

Oil maturity assessment using maturity indicators based on methylated dibenzothiophenes

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Abstract: Aromatic fractions of 140 oils and condensates that originated from different types of source rocks (marine shale, terrestrial shale and marine carbonate) were analyzed using gas chromatography–mass spectrometry (GC–MS) to investigate the relative distributions of methylated dibenzothiophenes with respect to thermal maturity. The positions of methyl groups of trimethyldibenzothiophene isomers (TMDBTs) including those used in the definition of maturity indicator TMDBT index in previous studies were firmly identified by co-elution of internal standards in GC–MS analysis and by comparing with reported retention indices. A new maturity ratio related to dimethyldibenzothiophenes (DMDBTs) is proposed on the basis of the differences in thermodynamic stability among different DMDBT isomers. Another maturity index (TMDBT-I2) based on TMDBTs is also suggested on the basis of our empirical observations and presumed thermodynamic stability of TMDBT isomers. These two newly proposed (2,6 + 3,6)-/1,4-DMDBT ratio and TMDBT-I2 correlate well with MDR (4-/1-methyldibenzothiophene) and 2,4-/1,4-DMDBT ratios, suggesting their common chemical reaction mechanisms and similar behavior with increasing maturity. Therefore, they can be effectively applied for maturity assessments. Furthermore, the TMDBTs related maturity parameters are more reliable for over-mature oils and condensates due to the relatively higher concentrations of thermodynamically unstable TMDBT isomers, i.e. 1,4,6-, 1,4,8- and 3,4,6-TMDBT in this study than those of 1-methyldibenzothiophene (1-MDBT) or 1,4-DMDBT. In contrast with 4,6-/1,4-DMDBT, the newly proposed (2,6 + 3,6)-/1,4-DMDBT ratios for oils that originated from different types of source rocks have approximately same relationship with the oil maturity (R_c %). This suggests that the lithology and organic facies may have relatively less influence on (2,6 + 3,6)-/1,4-DMDBT ratio compared to 4,6-/1,4-DMDBT. The maturity parameters based on methylated dibenzothiophenes are particularly useful in the maturity assessments of post- and over-mature oils and condensates and can complement maturity indicators based on steranes and terpanes.

Key words: Dibenzothiophene, maturity indicator, molecular thermodynamic stability, sedimentary organic matter

1 Introduction

Polycyclic aromatic sulfur heterocyclics (PASHs), such as benzothiophenes (BTs) and dibenzothiophenes (DBTs) are important organosulfur compounds in oils and ancient organic matter. In spite of the debate on the specific source precursors for PASHs and poor knowledge of the chemical mechanisms and kinetics of their formation during diagenesis, distributions of these compounds have been widely used as organic facies and maturity indicators of petroleum and source rocks (e.g. Ho et al, 1974; Radke et al, 1986; Radke, 1988; Budzinski et al, 1991; Hughes et al, 1995; Chakhmakhchev et al, 1997;

Santamaria-Orozco et al, 1998; Shen et al, 1998; Asif et al, 2009). Recently, the absolute concentrations of total DBTs and the relative abundances of methylated DBT isomers in petroleum are also suggested to be potential geochemical indicators for tracing petroleum migration pathways (Wang et al, 2004; Li et al, 2008a).

Compounds with thiophenic ring structures in kerogen are probably the precursors of BT, DBT and their homologues in oils (Orr, 1986). Proteins are among the sulfur-bearing constituents of organisms from which the kerogen was formed. Due to the insufficient amount of sulfur in plants and animals to produce the large amounts of sulfur compounds in petroleum and sedimentary organic matter, extra sulfur must have been incorporated into the organic matter from source rocks, especially from high sulfur source rocks (Gransch and

