



Impact of rock type on the pore structures and physical properties within a tight sandstone reservoir in the Ordos Basin, NW China

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Abstract

The pore–throat systems and physical properties of tight sandstone reservoirs are complex, and deposition is thought to be a fundamental control for them. In this study, the impacts of the full ranges of rock types (from pebbly coarse sandstone to fine sandstone) on the pore structures and physical properties of the Permian tight sandstone reservoir in the eastern Ordos Basin were investigated comprehensively through a series of experiments including conventional physical testing, thin-section analysis, scanning electron microscopy, nuclear magnetic resonance analysis and high-pressure mercury injection tests. The results showed that the coarser-grained sandstones tend to have higher feldspar content and lower percentage of cements, leading to strong dissolution, weak cementation and improved porosity and permeability. The medium sandstone has the highest level of quartz and the lowest average content of feldspar, resulting in strong heterogeneity of physical properties. Only those medium sandstone reservoirs with relatively high content of feldspars have better physical properties. Additionally, the coarser-grained sandstones contain relatively large dissolution pores (nearly 200 μm), whereas the finer-grained sandstones have more intercrystalline pores with a relatively more homogeneous pore structure. The pebbly coarse sandstone and coarse sandstone reservoirs are favorable targets with best physical properties.

Keywords Pore size distribution · Pore types · Physical properties · Tight sandstone · Ordos Basin

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1 Introduction

Compared to shale gas, tight sandstone gas is more practical and significant as an unconventional resource because tight sandstone can be hydraulically fractured more easily than shale (Song et al. 2013). The proven reserve of tight sandstone gas in China is approximately $3.5 \times 10^{12} \text{ m}^3$ with an annual production of $169 \times 10^8 \text{ m}^3$, demonstrating a great exploration potential (Zou et al. 2013). Hydraulic fracturing is an important technique for tight sandstone reservoirs (Guo et al. 2015; Feng et al. 2016; Zhao et al. 2016; Wang et al. 2017a, b, c; Tian et al. 2018). Nevertheless, in most basins of the world, reservoir “sweet spots” exist and searching for these “sweet spots” before hydraulic fracturing is crucial for successful exploration and development of tight sandstone gas (Zou et al. 2014; Sun et al. 2015, 2019; Wang et al. 2015; Tian et al. 2016; Zhang et al. 2016; Tang et al. 2017; Zhao et al. 2017). A “sweet spot” is defined as that has better physical properties, such as porosity and permeability. Essentially, pores and throats are controlling factors for reservoir physical properties (Li et al. 2017a, b; Huang et al. 2018). Multiple conventional and new methods have

been applied to characterize the types, sizes and connectivity of pores, as well as the geometry and sizes of throats within tight sandstones. These methods include geochemistry (Sonnenberg and Pramudito 2009; Bjørlykke and Jahren 2012; Clarkson et al. 2012a, b), thin-section analysis (Lame et al. 2004), scanning electron microscopy (Anovitz and Cole 2015; Xi et al. 2016; Zhou et al. 2016; Chen et al. 2017), high-pressure mercury intrusion (Xiao et al. 2011; Clarkson et al. 2012a, b; Shanley and Cluff 2015; Shao et al. 2017), nuclear magnetic resonance (Lai et al. 2016, 2018a; Li et al. 2017a, b; Xiao et al. 2018a, b; Xiao et al. 2018; Zhu et al. 2018;), nitrogen adsorption (Schmitt et al. 2015; Liu et al. 2017) and small angle neutron scattering (Anovitz et al. 2013; Anovitz et al. 2015).

The pore–throat systems and physical properties of tight sandstone reservoirs are complex, and deposition was thought to be a fundamental control for them. Diagenesis exerts more important controls on the pore–throat systems and physical properties. There is ample evidence that diagenesis controls pore structures, as well as the porosity and permeability of tight sandstones (Anovitz et al. 2013, 2015; Lai et al. 2015; Xi et al. 2015; Wang et al. 2017a, b, c; Patrick et al. 2017; Wang et al. 2017a, b, c; Yu et al. 2019). In addition, deposition plays a fundamental and innate role in porosity evolution (Yu et al. 2015). The rock types, mineral composition and the size and arrangement of grains further influence the diagenetic processes and the formation of natural fractures. For example, Wang et al. (2017a, b, c) found that coarse sandstone was prone to have a relatively weak compaction and strong dissolution due to the presence of more rigid grains and that fine sandstone generally undergoes strong compaction and cementation because of more plastic grains and high matrix content (Lai et al. 2018a, b). It has also been suggested that finer grain size is related to more matrix filling, resulting in more intense mechanical compaction, significant porosity loss and more intragranular micropore development within fine-grained tight sandstone (from fine sandstone to very fine sandstone and to coarse silt) (Shao et al. 2017).

Nevertheless, the previous studies focused on fine-grained sandstone were mostly qualitative, using thin-section methods and did not systematically reveal the lithologic controls on rock properties within the full ranges of rock types. The Shihezi Formation is presently the most significant reservoir of tight sandstone gas in the eastern Ordos Basin. It contains multiple rock types from pebbly coarse sandstone to siltstone, which lead to a diversity of pore–throat systems and strong heterogeneity of reservoir quality (Yue et al. 2018). This study aimed to investigate the lithological control on the pore structures and the quality of tight sandstone reservoir systematically using multiple methods, including thin-section analysis, scanning electron microscopy, nuclear

magnetic resonance (NMR) and high-pressure mercury intrusion (HPMI) in the study area.

2 Geological setting

The study area is located in the northwestern margin of the Jinxi Fault-fold Belt in the Ordos Basin, which is the second largest sedimentary basin with vast oil, gas and coal in China (Fig. 1). The northern Ordos Basin was uplifted due to Hercynian movement during the Late Paleozoic and provided sediments for the basin fill. The corresponding sedimentary sequence consists of the Upper Carboniferous Benxi Sequence (C₃b), the Lower Permian Taiyuan Sequence (P₁t) the Shanxi Sequence (P₁s), the Middle Permian Shihezi Sequence (P₂h) and the Upper Permian Shiqianfeng Formation (P₃q) (Wang et al. 2008; Jia et al. 2017).

The Shihezi Formation (P₂h), primarily containing braided river delta sediments, is the target interval. P₂h is subdivided into eight parts. The P₂h sandstone is tight, with an average porosity and permeability of 7.75% and 0.77 mD within the producing layers. Nevertheless, sweet spots with relatively high physical properties exist in the underwater distributary channels and river mouth bar. The P₂h reservoir has strong heterogeneity, as illustrated by its porosity and permeability, as well as varied pore types and structures observed in thin sections. Correspondingly, some wells yield a high gas flow, whereas others show little or no gas. The coal bed and shale within the P₂s, P₁t and C₃b are source rocks, which began to generate hydrocarbon at the end of the Triassic. The period from the Late Jurassic to the Early Cretaceous was the hydrocarbon expulsion peak and also was the primary reservoir-forming stage. The tectonic uplift that occurred during the stage from the Late Cretaceous to the present retarded the thermal evolution of the source rocks (Xie et al. 2016).

3 Samples and methods

According to the research objectives, samples from 16 wells covering a large range of lithologies from pebbly coarse sandstone to fine sandstone were collected from the P₂h tight gas sandstones in the eastern Ordos Basin, NW China. Particularly, 24 samples from 12 wells were measured by HPMI and proximal samples with the same depth were tested by NMR, in addition to other experiments (Table 1). The HPMI and NMR results were used to investigate the lithologic control on the pore structure and physical properties of the tight sandstones.

