1999). This diagram is a somewhat modified version of the Piper diagram (Piper 1944) and the expanded Durov diagram (Durov 1948). In Chadha's diagram, the difference in milliequivalent percentage between alkaline earths ($\operatorname{Ca}^{2+} + \operatorname{Mg}^{2+}$) and alkali metals ($\operatorname{Na}^+ + \operatorname{K}^+$), expressed as percentage reacting values, is plotted on the *X* axis and the difference in milliequivalent percentage between weak acidic anions ($\operatorname{CO_3}^{2-} + \operatorname{HCO_3}^-$) and strong acidic anions ($\operatorname{CI}^- + \operatorname{SO_4}^{2-}$) is plotted on the *Y* axis (Fig. 3). The square rectangular field describes the overall character of the water.

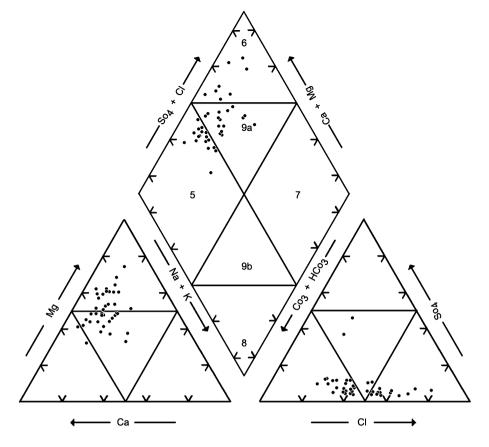
The diagram has all the advantage of the diamond-shaped field of the piper trilinear diagram and can be used to study various hydrochemical processes, such as base cation exchange, cement pollution, mixing of natural waters, sulphate reduction, saline water (end product water) and other related hydrochemical problem (Jain et al. 2001). In the present study, all the samples are falling in the field of 5 which belongs to the Ca²⁺–Mg²⁺–HCO₃⁻ type and such water has temporary hardness (Fig. 3).

Statistical analysis

Correlation between the different parameters of ground-water of Tummalapalle showed both positive and inverse relations between the parameters, some moderately correlated and some well correlated (Table 2). Highest positive correlation was observed among these parameters: hardness and Mg (1.00); hardness and CO₃ (0.905); hardness and HCO₃ (0.869); HCO₃ and CO₃ (0.968); Mg and CO₃ (0.900); Mg and HCO₃ (0.865). The high correlations among these parameters indicate that carbonates dissolution may be the dominant water–rock interaction occurring in the aquifer (Li et al. 2013a), which can be expressed by R1 and R3. Apart from these highly correlated parameters, there were other moderate correlations between other



Fig. 2 Trilinear diagram for representing the analyses of groundwater quality (Piper diagram)



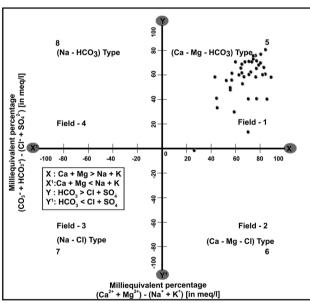


Fig. 3 Chadha's diagram (modified piper diagram)

parameters, such as Cl and Na (0.739), HCO₃ and Na (0.418), and F and Ca (0.378). The positive correlation between Cl and Na suggests that halite weathering (R3) is one of the reasons responsible for the concentration of Na and Cl in groundwater (Li et al. 2016a). Similarly, the

positive correlation between HCO_3 and Na indicates that the dissolution of albite, expressed in R4, is a possible origin of HCO_3 and Na in groundwater (Li et al. 2016b). The moderate correlation between F and Ca may be attributed to the dissolution of fluorite (Li et al. 2014b; Wu et al. 2015). Overall, the major chemical elements present in the groundwaters of the Tummalapalle area are due to the dolomitic rocks occurring in the study area.

$$CaCO_3 + CO_2 + H_2O \rightarrow Ca^{2+} + 2HCO_3^-$$
 (R1)

$$CaMg(CO_3)_2 + 2CO_2 + 2H_2O$$

 $\rightarrow Mg^{2+} + Ca^{2+} + 4HCO_3^-$ (R2)

$$NaCl \rightarrow Na^+ + Cl^-$$
 (R3)

$$CaF_2 = Ca^{2+} + 2F^-.$$
 (R5)

Drinking water quality assessment

Total dissolved solids (TDS) and total hardness (TH) are two important parameters in assessing drinking water quality. TDS represents the total weight of dissolved solids