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Posthuma, 1974 and references therein). A thermal reaction between elemental sulfur and the organic matter of sediments was proposed to be responsible for the genesis of the sulfur-containing organic compounds (Hanson, 1960; Douglas and Mair, 1965). Recently, laboratory simulation experiments showed that biphenyls and sulfur can form dibenzothiophenes with carbon catalysis (Xia and Zhang, 2002; Asif et al, 2009). This suggests that the widespread distributions of dibenzothiophene and alkylated dibenzothiophenes in sediments and petroleum may be the result of a catalytic reaction of biphenyl ring systems and surface-adsorbed sulfur on the surface of carbonaceous material (Asif et al, 2009; Li et al, 2013a; 2013b). Abundant BTs and DBTs in petroleum and sedimentary organic matter are commonly associated with anoxic and sulfur-rich depositional environments (Hughes, 1984; Hughes et al, 1995; Radke et al, 2000; Li et al, 2011). The cross plots of dibenzothiophenes/phenanthrene (DBT/Phen) vs. pristane/phytane (Pr/Ph) ratios (Hughes et al, 1995), alkylated dibenzothiophene/alkylated dibenzofuran (ADBT/ADBF) ratios (Radke et al, 2000) and the ternary diagram of the relative abundance of fluorenes, dibenzothiophenes and dibenzofurans (Li et al, 2013c) have been introduced as indicators of the depositional environment and lithology of petroleum source rocks.

The molecular thermodynamic stability of methylated dibenzothiophenes varies with the positions of the methyl groups (Radke, 1988; Budzinski et al, 1991; Richard, 2001). Thus, the distribution and relative abundances of methylated DBT isomers can be used as maturity indicators for both crude oils and source rocks (Radke et al, 1986; Radke, 1988; Schou and Myhr, 1988; Budzinski et al, 1991; Radke et al, 1991; Radke and Willsch, 1994; Bao et al, 1996; Chakhmakhchev et al, 1997; Santamaría-Orozco et al, 1998; Shen et al, 1998; Krüge, 2000). Some maturity indicators on the basis of methylated dibenzothiophene (MDBT), dimethyldibenzothiophene (DMDBT) and ethyldibenzothiophene (EDBT) such as MDR (4-/1-MDBT ratio) (Radke et al, 1986; Radke, 1988), DBTI ((2 + 3)-MDBT/2×(1-MDBT)) (Payzant et al, 1989), MDR' (4-/((4 + 1)-MDBT) and EDR' (4,6-DMDBT/(4-EDBT + 4,6-DMDBT)) (Radke and Willsch, 1994), 4,6-/1,4-DMDBT and 2,4-/1,4-DMDBT ratios (Chakhmakhchev et al, 1997; Santamaría-Orozco et al, 1998) were proposed to assess the thermal maturity of petroleum and organic matter. These maturity parameters change in a regular fashion with increasing maturity up to high levels. Therefore, they are effective maturity indicators, and particularly useful for high to over-mature oils and condensates lacking biomarkers (Chakhmakhchev et al, 1997). Some discrepancies in calibration of the MDR maturity parameter with vitrinite reflectance (R_o) have also been reported (Dzou et al, 1995). Chakhmakhchev et al (1997) attributed these discrepancies to differences in analytical procedures and/or organic facies control in different geological settings. Huang and Pearson (1999) showed that high MDRs are associated with fresh water source deposition whilst low MDRs are found in oils of hypersaline origin in the Bohai Bay Basin, East China.

The empirical TMDBT index based on the relative abundances of some trimethyldibenzothiophene (TMDBT)

isomers (peak 3 and peak 5) in m/z 226 mass chromatograms (Fig. 1) was introduced by Chakhmakhchev et al (1997). The TMDBT index correlates well with the MDR and other DBT-based maturity indicators, suggesting its similar behavior with increasing maturity. However, the specific positions of these three methyl groups for those TMDBT isomers were not assigned due to the lack of internal standards and reference retention indices at that time.