Routine core testing to measure the porosity and permeability of the P₂h tight sandstone was performed on 72 samples at 25 °C and 35–50% RH. A PoroTm300 porosimeter

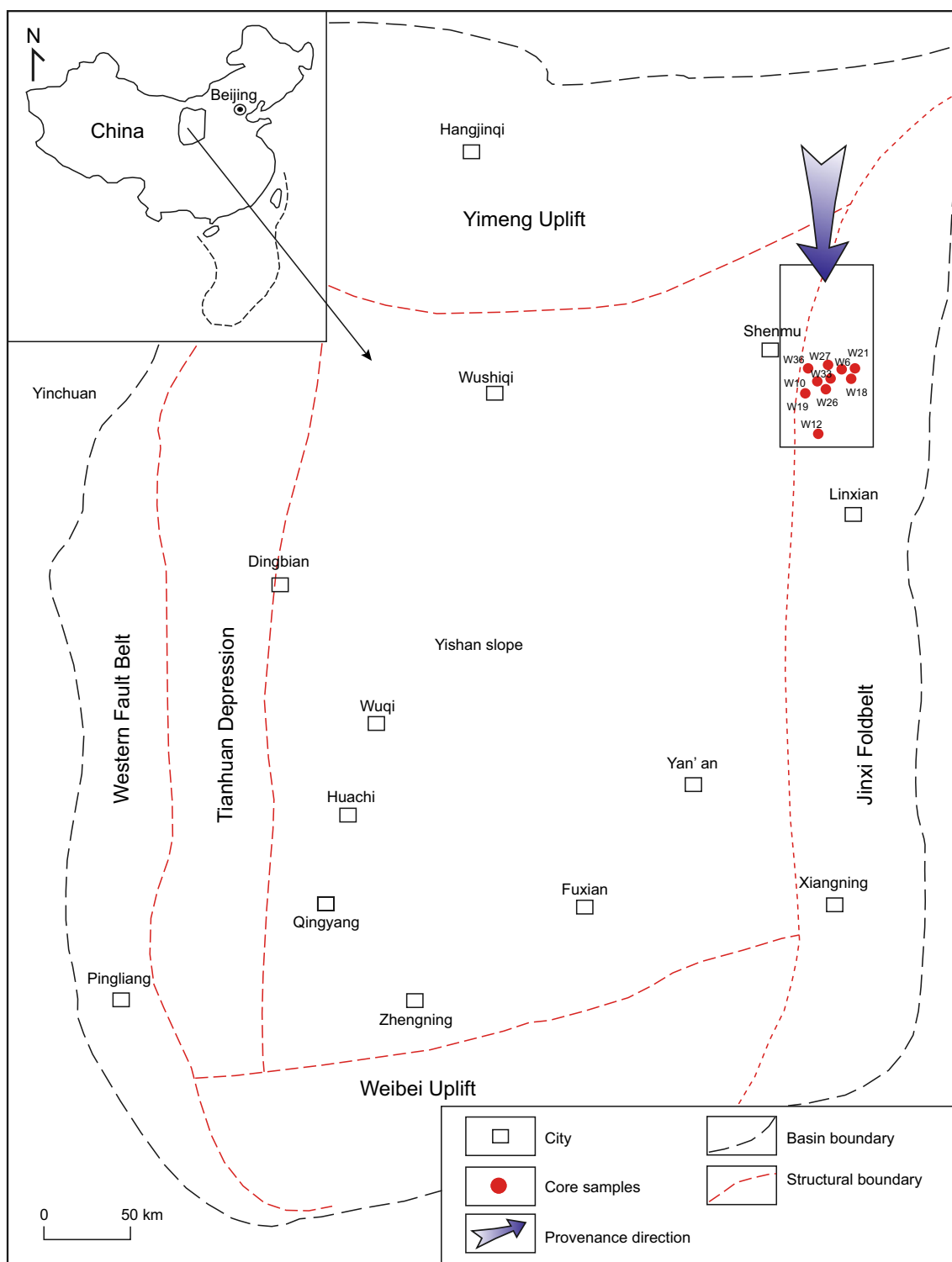


Fig. 1 Map of the Ordos Basin and location of the study area

and a low perm-meter with flowing air were used to obtain the porosity and permeability of these samples, respectively.

Seventy-two samples from 20 wells were cut to 2.54 cm × 1.5 cm × 0.5 cm volume for observation with

the polarizing microscope. The thin-section chips were then casted with blue stain to highlight the pores (calcite stained with Alizarin Red S) using an LJS-3 casting machine and ground to thin sections by SPM-300

Table 1 Representative samples and their primary parameters from HMPI

Sample	Depth, m	Lithology	Facies	Porosity, %/avg		Permeability, mD/avg		Displacement pressure, MPa/avg		Maximum throat radius, nm/avg	
W19-1	1641.12	Pebbly coarse sandstone	Distributary channel	9.9	9.9	0.54	9.60	0.72	0.96	1.04	5.63
W21-1	1654.34	Pebbly coarse sandstone	Channel lag	16.0		36.72		0.04		20.06	
W33-1	1371.67	Pebbly coarse sandstone	Channel lag	4.1		0.67		0.68		1.10	
W33-2	1539.1	Pebbly coarse sandstone	Channel lag	9.5		0.52		2.40		0.31	
W21-2	1757.78	Coarse sandstone	Under water distributary channel	15.2	14.6	4.85	3.06	0.20	0.36	3.81	2.42
W10-1	1480.4	Coarse sandstone	Distributary channel	18.6		2.90		0.39		1.92	
W19-2	1521.27	Coarse sandstone	Under water distributary channel	12.8		3.59		0.29		2.60	
W21-3	1759.84	Coarse sandstone	Under water distributary channel	11.6		0.91		0.56		1.35	
W12	1791.13	Pebbly medium sandstone	Under water distributary channel	4.8	9.6	4.09	2.26	0.31	1.24	2.41	1.00
W18	1762.81	Pebbly medium sandstone	Under water distributary channel	5.6		0.08		2.28		0.33	
W33-3	1360.83	Pebbly medium sandstone	Channel lag	18.1		3.80		1.12		0.67	
W36	1656.85	Pebbly medium sandstone	Under water distributary channel	10.0		1.06		1.26		0.60	
W6-1	1646.85	Medium sandstone	Under water distributary channel	6.0	8.7	0.39	1.74	1.02	0.92	0.72	0.92
W10-3	1484.1	Medium sandstone	Distributary channel	7.3		0.22		1.21		0.62	
W10-2	1213.6	Medium sandstone	Distributary channel	5.1		0.33		0.98		0.76	
W35	1442.47	Medium sandstone	Distributary channel	16.6		6.02		0.47		1.57	
W30-1	1383.84	Pebbly fine sandstone	Distributary channel	11.5	10.5	2.30	1.57	0.83	0.91	0.90	1.16
W26-1	1603.1	Pebbly fine sandstone	Distributary channel	9.8		1.33		0.77		0.98	
W26-2	1604.76	Pebbly fine sandstone	Distributary channel	10.5		1.84		0.65		1.16	
W27	1547.75	Pebbly fine sandstone	Under water distributary channel	10.1		0.79		1.39		0.15	
W6-2	1470.81	Fine sandstone	Distributary channel	3.0	8.5	0.23	0.51	0.33	1.09	2.22	1.20
W21-4	1651.12	Fine sandstone	Point bar	11.3		0.16		2.21		0.34	
W30-2	1447.37	Fine sandstone	Distributary channel	11.2		1.13		0.73		1.03	

W26-1, W26-2, and W6-2 have no corresponding NMR test

double-deck grinding machine. Finally, petrologic investigation of these samples was conducted using a DMLP polarizing microscope at the test temperate and humidity of 25 °C and 40%, respectively. The detailed procedure was referred from the China Petroleum and Natural Gas Industry Standard SY/T 5913-2004. In addition, 24 samples, each weighting 10–12 g, were crushed to radii smaller than 1 mm. Then, a D/max-2200 powered X-ray diffractometer was employed for X-ray diffraction testing (XRD) of mineral compositions, and the absolute and relative contents of clay minerals of these samples following SY/T 5163-2010.

Field-emission scanning electron microscopy (FE-SEM) analysis of 72 samples was performed using a JSM-5500LV scanning electron microscope, equipped with a QUANTAX400 energy-dispersive X-ray spectrophotometer following SY-T5162-1997. Small samples were cut from large core samples (diameter = 1 mm, thickness = 1 mm) and then were analyzed at a temperature of 25 °C and humidity level of 40%. The SEM micrographs, with magnifications ranging from 100 to 15,000 under

20 kV, could provide critical information on the size, morphology and even connectivity of pores.

The pore–throat distribution was investigated by HMPI using an AutoPore IV9500 mercury porosimeter at 18.9–19.8 °C following SY/T 5346-2005. Twenty-four samples from 12 wells (diameter = 2.53 cm, length = 2.5 cm), with six diverse lithologies (from pebbly coarse sandstone to fine sandstone), were measured. Pore–throat structure was characterized according to the patterns of the mercury intrusion and extrusion curves. The permeability contribution distribution was also acquired from the fluid flow capacity for different pore–throat radii using Eq. (1). The method can be specifically followed the work of Wall (1965).

$$c = \frac{(2i - 1) \times r_n^2}{\sum_{i=1}^n (2i - 1) \times r_n^2} \times 100 \tag{1}$$

In Eq. (1), *c* is the permeability contribution of pores with the same radius, *n* is the group of pores divided on the basis of pore size, and *r_n* is the pore radius (μm).