Table 2 Correlation matrix for groundwater samples of Tummalapalle area

	EC	pН	Ca	Mg	Na	K	HCO ₃ CC	O ₃ Cl	SO ₄	F	TDS	Hardness Alkalinity
EC	1											
pН	0.488^{*}	* 1										
Ca	0.122	0.043	5 1									
Mg	-0.151	-0.020	0.455**	* 1								
Na	0.128	-0.248	8 0.300	0.303	1							
K	-0.157	-0.23	1 0.028	0.306	0.361^{*}	1						
HCO_3	-0.031	0.00	7 0.695**	0.865**	0.418**	0.392*	1					
CO_3	-0.042	0.02	7 0.678**	* 0.900**	0.391^{*}	0.308	0.968** 1					
Cl	-0.049	-0.29	1 0.544**	* 0.475**	0.739**	0.361*	0.460** 0.4	34** 1				
SO_4	0.142	0.179	9 0.097	0.221	-0.219	-0.157	0.010 0.1	-0.03	108 1			
F	0.176	0.046	5 0.378*	-0.086	0.128	-0.083	0.132 0.0	0.83	-0.232	1		
TDS	0.359^{*}	0.094	4 0.318*	-0.019	0.327^{*}	0.136	0.174 0.1	21 0.	241 -0.002	0.346*	1	
Hardness	-0.151	-0.020	0.467**	* 1.000**	0.303	0.304	0.869** 0.9	05** 0.	478** 0.222	-0.082	-0.017	1
Alkalinity	0.440**	* 0.15	7 0.298	0.031	0.315^{*}	0.440*	* 0.309 0.1	53 0.	255 -0.171	0.197	0.456**	0.032 1

^{**} Correlation is significant at the 0.01 level (two-tailed)

^{*} Correlation is significant at the 0.05 level (two-tailed)

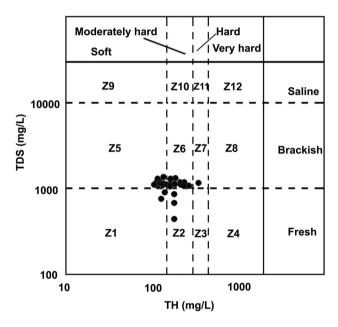
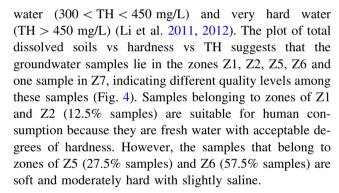


Fig. 4 Plot of TDS vs TH expressed in mg/l as CaCO₃

in a solution and expresses the degree of salinity of a medium (Mitra et al. 2007). According to the level of TDS, groundwater can be classified as fresh groundwater (TDS < 1000 mg/L), brackish water (1000 > TDS < 10000 mg/L) and saline water (TDS > 10000 mg/L) (Wanda et al. 2011). Hardness is a measure of dissolved Ca^{2+} and Mg^{2+} in water and is expressed as $CaCO_3$. According to the grading standards of TH (as $CaCO_3$), groundwater can be classified as soft water (TH < 150 mg/L), moderately hard water (150 < TH < 300 mg/L), hard



Irrigation water quality

Sodium adsorption ratio (SAR)

Sodium hazard is also usually expressed in terms of the sodium adsorption ratio (SAR), which is calculated from the ratio of sodium to calcium and magnesium (Todd and Mays 2005; Li et al. 2013a). The SAR is an important parameter for the determination of the suitability of irrigation water because it is responsible for the sodium hazard (Li et al. 2016c, d; Todd and Mays 2005). The waters were classified in relation to irrigation based on the ranges of SAR values (Richards 1954). Continued use of water having a high SAR leads to a breakdown in the physical structure of the soil. Sodium is adsorbed and becomes attached to soil particles. The soil then becomes hard and compact when dry and increasingly impervious to water



penetration. The degree to which irrigation water tends to enter into cation-exchange reactions in soil can be indicated by the sodium adsorption ratio. Sodium replacing adsorbed calcium and magnesium is a hazard as it causes damage to the soil structure.

SAR is computed as

$$SAR \ = \ \frac{Na}{\sqrt{Ca \ + \ Mg/2}}$$

where all ionic concentration is expressed in meq/l.