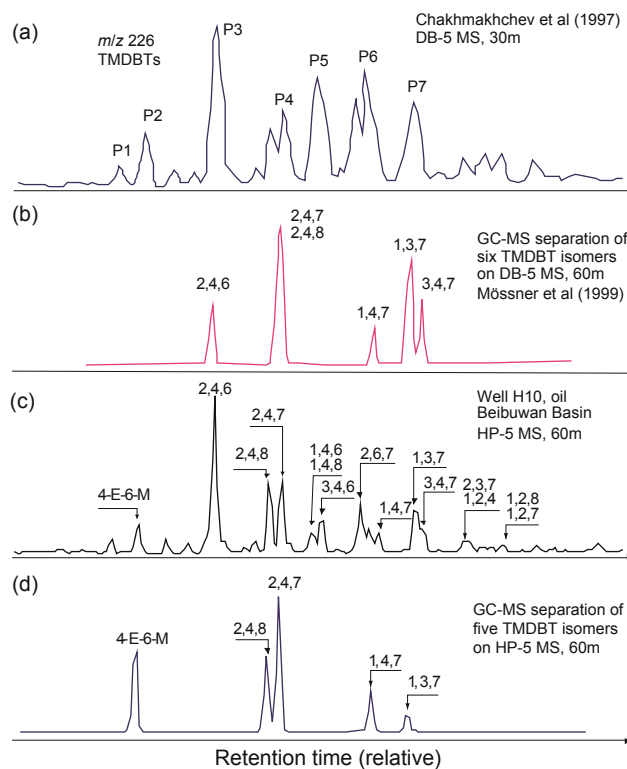


Fig. 1 Identification of trimethyldibenzothiophenes in crude oils. Numbers on peaks indicate the methyl substitution patterns on dibenzothiophene

In this study, we firmly identify a series of TMDBT isomers including peak 3 and peak 5 compounds assigned by Chakhmakhchev et al (1997) by co-elution of internal standards and by comparison of the calculated retention indices (Fig. 1) with those reported in references (Lee et al, 1979; Mössner et al, 1999; Schade and Andersson, 2006). Another TMDBT based maturity parameter is proposed on the basis of our empirical observations and in the light of the differences in thermodynamic stability between different TMDBT isomers. Furthermore, a new DMDBT maturity indicator is also suggested. A large data set, comprising 50 oils and condensates originated from typical terrestrial source rocks, 45 oils from marine shales, and 45 oils from marine carbonates was utilized in this study. The results obtained may provide additional geochemical parameters for maturity evaluation, oil-to-oil and oil-to-source correlations, especially for the characterization of light oils and condensates with low concentrations of biomarkers. The newly proposed DMDBT maturity parameters have some advantages, as they may be more organic facies independent and more reliable for over-mature petroleum.

2 Samples and methods

2.1 Samples

A total of 50 oil samples collected from the Fushan Depression, Beibuwan Basin (South China Sea) were analyzed for the investigation of DBTs and biomarkers distributions. The Fushan Depression is one of the many Mesozoic–Cenozoic rifting half-grabens in the northern continental shelf of the South China Sea. The discovered oil pools mainly occur in the third ($E_2/3$) and first ($E_2/1$) members

of the Eocene Liushagang Formation (Li et al, 2008b). Oils are dominated by light oils or condensates with an average gravity of 47° (API) (Table 1). They are typical terrestrial-sourced high-mature oils with a highest $R_c\%$ equivalent of 1.0-1.1 (Li et al, 2007a; 2007b). The biomarkers are present in very low concentrations or below the detection limit in GC–MS analyses (Li et al, 2010). The light oils and condensates were derived from type II to III kerogen in source rocks of the second member of the Eocene Liushagang Formation (Ding et al, 2003; Li et al, 2008c).