Low-field NMR technology is now used increasingly to characterize the pore size distribution (PSD) within unconventional reservoirs. The NMR T_2 relaxation time can be applied to indicate the PSD of tight sandstone. T_2 is controlled by three relaxation mechanisms, the surface, bulk and diffusion relaxations.

The bulk relaxation can be neglected because it takes a longer amount of time than the surface relaxation for nonviscous kerosene. The diffusion relaxation can also be ignored in a low and uniform magnetic field. Thus, the NMR T_2 relaxation time is primarily determined by the surface relaxation (Eq. 2):

$$\frac{1}{T_2} = \rho_2 \frac{S}{V} \quad (2)$$

Twenty-two samples of six different lithologies (diameter = 2.54 cm, length = 2.54 cm) were selected for NMR T_2 distribution determination after crude oil and salt contamination were removed using an NMR rock core analyzer. The NMR parameters were as follows: waiting time, 3 s; echo spacing, 0.6 ms; number of scans, 256; echo numbers, 1024; water salinity, 49,600 mg/g, experimental temperature 35 °C.

4 Results and discussion

4.1 Rock types, mineral compositions and pore systems

The bulk-rock mineralogy obtained by thin-section observation demonstrated that samples of P₂h were mainly feldspathic litharenites and lithic arkoses (Fig. 2). The upper Shihezi Formation had less feldspar and more rock fragments than the lower Shihezi Formation. As shown in Fig. 3, quartz, metamorphic rock fragments and potash feldspar were three the primary compositions within the P₂h reservoir. Nevertheless, as illustrated in Fig. 4, the pebbly sandstones (pebbly coarse sandstone, pebbly medium sandstone and pebbly fine sandstone) contained less quartz and more metamorphic rock fragments, volcanic rock fragments and plagioclase than the non-pebbly sandstones (coarse sandstone, medium sandstone and fine sandstone). Pebbly coarse sandstone had the largest content of feldspar (plagioclase and potash feldspar), whereas medium sandstone had the lowest feldspar but highest quartz content. Among the pebbly sandstones, pebbly fine sandstone had the more quartz and metamorphic rock fragments, but less plagioclase. The non-pebbly sandstones showed similar compositions, except that volcanic rock fragments increase as the lithology became finer.

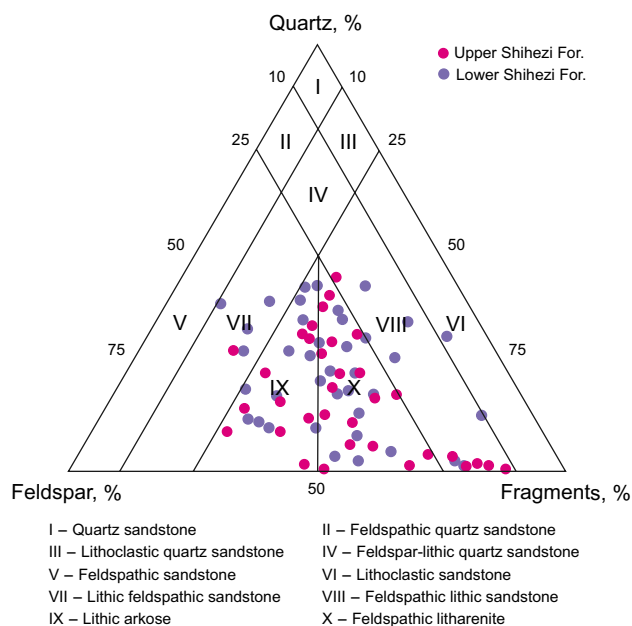


Fig. 2 Rock type of the tight sandstone of P₂h reservoir

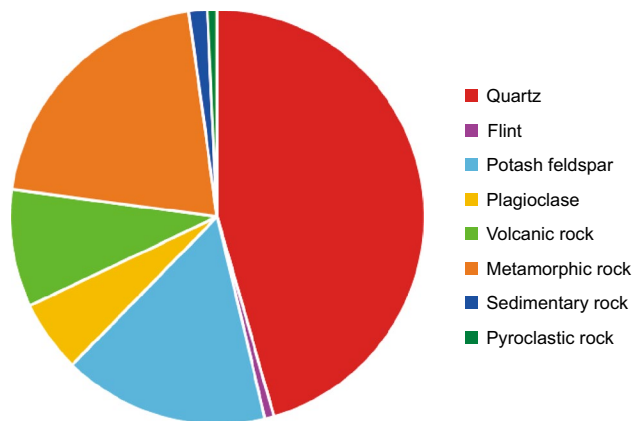


Fig. 3 Rock compositions of the P₂h tight sandstone reservoir

The content of interstitial fillings (including clay minerals and cements) increased as grain size became smaller, except for the medium sandstone (Fig. 5). The reason that the fine sandstone contained more interstitial fillings is that it has more cement and lower clay mineral content. Figure 6 illustrates that most samples contained interstitial fillings less than 15%. Nevertheless, some samples of pebbly fine sandstone and fine sandstone had a larger content of interstitial fillings with some samples exceeding 30%. Overall, 5–10% and 10–15% were the primary distribution intervals of all samples. However, it varied with grain size. The main distribution interval of pebbly coarse sandstone was 5–10%, whereas both coarse sandstone and fine sandstone contained interstitial fillings with a peak interval of 10–15%.