The SAR values of the study area are presented in Table 1 and are varying from 0.15 to 1.54. Groundwater could be also classified based on sodium adsorption ratio (SAR) as excellent (10), good (10–18), doubtful (18–26) and unsuitable (>26) (Sadashivaiah et al. 2008). A high SAR in irrigation water has the potential to impair soil structure and thus the permeability of the soil leading to a lack of soil moisture (Compton 2011). The sodium adsorption ratio (SAR) greater than 12.0 is considered sodic and threatens the survival of vegetation by increasing soil swelling (dispersion) and reducing soil permeability (Kuipers et al. 2004).

Integrated effect of EC and SAR

Richard's classification of sodium adsorption ratio proposes that the groundwaters of the study area are excellent for irrigation. The graphical diagram of irrigated water (US Salinity Laboratory 1954) and the integrated effect of EC and SAR conclude that the groundwaters of the study area reveal that the groundwaters possess high salinity with low sodium. Further, these waters are rated as satisfactory for agricultural purposes. SAR is plotted against EC, which is designated as after US Salinity Laboratory 1954 and is illustrated in Fig. 5. In the present study, all the water samples are falling under high specific conductance and low sodium [C3S1] as per US Salinity Laboratory (1954). The excess sodium in waters produces the undesirable effects of changing soil properties and reducing soil permeability (Kelly 1951). Hence, the assessment of sodium concentration is necessary while considering the suitability for irrigation. The degree to which irrigation water tends to enter into cation-exchange reactions in soil can be indicated by the sodium adsorption ratio (US Salinity Laboratory 1954). Sodium replacing adsorbed calcium and magnesium is a hazard as it causes damage to the soil structure. It becomes compact and impervious.

Adjusted sodium adsorption ratio (adj.SAR)

The high concentration of sodium in irrigation water may negatively affect the soil structure and decrease the soil hydraulic conductivity in fine-textured soil. The degree to

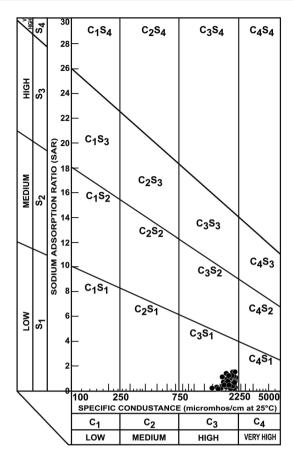


Fig. 5 Quality of groundwater samples in relation to salinity and sodium hazard (after US Salinity Laboratory 1954)

which sodium will be absorbed by a soil is a function of the amount of sodium to divalent cations (Ca²⁺ and Mg²⁺) and is regularly stated by the sodium adsorption ratio (SAR) (Bouwer and Idelovitch 1987). This parameter is basically used for assessment of alkalinity hazard in irrigation water and it is ranging from 0.33 to 4.59 (Table 1). The result showed that the concern due to sodium hazard of the water became more emphatic because in all water samples adj.SAR is higher than SAR (Suarez 1981; Lesch and Suarez 2009).

This can be calculated with the following formula (Ayers and Westcot 1985):

Adjusted sodium adsorption ratio (adj.SAR) = SAR [1 + (8.4 - pHc)]

$$pHc = (pK2 - pKc) + p(Ca + Mg) + p(Alk)$$

where pK2 = negative logarithm of the second disassociation constant for carbonic acid, pKc = solubility constant for calcite, p = negative logarithm of ion concentration. All ionic concentration is in meq/L.

Based on Ayers and Tanji (1981) classification, majority of samples have adj.SAR values are showing <3 and are safe for irrigation. The minimum value of adj.SAR is 0.33 and maximum value is 4.59 and average is about 1.79.



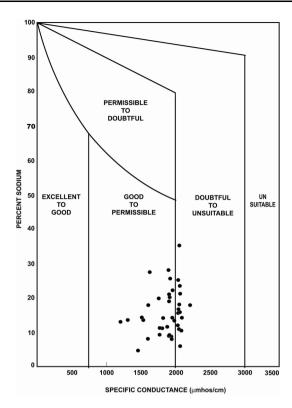


Fig. 6 Quality of water in relation to electrical conductivity and per cent sodium (Wilcox diagram)