Table 1 The methylated dibenzothiophenes related maturity parameters in oils from the Beibuwan Basin, South China Sea

Sample No.	Well	Depth, m	API, °	4-/1-	4,6-/1,4-	2,4-/1,4-	(2,6+3,6)/1,4-	TMDBT-II	TMDBT-I2	MPR	MPI-1	R_c %	Ts/Tm	22S/(22S+22R)	Pr/Ph	ADBT/ADBF	Oleanane/ $C_{30}H$
FS1	H1	3151.2-3153.4	45	8.31	2.07	1.77	4.23	2.36	2.11	1.48	0.90	0.94	2.04	0.60	3.29	0.76	0.41
FS2	H1-1	3347.2-3401.1	46	8.03	2.07	1.51	4.07	2.59	1.99	1.43	0.86	0.92	8.70	0.65	3.69	0.81	0.56
FS3	H2	2970.6-3014.0	49	7.41	1.83	1.51	3.81	2.70	2.09	1.31	0.83	0.90	3.05	0.61	3.97	0.48	0.32
FS4	H2-1	2909.8-2954.6	57	6.10	1.89	1.36	3.42	2.54	1.90	1.35	0.82	0.89	5.85	0.62	3.68	0.56	0.38
FS5	H2-2	2973.0-3021.0	46	5.67	1.76	1.38	3.24	2.64	2.06	1.26	0.76	0.86	3.31	0.60	3.71	0.51	0.31
FS6	H2-3	2994.6-3084.0	55	9.26	3.07	2.33	5.38	2.84	1.76	1.37	0.82	0.89	3.38	0.65	3.97	0.47	0.38
FS7	H3	3015.0-3182.6	51	6.69	1.59	1.32	3.48	1.92	1.86	1.18	0.76	0.86	3.04	0.52	3.49	0.69	0.37
FS8	H3-1	3949.6-3286.2	43	6.93	1.83	1.55	3.61	2.42	1.85	1.31	0.78	0.87	6.40	0.62	3.76	0.62	0.44
FS9	H3-2	3087.0-3144.6	57	6.08	1.47	1.20	3.15	2.05	1.59	1.26	0.76	0.86	4.23	0.58	3.83	0.54	0.30
FS10	H3-3	3158.6-3221.0	47	5.85	1.59	1.43	3.27	2.06	1.73	1.23	0.77	0.86	4.37	0.60	3.74	0.54	0.37
FS11	HX4	3253.0-3259.6	41	5.14	1.23	1.02	2.40	1.60	1.56	0.98	0.66	0.79	2.61	0.61	4.00	0.77	0.62
FS12	HX4	2965.5-3073.4	43	6.64	1.73	1.43	3.32	2.55	2.00	1.26	0.74	0.84	3.83	0.59	4.00	0.49	0.28
FS13	H4-1	2851.0-2862.2	53	7.55	1.53	1.34	3.16	2.77	2.14	1.17	0.75	0.85	2.77	0.60	3.99	0.44	0.26
FS14	H5	3008.6-3011.8	44	4.75	1.21	1.01	2.53	1.81	1.42	0.98	0.65	0.79	3.67	0.62	3.93	0.45	1.02