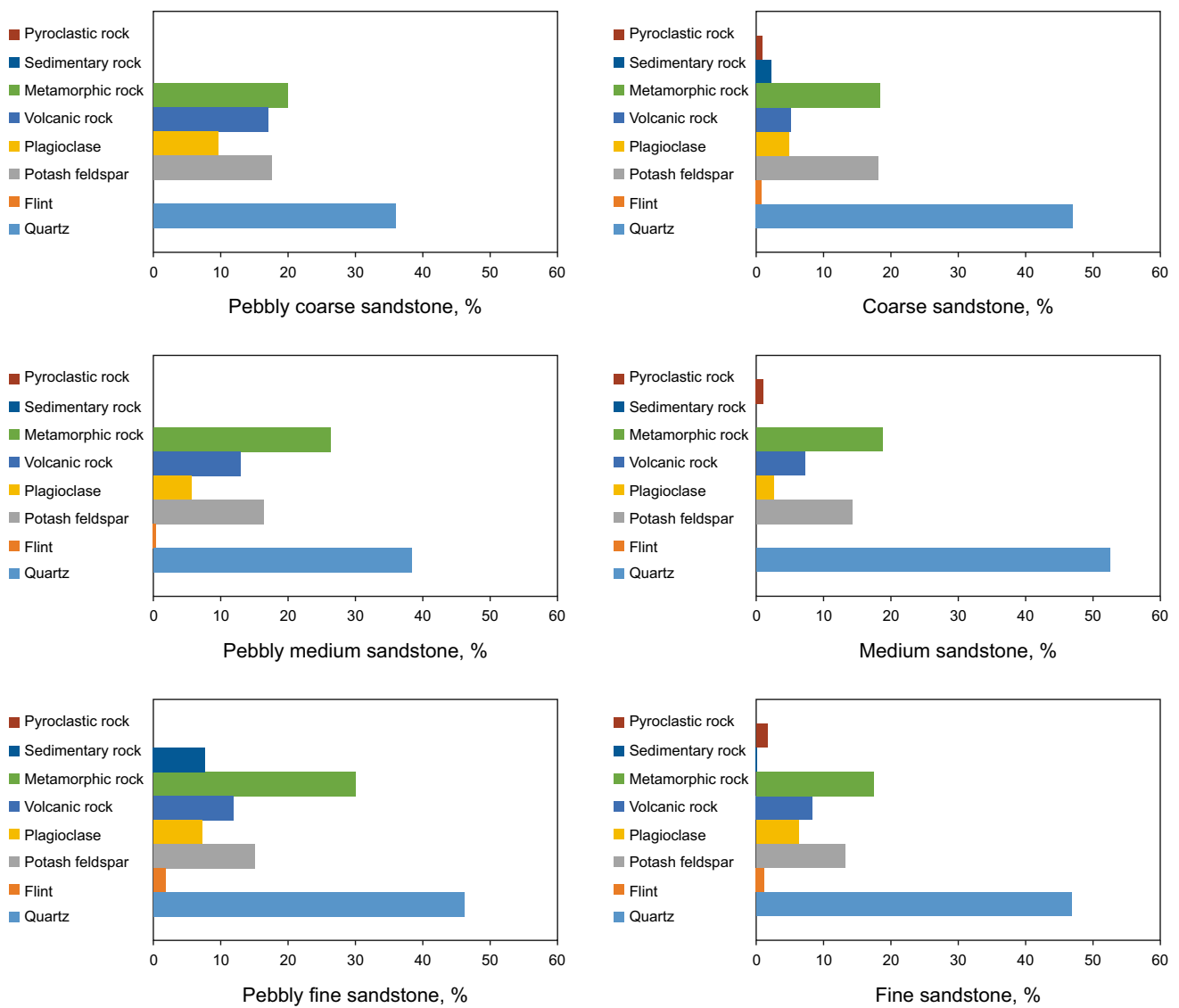


Fig. 4 Rock compositions for 24 samples with diverse lithologies obtained from thin sections

The pebbly medium sandstone, medium sandstone and pebbly fine sandstone had relatively less content of interstitial fillings, with a primary distribution of 5–10%.

As illustrated in Fig. 6b, smectite was not developed within the coarse-grained reservoir in this area. The clay mineral composition varied greatly among these samples of different lithologies. The pebbly coarse sandstone had the lowest contents of illite and kaolinite and the highest illite/smectite-mixed layer (*I/S*) content (exceeding 55%), whereas the medium sandstone contained the highest percentage of kaolinite and lowest percentage of *I/S*. The pebbly fine sandstone had the highest chlorite/smectite-mixed layer (*C/S*) content of about 20%. The chlorite content was more stable among all lithologies, with an average content of approximately 30%.

The casting thin-section and scanning electron microscope analysis provided additional visual details of the characteristics of pore types and structures, as shown in Fig. 7. The P_2h tight sandstone primarily contained four pore types: residual intergranular pores, intergranular dissolution pores, intragranular dissolution pores and intercrystalline pores. Plenty of intragranular dissolution pores and intergranular dissolution pores were developed in the pebbly coarse sandstone (Fig. 7a–c), although these were frequently mixed with residual intergranular pores (Fig. 7c). The intergranular dissolution pores, ranging in size from 10 to 200 μm , were commonly developed within the feldspar and rock fragment. The combination of intragranular pores and fractures within feldspar could largely improve the physical properties (Fig. 7b). Nevertheless, there were more residual intergranular pores within the coarse sandstone

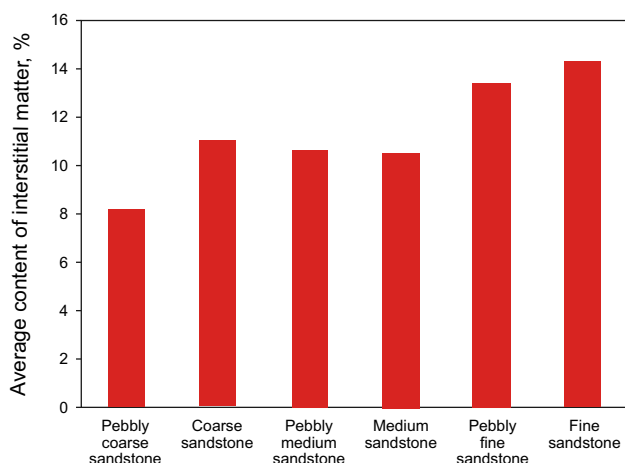


Fig. 5 Average interstitial fillings contents of 72 samples with diverse lithologies obtained from thin-section analysis

in addition to intergranular dissolution pores and intragranular dissolution pores (Fig. 7d–f). The residual intergranular pores might contribute to the large porosity and permeability of the four samples of coarse sandstone. The intergranular dissolution pores were as large as 400 μm in long dimension and maybe developed with intragranular dissolution pores (Fig. 7d). The size of the dissolution pores was identical to that of the pebbly coarse sandstone, with a range of 20–200 μm . The medium sandstone of P_2h reservoir had a large amount of intergranular and intragranular dissolution pores and some residual intergranular pores (Fig. 7g–i). Nevertheless, the size of the intergranular dissolution pores and the intragranular dissolution pores (range of 10–100 μm) was smaller than that of the pebbly coarse sandstone and coarse sandstone. However, the pore number of the medium sandstone was larger than that of the latter two sandstones. Unlike the case of the pebbly coarse sandstone and coarse sandstone, intercrystalline pores, particularly pores between kaolinite (Fig. 7i), contributed a rather large portion of the total porosity in the medium sandstone.

There were relatively few intergranular dissolution pores, intragranular dissolution pores and residual intergranular pores within the samples of fine sandstone (Fig. 7j, k), because the fine sandstone was tightly cemented (Fig. 7k). This phenomenon may account for the low porosity and permeability within the fine sandstone reservoirs. However, a large amount of intercrystalline pores was developed within the fine sandstone samples, as shown in Fig. 7i.

4.2 Lithological control on the physical properties of the tight sandstone

The porosity varied with grain size, as shown in Fig. 7c. Although the pebbly coarse sandstone contained the highest amount of clay minerals, it had the largest porosity, with an

average value of 13%. This likely resulted from the lower cement content. The medium sandstone had the second largest porosity, exceeding 10% on average, whereas the fine sandstone had the lowest porosity. Figure 7d also shows that the pebbly coarse sandstone and the medium sandstone had the largest proportions of large pore, exceeding 10% and 20%, respectively. The peak porosity intervals of the pebbly coarse sandstone, coarse sandstone, pebbly medium sandstone and pebbly fine sandstone were 10–14%, whereas that of the fine sandstone was 6–10%. The medium sandstone had no obvious peak interval and the content of each portion exceeded 20%, showing strong heterogeneity of porosity distribution. Analogously, the pebbly coarse sandstone and medium sandstone also had the highest average permeability values (Fig. 7e). The fine sandstone had the lowest average permeability. The medium sandstone had the second largest average permeability value of approximately 0.06 mD. All samples, except for fine sandstone, had a large proportion of high-permeability interval (Fig. 7f), especially in the case of the pebbly coarse sandstone (> 50%). The permeability of the fine sandstone was lowest, primarily between 0.01–0.1 mD and 0.1–0.5 mD. Similar to the distribution of porosity, the permeability of the medium sandstone was also relatively uniform, with a high percentage of 0.01–0.1 mD interval (> 30%). In addition, a good relationship between porosity and permeability occurred for all lithologies, as shown in Fig. 8, and there was no obvious diversity of the relationship law among the different lithologies.

The impact of clay minerals on the porosity and permeability was also a concern of previous studies (Zhou et al. 2016). The relationships between clay minerals and porosity and permeability in the study area were nonlinear and complex. As shown in Fig. 9, illite content first increased and then decreased with increase in porosity. Nevertheless, illite content was negatively correlated with permeability. Chlorite content was a little positively correlated with porosity and permeability. *I/S* content was weakly positively associated with porosity.