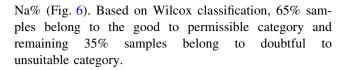
Sodium percentage

Sodium concentration is important in classifying irrigation water. Soils containing a large proportion of sodium with carbonate as the predominant anion are termed alkali soils; those with chloride or sulphate as the predominant anion are saline soils. The role of sodium in the classification of groundwater for irrigation was emphasized because of the fact that sodium reacts with soil and as a result clogging of particles, thereby reducing the permeability (Domenico and Schwartz 1990; Todd and Mays 2005). When the concentration of Na⁺ is high in irrigation water, Na⁺ tends to be absorbed by clay particles displacing Mg²⁺ and Ca²⁺ ions. This exchange process of Na⁺ in water for Ca²⁺ and Mg²⁺ in soil reduces the permeability and eventually results in soil with poor internal drainage. Hence, the assessment of sodium concentration is necessary while considering the suitability for irrigation. The Na% is calculated using the formula given as follows:

$$\% Na \, = \, \frac{(Na \, + \, K)}{(Ca \, + \, Mg \, + \, Na \, + \, K)} \, \times \, 100$$

where all ionic concentrations are expressed in meq/l.

In the present study, the per cent sodium values ranges from 4.73 to 35.03 with a mean value of about 15.86 (Table 1). Wilcox's (1955) diagram is adopted for classification of irrigation, wherein the EC is plotted against



Potential salinity

Potential salinity is defined as the chloride concentration plus half of the sulphate concentration. Doneen (1954) explained that the suitability of water for irrigation is not dependent on soluble salts. It is true that the low solubility salts may precipitate in the soil and accumulate with each successive irrigation, whereas the concentration of highly soluble salts increases the soil salinity (Doneen 1962). The potential salinity of the water samples ranges from 0.72 to 6.13 (Table 1).

Residual sodium carbonate (RSC)

It is another parameter used to classify groundwater for irrigation purposes (Siddiqui et al. 2005). The RSC in groundwater is mainly due to the higher concentration of bicarbonate ions, which precipitates Ca²⁺ and Mg²⁺ ions as their carbonates and elevates Na⁺ ions, which increases the sodium carbonate in the groundwater (Srinivas et al. 2014). In addition to the total dissolved solids, the relative abundance of sodium with respect to alkaline earths and boron and the quantity of bicarbonate and carbonate in excess of alkaline earths also influence irrigation water quality. This excess is denoted by residual sodium carbonate and determined as suggested by Richards (1954). The water with high RSC has high pH and land irrigated by such waters becomes infertile owing to deposition of sodium carbonate as indicated by the black colour of the soil (Eaton 1950). In waters having high concentration of bicarbonate, there is tendency for calcium and magnesium to precipitate as the water in the soil becomes more concentrated. As a result, the relative proportion of sodium in the water is increased in the form of sodium carbonate.

RSC is calculated using the following formula:

$$RSC = (HCO3 + CO3) - (Ca + Mg) (meq/l)$$

Lloyd and Heathcoat (1985) have classified irrigation water based on RSC as (1) suitable (<1.25) (2) marginal (1.25–2.5) and (3) not suitable (>2.5). In the present study, from Table 1, it can be interpreted that the groundwaters in the study area shows RSC values of ranging from -3.92 to 1.03 meq/l. Based on RSC values, all 40 samples have values less than 1.25 and are safe for irrigation.



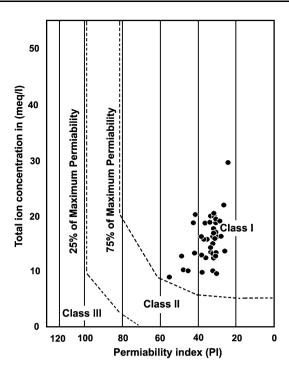


Fig. 7 Doneen's diagram of groundwater quality

Permeability index (PI)

The soil permeability is affected by the long-term use of irrigation water as it is influenced by Na⁺, Ca²⁺, Mg²⁺ and HCO₃⁻ content of the soil. Doneen (1964) and WHO (2006) gave a criterion for assessing the suitability of groundwater for irrigation based on the PI, where concentrations are in meq/l.

$$PI = (Na^{+} + \sqrt{HCO_{3}^{-}} \times 100/(Ca^{2+} + Mg^{2+} + Na^{+})$$

The analytical data are plotted on the Doneen's chart (Fig. 7). All samples were plotted in Class I zone, indicating that all groundwater samples were suitable for irrigation in terms of the PI (Domenico and Schwartz 1990).