FS15	H5	2715.8-2731.0	53	9.55	2.50	1.92	4.71	3.23	2.56	1.55	0.92	0.95	3.25	0.49	3.93	0.74	0.51
FS16	H6	2884.6-2899.0	39	3.03	1.07	1.00	2.30	1.56	1.32	1.01	0.63	0.78	3.17	0.59	4.54	0.38	0.18
FS17	H7	3516.6-3521.0	43	21.56	6.31	4.63	10.37	4.48	3.27	2.45	1.25	1.15	5.67	n.d.	3.98	0.19	0.50
FS18	H7	3383.2-3493.4	55	55.36	19.18	12.24	28.30	10.28	6.04	3.77	1.59	1.35	19.00	n.d.	3.98	0.17	1.30
FS19	HD1	3329.4-3372.8	56	14.57	3.46	2.48	5.83	3.61	2.59	1.95	1.07	1.04	7.33	0.61	3.48	0.68	1.25
FS20	HD1-1	3347.0-3384.0	45	12.90	4.82	3.24	7.71	3.28	2.22	2.00	1.06	1.04	1.78	0.57	3.24	0.86	1.09
FS21	HD2	2542.0-2563.0	40	5.96	2.15	1.64	3.79	2.06	1.68	1.34	0.78	0.87	4.24	0.61	4.22	0.47	0.31
FS22	M1	3105.2-3108.0	38	3.42	1.14	1.06	2.27	1.63	1.39	0.96	0.62	0.77	1.25	0.59	3.45	0.41	0.11
FS23	M1	3084.0-3086.0	37	3.32	1.03	0.98	2.38	1.39	1.33	0.98	0.65	0.79	1.37	0.59	3.45	0.46	0.11
FS24	M2	3011.8-3047.6	39	3.78	1.29	1.25	2.86	1.45	1.41	0.98	0.65	0.79	1.40	0.59	3.68	0.39	0.11
FS25	JFn1	2213.5-2216.0	38	4.49	1.29	1.17	2.68	2.02	1.76	1.26	0.78	0.87	1.84	0.59	4.71	0.53	0.32
FS26	H10	2961.6-2981.8	56	7.99	1.78	1.54	3.10	2.29	2.23	1.51	0.96	0.99	2.57	0.59	4.06	0.46	0.32
FS27	H2-13	3043.4-3085.4	46	4.46	1.39	1.44	3.13	2.02	1.94	1.17	0.80	0.92	2.33	0.59	4.28	0.28	0.26
FS28	H2-16	2892.6-2972.0	45	4.78	1.49	1.49	3.01	1.99	2.03	1.26	0.87	0.95	2.85	0.58	4.24	0.42	0.30
FS29	H2-17X	2998.6-3020.0	52	7.17	2.23	1.87	4.42	2.66	2.51	1.56	1.00	1.00	4.26	0.53	4.00	0.57	0.45
FS30	H2-18	3151.6-3160.0	n.d.	6.74	1.67	1.37	3.67	1.87	1.86	1.21	0.82	0.93	1.56	0.46	4.49	0.60	0.35