4.3 Pore size distributions of the different lithologies

As shown by the mercury intrusion/extrusion curves for the samples (Fig. 10a–f), the pore–throat structure of each lithology showed diversity to some extent. Nevertheless, the pebbly coarse sandstone and coarse sandstone had obviously lower displacement pressure and larger maximum connected pore–throat radius (MCPR). The coarse sandstone had a lower average value of displacement pressure (0.36 MPa) than the pebbly coarse sandstone (0.96 MPa) because all four samples of the coarse sandstone had good physical properties. The worst sample of these had porosity and permeability values of 11.6% and

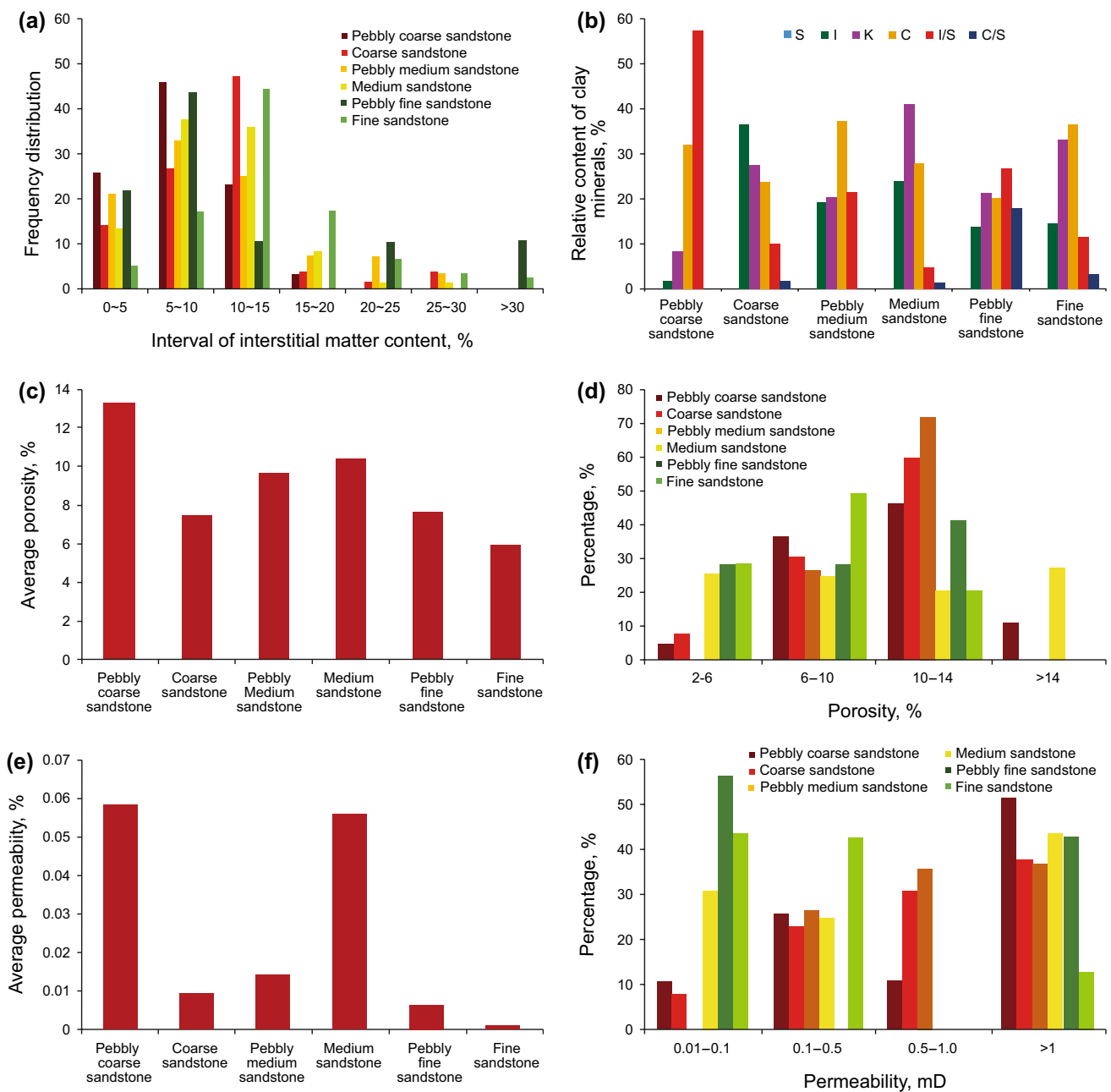


Fig. 6 **a** Frequency distributions of interstitial fillings for 72 samples with diverse lithologies; **b** clay mineral compositions for 24 samples with diverse lithologies; **c** average porosity for 72 samples with diverse lithologies; **d** distribution of porosity for different lithologies; **e** average permeability for 72 samples with diverse lithologies; **f** distribution of permeability for different lithologies

0.91 mD, respectively. The pebbly medium sandstone and medium sandstone had the largest diversity of mercury saturation, with a range of 19–82%, showing the greatest heterogeneity. The pebbly fine sandstone and fine sandstone had relatively low porosity and permeability, as well as low MCPR. Their mercury saturations varied the least, demonstrating the weakest heterogeneity among all lithologies, although these lithologies had a variation range of 20% and 12%, respectively.

The high porosity within the coarser sandstones was generally contributed by large pores (Fig. 11a, b), whereas the porosity in samples with small grain size was more likely due to a large amount of small pores (Fig. 11c–f), although some exceptions existed (W30-1 in Fig. 11e). As shown in Fig. 11, according to the Windland R35 method (Koldziej 1980), coarse sandstone (W19-2) and medium sandstone (W10-3) had larger R35 values in a range of 0.5–2 μm and belonged to “Mesopore” type, whereas other sandstones