Kelly's ratio

Groundwater was also classified for irrigation based on Kelly's ratio in this study. Sodium measured against calcium and magnesium was considered by Kelly (1940) and Paliwal (1967) to calculate this parameter. A Kelly's ratio of more than one indicates excessive sodium in waters, which is not suitable for irrigation. Kelly's ratio can be calculated as follows:

$$\{\text{Kelly's ratio} = \text{Na}/(\text{Ca} + \text{Mg})\}$$

where all the ionic constituents are expressed in meq/l.

From Table 1, it can be suggested that the Kelly's ratio varies from 0.04 to 0.50, demonstrating that all water samples are suitable for irrigation.

Sources of major ions

Correlation analysis and bivariate diagrams are helpful in understanding the processes of groundwater chemical evolution (Wu et al. 2014; Tiwari and Singh 2014). In this study, bivariate diagrams were generated based on Pearson correlation analysis to get a better understanding of the groundwater chemical evolution processes, and the results are shown in Fig. 8.

The correlation and trend line analyses between the groundwater constituents are exhibited in Fig. 8. The existence of a very strong correlation exist between HCO_3^- and $Ca^{2+} + Mg^{2+}$ (r = 0.92); between $HCO_3^- + SO_4^{2-}$ and $Ca^{2+} + Mg^{2+}$ (r = 0.95); total cations and Ca^{2+} and Mg^{2+} (r = 0.96). There is a strong relation between HCO_3^- and $Ca^{2+} + C_3^{2-} + C_3^$

The plot of $Ca^{2+} + Mg^{2+}$ vs HCO_3^- shows a abundance of $Ca^{2+} + Mg^{2+}$ relative to HCO_3^- , and Fig. 8b shows that the abundance of (Ca + Mg) in most of the groundwater samples probably can be attributed to carbonate weathering. The climate also plays a vital role in the arid and semi arid areas. The plots between total cations and Ca + Mg as well as Na + K are 0.96 and 0.40, respectively (Fig. 8d, e).

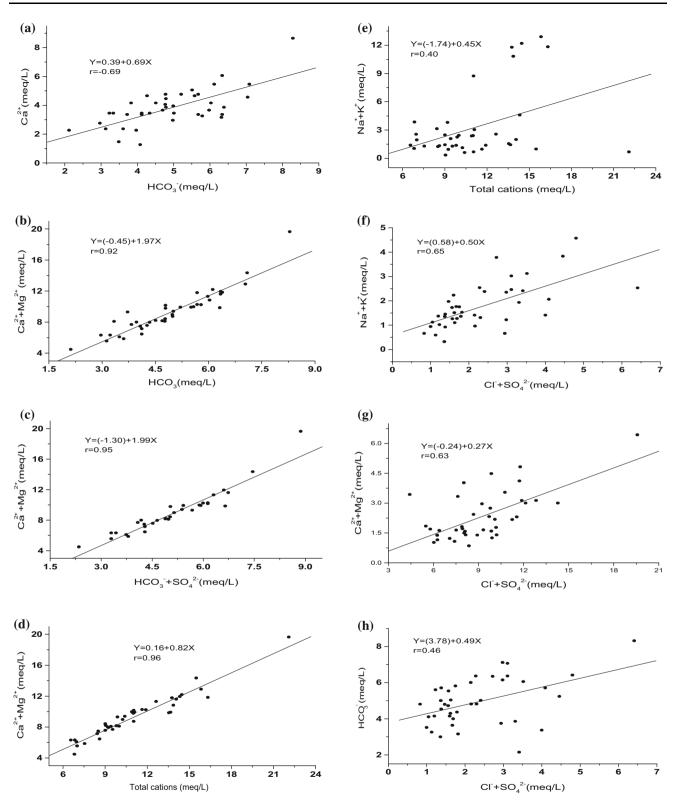
The chemical composition of groundwater is the imprints of the rock-water interaction and chemical processes. Generally, the presence of rock-water interaction was identified using TDS vs Na/(Na + Ca) and TDS vs Cl/(Cl + HCO₃) scatter diagrams as reported by Gibbs (1970) (Fig. 9). These diagrams are widely employed to assess the functional sources of dissolved chemical constituents, such as precipitation, rock, and evaporation dominance. In these diagrams, samples falling in the centre of the curve are derived from rock-water interaction. Figure 9 presents Gibbs plots of Group (A) and (B) groundwater quality data of Tummalapalle area.

This observation showed the involvement of silicate weathering in the geochemical processes, which contribute mainly sodium, calcium and potassium ions to the groundwater (Stallard and Edmond 1983; Sarin et al. 1989).