(To be continued)

(Continued)

Sample No.	Well	Depth, m	API, °	4-/1-	4,6-/1,4-	2,4-/1,4-	(2,6+3,6)/1,4-	TMDBT-I1	TMDBT-I2	MPR	MPI-1	R _c %	Ts/Tm	22S/(22S+22R)	Pr/Ph	ADBT/ADBF	Oleanane/C ₃₀ H
FS31	H2-5	3114.0-3148.8	59	6.32	1.77	1.52	3.20	2.51	2.25	1.35	0.89	0.96	1.08	0.54	4.27	1.04	0.17
FS32	H2-6	3147.8-3192.0	52	6.89	1.83	1.60	3.52	2.34	2.20	1.31	0.90	0.96	2.45	0.50	4.19	0.56	0.33
FS33	H3-12x	3194.4-3238.8	47	5.38	1.37	1.27	3.38	1.77	1.80	1.14	0.78	0.91	3.55	0.51	3.98	0.55	0.30
FS34	H3-5	3109.1-3128.2	54	5.44	1.53	1.41	3.49	1.74	1.65	1.29	0.84	0.94	1.86	0.59	4.31	0.55	0.20
FS35	H3-6	3114.4-3148.2	52	5.57	1.56	1.47	3.51	1.99	1.95	1.25	0.85	0.94	3.00	0.59	3.98	0.47	0.29
FS36	H4-3	3015.4-3060.6	39	4.50	1.12	1.08	2.77	1.90	1.64	1.15	0.76	0.90	1.94	0.58	4.58	0.31	0.27
FS37	H8-1	3585.2-3669.2	48	11.34	2.81	2.39	4.66	3.25	2.68	1.73	1.15	1.06	6.14	0.32	3.36	1.23	1.07
FS38	H8-1X	3678.0-3733.6	n.d.	13.27	3.48	2.70	6.87	2.92	2.52	2.10	1.26	1.11	n.d.	n.d.	3.19	1.75	n.d.
FS39	HD1-5	3486.2-3510.4	59	9.83	2.63	2.25	5.15	2.72	2.66	1.80	1.09	1.04	1.33	0.51	4.34	1.21	0.24
FS40	HD1-7	3450.7-3462.5	n.d.	9.69	2.55	2.08	5.01	2.51	2.20	1.77	1.07	1.03	1.33	0.50	4.18	1.36	0.25
FS41	Hong3	2504.0-2513.4	n.d.	4.84	1.23	1.23	2.43	1.51	1.57	1.10	0.81	0.92	1.38	0.52	4.64	0.50	0.23
FS42	Hong5	2504.0-2513.4	n.d.	3.35	0.99	1.10	2.62	1.65	1.77	1.06	0.76	0.90	2.33	0.60	4.35	0.28	0.29
FS43	Hong5-1	2725.4-2775.0	n.d.	2.01	0.60	0.76	2.05	0.90	1.02	0.91	0.65	0.86	1.56	0.55	4.38	0.26	0.93
FS44	L11	2842.4-2846.0	42	5.21	1.03	1.36	2.42	1.57	1.67	1.51	0.71	0.88	1.22	0.55	4.95	0.16	0.62
FS45	L12	2794.8-2801.0	32	5.20	0.87	2.02	2.86	2.28	1.99	1.90	0.70	0.88	1.27	0.55	5.12	0.16	0.81
FS46	M4	3208.2-3275.8	36	2.77	1.02	0.96	2.50	1.14	1.17	1.00	0.64	0.85	1.27	0.57	3.70	0.38	0.15
FS47	M4-1	3459.6-3493.4	39	2.94	0.82	0.89	1.89	1.05	1.04	0.86	0.61	0.84	1.38	0.52	3.99	0.39	0.18
FS48	M5	3217.3-3219.0	n.d.	3.47	0.97	1.04	2.82	1.30	1.35	0.98	0.69	0.87	1.50	0.52	4.06	0.54	0.16
FS49	M5-1	3150.0-3160.1	n.d.	3.88	1.01	1.00	2.57	1.31	1.40	0.97	0.67	0.87	1.63	0.56	4.07	0.45	0.16
FS50	M6X	2628.2-2631.0	n.d.	3.68	1.03	1.03	2.32	1.29	1.40	0.94	0.68	0.87	1.38	0.52	4.24	0.47	0.19

Notes: 4-/1-: 4-/1-methylidibenzothiophene; 4,6-/1,4-: 4,6-/1,4-dimethylidibenzothiophene; 2,4-/1,4-: 2,4-/1,4-dimethylidibenzothiophene; (2,6 + 3,6)-/1,4-: (2,6 + 3,6)-/1,4-dimethylidibenzothiophene; TMDBT-I1: 2,4,6-/(1,4,6 + 1,4,8 + 3,4,6)-TMDBT; TMDBT-I2: (2,4,8 + 2,4,7)-/(1,4,6 + 1,4,8 + 3,4,6)-TMDBT; MPR: 2-/1-MP (MP: methylphenanthrene); MPI-1: $1.5 \times (2\text{-MP} + 3\text{-MP}) / (\text{Phen} + 9\text{-MP} + 1\text{-MP})$ (Phen: phenanthrene); R_c (%) = $0.4 + 0.6 \times \text{MPI-1}$; 22S/(22S+22R): C₃₁ hopane 22S/(22S+22R); Pr/Ph: Pristane/Phytane; ADBT/ADBF: alkylated dibenzothiophene/alkylated dibenzofuran; Oleanane/C₃₀H: oleanane/C₃₀ hopane

All oils are characterized by extremely high Pr/Ph ratios (with an average of 4.01), high abundance of oleananes (Table 1). The abundance of oleananes is likely to be sensitive to changes in Eh and pH during early diagenesis (Murray et al, 1997), but the extremely high oleanane abundances in oils and sediment extracts are usually related to the organic matter input from angiosperm higher land plants (Moldowan et al, 1994; Li et al, 2008c). In the cross plot of Pr/Ph-ADBT/ADBF (Radke et al, 2000), all data points of the oils and condensates from the Fushan Depression fall into the Zone 4, suggesting their terrestrial/deltaic origin.