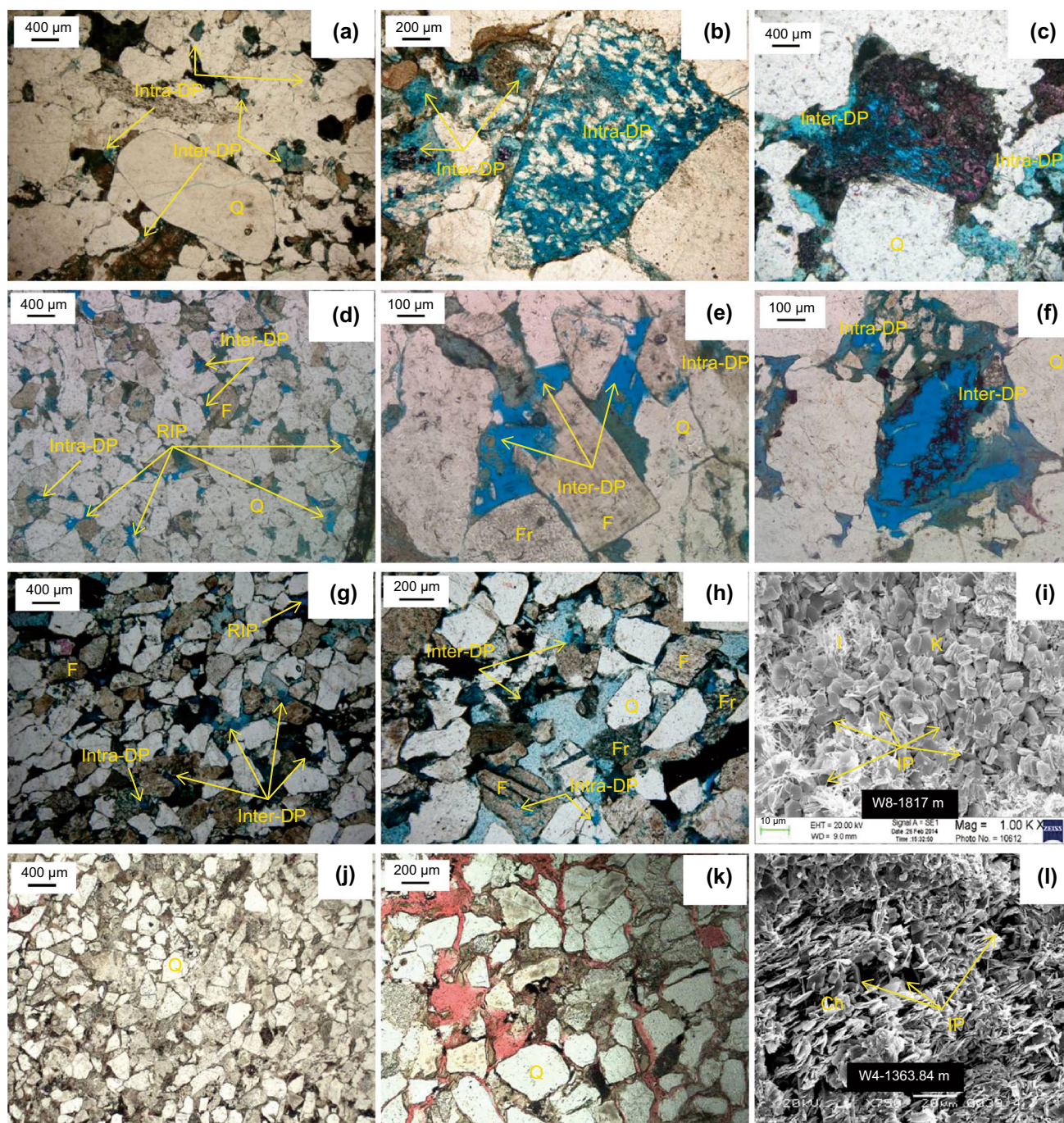


Fig. 7 Photomicrographs showing the pore types within the P_2h tight sandstone: **a, b** intergranular dissolution pores (Inter-DP) and intragranular dissolution pores (Intra-DP) within the pebbly coarse sandstone, PPL, W8, 1797 m; **c** Inter-DP and Intra-DP within the pebbly coarse sandstone, PPL, W19, 1641.12 m; **d** residual intergranular pores (RIP), Inter-DP and Intra-DP within the coarse sandstone, PPL, W21, 1759.75 m; **e** Inter-DP and Intra-DP within the coarse sandstone, PPL, W21, 1759.75 m; **f** Inter-DP and Intra-DP within the coarse sandstone, PPL, W21, 1628.4 m; **g** RIP, Inter-DP and Intra-DP within the medium sandstone, PPL, W10, 1406.5 m; **h** Inter-DP and Intra-DP within the medium sandstone, PPL, W10, 1406.5 m; **i** intercrystalline pores within Kaolinite in the medium sandstone, SEM, W8, 1817 m; **j** strong compaction in the fine sandstone with fewer pores, PPL, W4, 1305.5 m; **k** calcite cementation in the fine sandstone, PPL, W4, 1305.5 m; **l** intercrystalline pores within chlorite in the fine sandstone, SEM, W4, 1363.84 m. PPL plane polarized light, Q quartz, F feldspar, Fr rock fragment, I illite, K kaolinite, Ch chlorite, RIP residual intergranular pore, Inter-DP intergranular dissolution pore, Intra-DP intragranular pore, IP intercrystalline pore

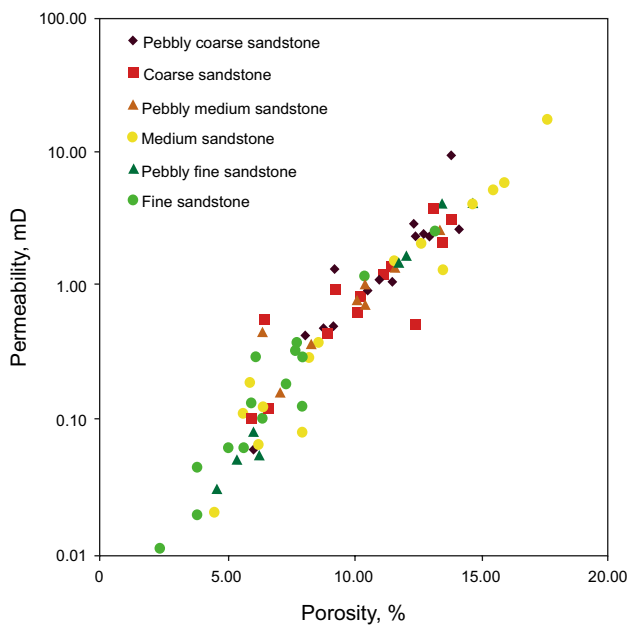


Fig. 8 Relationship between porosity and permeability for samples with diverse lithologies

had R35 values in a range of 0.1–0.5 μm belonged to “Microport” type.

The T_2 spectra for samples with diverse lithologies in the study area were not identical. Additionally, there was a difference between the T_2 distributions of the same lithology. Overall, the T_2 distributions of pebbly coarse sandstone, coarse sandstone, pebbly medium sandstone and medium sandstone showed a wide distribution and large signal amplitude, which were around 0.1–1000 ms and 0.6–1.1%, respectively (Fig. 12a–c). The maximum values of the T_2 distributions and signal amplitudes of pebbly fine sandstone and fine sandstone were approximately 400 ms and 0.63%, respectively, indicating a lower porosity within this type of reservoir (Fig. 12e, f). Therefore, similar to the HPMI results, the reservoirs with large grain sizes tended to have high porosity and vice versa. In addition, the medium sandstone had the largest difference of T_2 distributions, with a signal amplitude range of 0.15–1.2% (Fig. 12d). The NMR results also reveal the pore size distribution and connectivity of pores. Most of the samples had two peaks in the T_2 distribution. The first peak with a smaller T_2 relaxation time represented small pores, and the second peak corresponding to a larger T_2 relaxation time indicated large pores. Although the coarser sandstones had high porosity on the whole, the T_2 relaxation times of the first peak of all of the samples were all located within a relatively narrow range of 0.8–1.5 ms. The samples with large porosity had a T_2 relaxation time of the first peak of 1.2–1.5 ms, and those with relatively small porosity had a smaller T_2 relaxation time corresponding to the first peak of approximately 0.8–1.2 ms. Unlike the first peak of the T_2

relaxation time, the second peak had a rather wide range of T_2 time of 10–110 ms. The pebbly coarse sandstone, coarse sandstone and pebbly medium sandstone had a large T_2 time value of the second peak, greater than 100 ms. Conversely, the other sandstones with smaller grain size had a T_2 time value of the second peak of 10–100 ms, with an average of 50 ms. Additionally, strong heterogeneity existed within each lithological group. In the pebbly medium sandstone group, the four samples had entirely different T_2 distributions. The W18 sample had a narrow T_2 time distribution of 0.1–100 ms, a small signal amplitude of 0.26% and a rather small second peak corresponding to a signal amplitude of 0.6%. However, the W12 and W33-3 samples had a wide T_2 time distribution of 0.1–1100 ms and large signal amplitude of 0.4–0.6%. The latter even had a second peak with larger signal amplitude than the first peak. Unlike these three types, the T_2 distribution of W36 was unimodal with a large-signal amplitude of 0.5%.

The connectivity between the two peaks was diverse within each group. The samples with large porosity and permeability, such as the W21-1 in the pebbly coarse sandstone group, the W19-2 and W21-2 in the coarse sandstone group, the W12 in the pebbly medium sandstone, the W30-1 in the pebbly fine sandstone group and the W30-2 in the fine sandstone group generally had better connectivity of the two peaks. However, the good connectivity between small pores and larger pores did not indicate large permeability, such as for W35 in the medium sandstone group. Previous studies showed that permeability was mainly related to large pores and that the MCPR from HPMI may indicate permeability more accurately³⁵. Therefore, the connectivity between two peaks from NMR is not an ideal indicator of rock permeability.

4.4 Comparison of pore size distribution between coarse sandstones and fine sandstones

Zhu et al. (2018) analyzed the distribution of pore size of fine-grained samples in the southwestern Ordos Basin, including the fine sandstone, very fine sandstone and coarse siltstone, using NMR and thin sections. These samples contained more quartz and feldspars and less interstitial material. The amount of cements was less than that of the coarse-grained samples in this area. In addition, the porosity of the fine-grained samples was also much lower. The T_2 distributions of the fine samples had a narrower body and lower peak value. The porosity at the radii connecting two peaks was almost lower than 0.1 within the fine-grained samples, whereas the value was larger than 0.2 within most coarse-grained samples. Therefore, by integrating the two aspects, it can be concluded that the fine-grained reservoir tends to have poor pore structure and correspondingly low porosity and permeability, except for the medium sandstone.