Conclusions

The groundwater in Tummalapalle area is alkaline in nature. In majority of the groundwater samples, concentrations of alkaline earths $(Ca^{2+} + Mg^{2+})$ exceed alkali



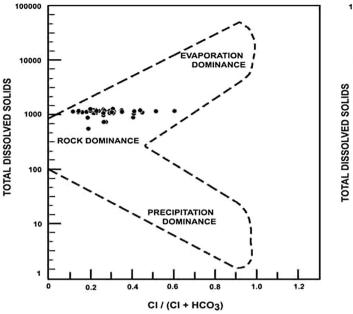


 $\begin{array}{l} \textbf{Fig. 8} \ \ \text{Scatter plots between. a HCO}_3^- \ \text{vs Ca}^{2+}; \ \textbf{b HCO}_3^- \ \text{vs Ca}^{2+} + \ \text{Mg}^{2+}; \ \textbf{c HCO}_3^- + \ \text{SO}_4^{\,2-} \ \text{vs Ca}^{2+} + \ \text{Mg}^{2+}; \ \textbf{d total cations vs Ca}^{2+} + \ \text{Mg}^{2+}; \ \textbf{e total cations vs Na}^+ + \ \text{K}^+; \ \textbf{f Cl}^- + \ \text{SO}_4^{\,2-} \ \text{vs Na}^+ + \ \text{K}^-; \ \textbf{g Cl}^- + \ \text{SO}_4^{\,2-} \ \text{vs Ca}^{2+} + \ \text{Mg}^{2+}; \ \textbf{h Cl}^- + \ \text{SO}_4^{\,2-} \ \text{vs HCO}_3^- \end{array}$

cations $(Na^+ + K^+)$ and HCO_3^- dominate over $(SO_4^{\ 2^-} + Cl^-)$. Ca-Mg-HCO $_3$ is the dominant hydrogeochemical facies. The statistical analyses were used for

determining the groundwater quality variations. Water chemistry of the study area strongly reflects the dominance of weathering of rock-forming minerals. The Gibbs





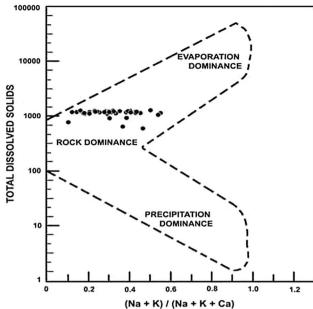


Fig. 9 Mechanism controlling the quality of groundwater

diagram revealed that the hydrochemistry of groundwater falls in the rock weathering region and is due to dissolution with rock-forming minerals. Pearson correlation analysis reveals that natural processes such as mineral dissolution/precipitation and cation exchange are dominant factors influencing the groundwater chemistry. The high contribution of $(Ca^{2+} + Mg^{2+})$ to the total cations suggest that the chemical composition of the water is largely controlled by bicarbonate weathering with limited contribution from silicate weathering. The Piper diagram has revealed that all the samples are characterized as carbonate hardness (secondary alkalinity) exceeds 50% and are due to the dolomitic rocks which are responsible for release of chemical elements into the groundwaters of the study area.

In this study, the assessment of groundwater for irrigational uses has been evaluated on the basis of various parameters. Sodium adsorption ratio (SAR) values are to be less than 10; and adj. SAR also less than 3; Residual sodium carbonate (RSC) values on the whole are less than 1.25 meg/l; Permeability index (PI) are also in acceptable range for irrigation. All these parameters indicating the water samples of the study are good for irrigation. According to the as per USSL diagram all water samples are falling under very high specific conductance and low sodium [C3S1] and are suitable for irrigation. The Wilcox classification has shown 65% of groundwater under "good to permissible" zone. Thus, the overall groundwater quality in the basin is fresh and suitable for irrigation use. The plots of dissolved solids (TDS) and total hardness (TH) suggest that about 12.5% of the samples are suitable for human consumption because they are fresh water with acceptable degrees of hardness. Suitable water treatment process such as water softening, ion exchange, and demineralization should be applied to reduce the concentration of contaminants. The various indices derived in the study indicate that the most of groundwater of the study area is suitable for agriculture irrigation use. The long-term use of such groundwater for irrigation will induce sodium hazard to soils. It will have negative impacts on the yields of crops and properties of soils. However, mixing of low and high salinity water is recommended before irrigation to reduce the salinity hazard in local areas.

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