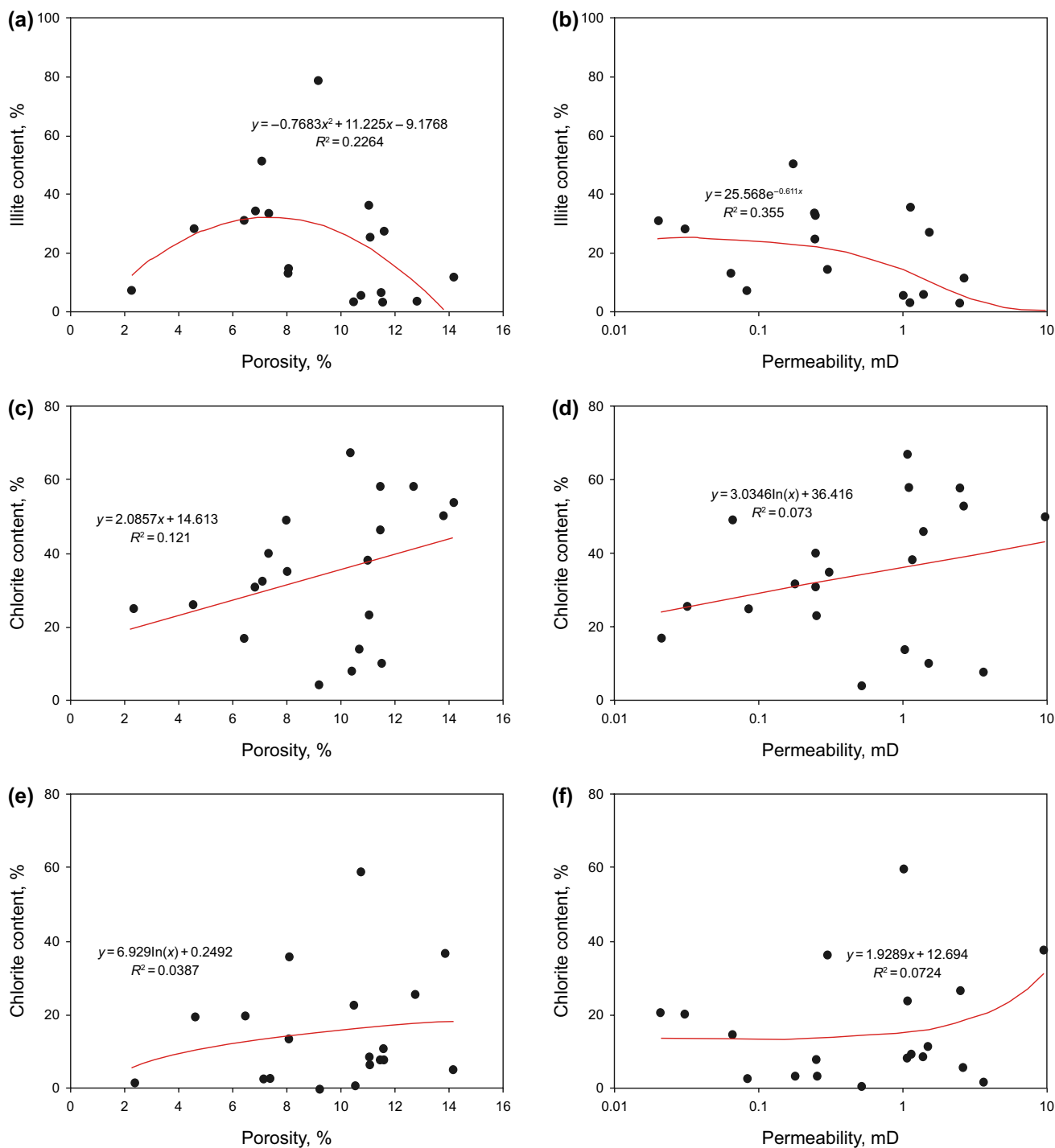


Fig. 9 Plots showing the relationship between **a** illite content and porosity, **b** illite content and permeability, **c** kaolinite content and porosity, **d** kaolinite content and permeability, **e** chlorite content and porosity, **f** chlorite content and permeability, **g** andreattite content and porosity, **h** andreattite content and permeability

5 Conclusion

A series of experiments were conducted to investigate the lithological control on the pore structures and physical properties of the P₂h tight sandstone reservoir in the eastern

Ordos Basin. This reservoir has strong heterogeneity in the aspects of both macroscopic porosity and permeability and microscopic pore–throat structures within all types of rock. Although many of the sandstone samples were fairly tight, some types of reservoirs with relatively high

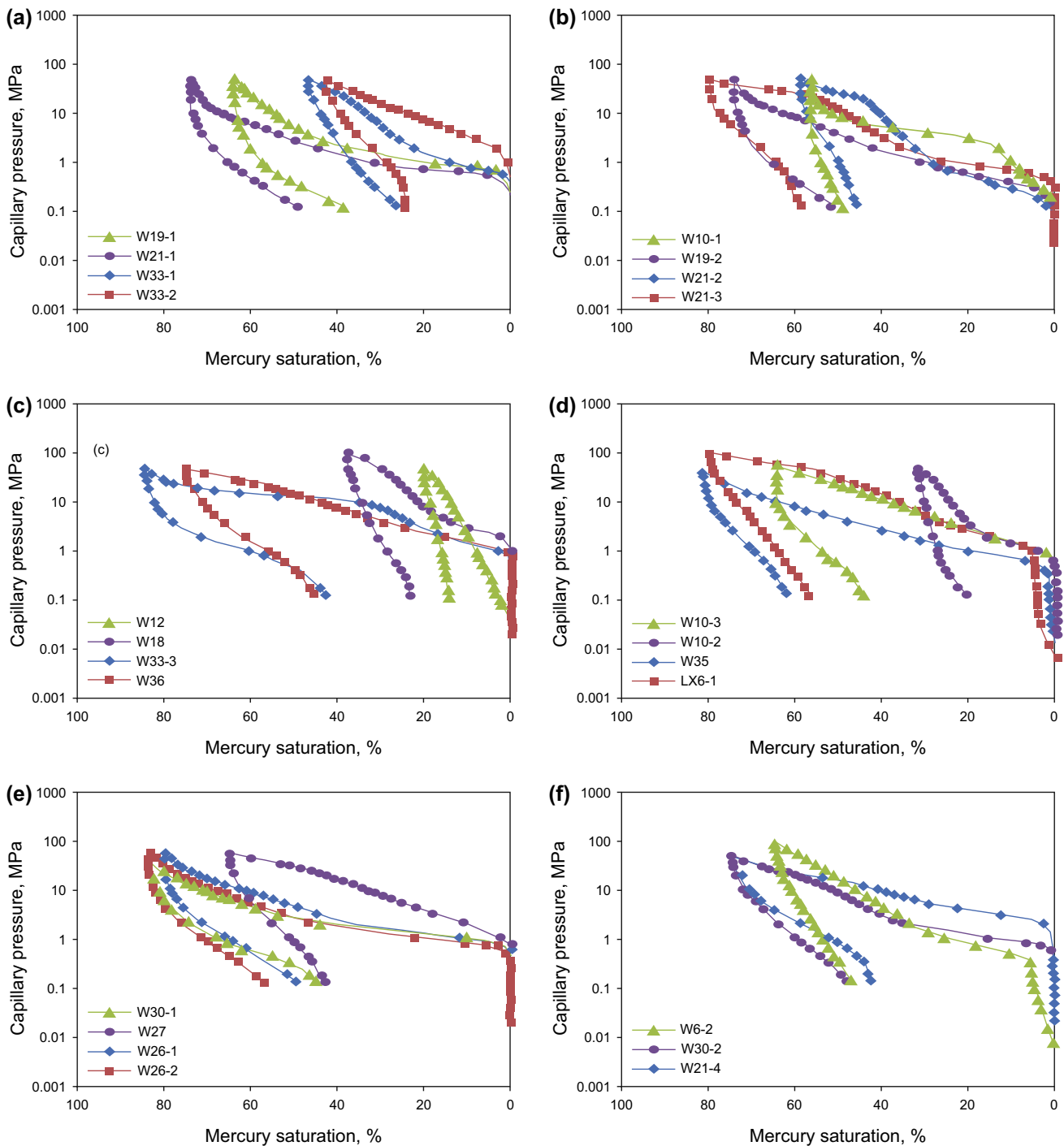


Fig. 10 Mercury injection curves for the samples of **a** pebbly coarse sandstone, **b** coarse sandstone, **c** pebbly medium sandstone, **d** medium sandstone, **e** pebbly fine sandstone, **f** fine sandstone

average porosity and permeability can be viewed as spot sweets, such as pebbly coarse sandstone, coarse sandstone and medium sandstone reservoirs. Nevertheless, the medium sandstone reservoir has the strongest heterogeneity, which may bring some risks of tight sandstone gas exploration and development.

The P₂h tight sandstone contains mainly intergranular dissolution pores, intragranular dissolution pores and inter-crystalline pores. The former is as large as 200 μm. Some coarser-grained samples had a small amount of residual

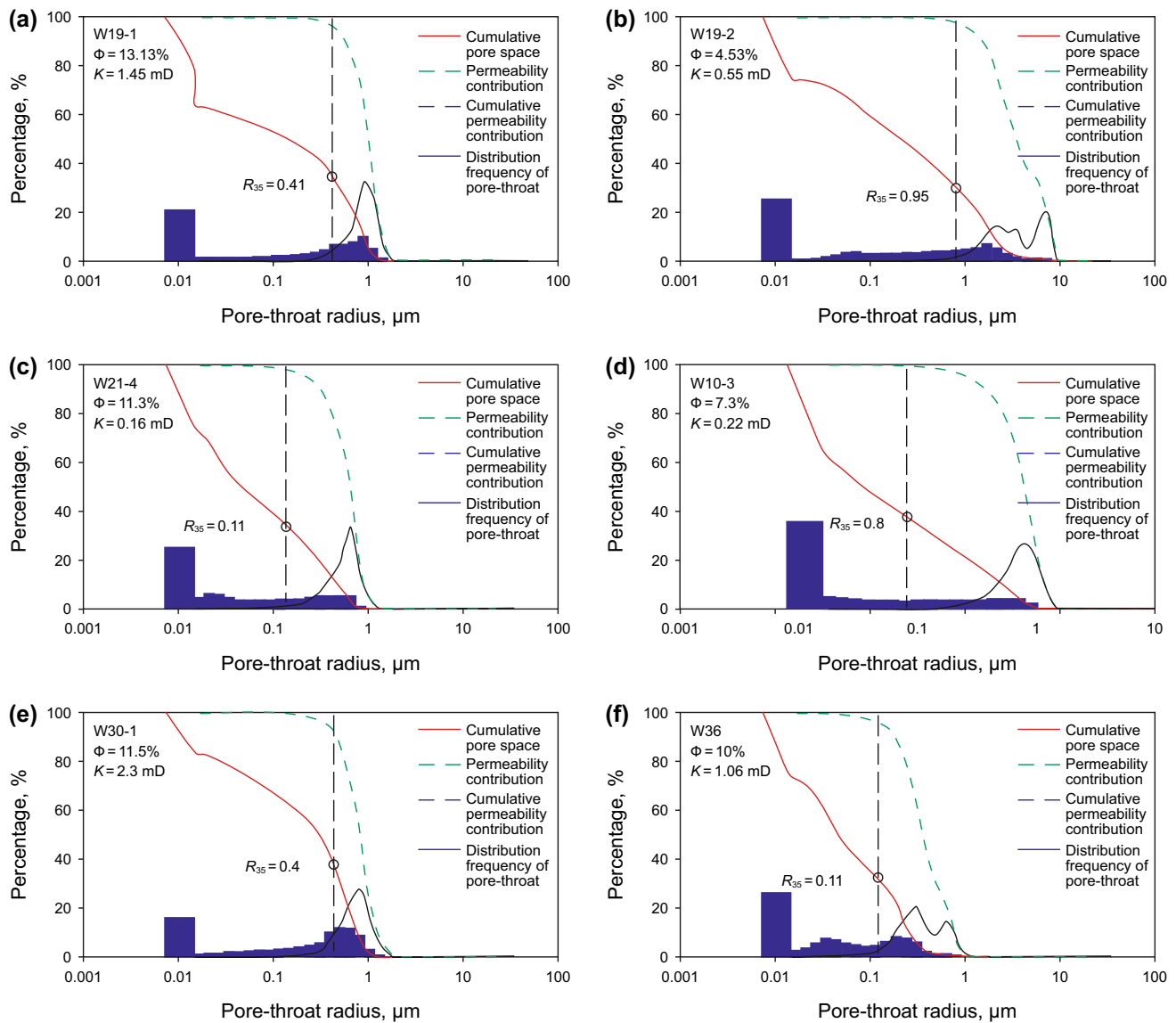


Fig. 11 Pore size distribution (PSD) and its contributions to porosity and permeability for samples of **a** pebbly coarse sandstone, **b** coarse sandstone, **c** pebbly medium sandstone, **d** medium sandstone, **e** pebbly fine sandstone and **f** fine sandstone

intergranular pores. The coarse sandstone and medium sandstone contain relatively large intragranular dissolution pores, whereas the fine sandstone has more intercrystalline pores with a more homogeneous pore structure.

As the T_2 spectra from the NMR and HPMI showed, the coarse-grained sandstones tend to have large pores and

throat radii, whereas the fine-grained sandstones generally have more small pores and throat radii. Large pores contribute most to the permeability within all samples. Nevertheless, small pores in the fine-grained sandstones contribute more to the porosity than small pores in the coarse-grained sandstone.

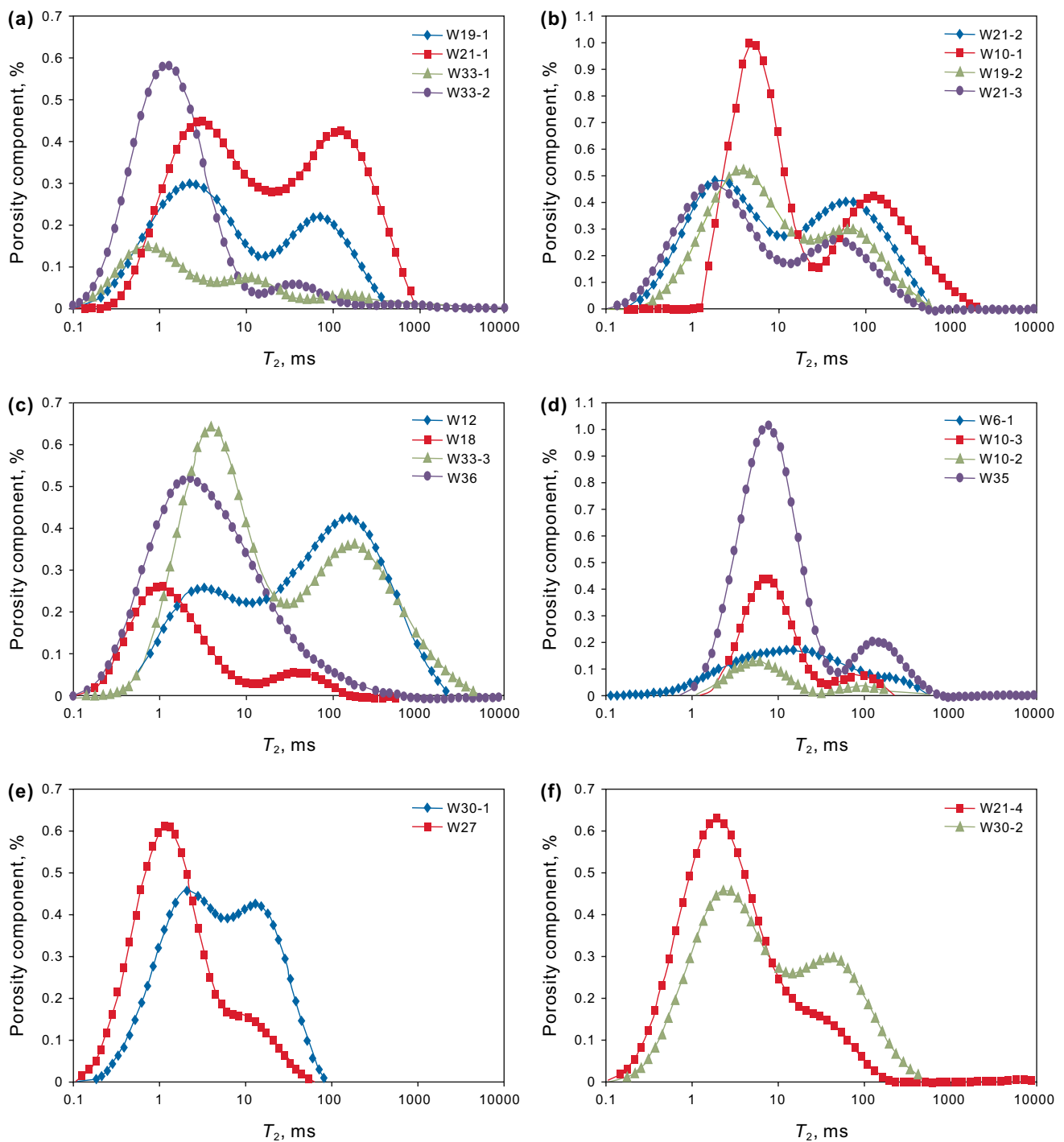


Fig. 12 T_2 distributions for samples of **a** pebbly coarse sandstone, **b** coarse sandstone, **c** pebbly medium sandstone, **d** medium sandstone, **e** pebbly fine sandstone and **f** fine sandstone

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