A total of 45 oils were selected from the Tahe oil field, Tarim Basin, NW China. All the oils mainly occur in the Ordovician carbonates and Triassic sandstones. The majority of the discovered oil accumulations originated from the Middle-Upper Ordovician marine carbonate source rocks (Zhang and Huang, 2005; Wang et al, 2008; Cui et al, 2013). They are characterized by low Pr/Ph and high gammacerane/C₃₀ hopane ratios (Table 2), which indicate anoxic and saline depositional environments of the source rocks. The higher C₂₉/C₃₀ hopane (>0.6) and C₃₅/C₃₄ hopane (>0.8) ratios (Table

2) suggest their marine carbonate origin (Peters et al, 2005). All oils collected from the Tahe oil field fall into Zone 1A of the cross plot of Pr/Ph-ADBT/ADBF (Fig. 2), confirming their marine carbonate origin.

Some oils in the Tahe oil field have undergone severe biodegradation (Wang et al, 2008). The distribution of PAHs may be influenced at advanced stages of biodegradation. For example, with more severe microbial attack, methylphenanthrene isomers of the α -type are more refractory than β -type isomers (Rowland et al, 1986; Radke, 1988). Bao and Zhu (2009) also reported that the relative abundances and distribution patterns may be influenced by secondary alteration processes, such as water washing and biodegradation. Because this study mainly focuses on the DBT related maturity parameters, those samples which have potential biodegradation signatures, such as very low API°, extremely high Pr/nC₁₇ and Ph/nC₁₈ ratios and apparent unresolved complex mixture (UCM) baseline humps of the gas chromatograms are therefore excluded and biodegradation is not considered in the following discussion.

Forty-five oil samples collected from the Termit Basin,

Table 2 Methylated dibenzothiophenes related maturity parameters in oils from the Tarim Basin, NW China

Sample No.	Well	Formation	API, °	4-/1-	4,6-/1,4-	2,4-/1,4-	(2,6+3,6)-/1,4-	TMDBT-I1	TMDBT-I2	MPR	MPI-1	R _c , %	Ts/Tm	22S/(22S+22R)	Pr/Ph	ADBT/ADBF	C ₂₉ /C ₃₀ hopane	C ₃₅ /C ₃₄ hopane
TH01	S1066	O _{3q} -O _{2yj}	26	1.97	1.02	0.51	1.40	1.92	1.46	0.79	0.55	0.82	0.54	0.61	0.74	5.66	0.74	0.69
TH02	S7202	O _{2yj}	31	2.88	1.46	0.66	2.28	2.41	1.66	0.82	0.57	0.83	0.52	0.62	0.73	6.76	0.78	0.92
TH03	TK1010	O _{2yj}	26	2.92	1.48	0.66	2.13	2.27	1.69	0.80	0.57	0.83	0.56	0.61	0.81	7.40	0.90	0.89
TH04	TK241	O _{2yj}	19	3.03	1.68	0.75	2.09	2.22	1.66	0.87	0.58	0.83	0.47	0.61	0.72	6.35	0.88	1.10
TH05	S108	O _{2yj}	33	3.12	1.07	0.53	1.77	1.98	1.51	0.78	0.55	0.82	0.67	0.60	0.68	7.60	0.69	0.80
TH06	TK207	C _{1b} -O _{1,2y}	15	3.24	1.53	0.74	1.79	2.47	1.65	0.81	0.55	0.82	0.39	0.59	0.80	9.40	0.91	1.05
TH07	S116-1	O _{3q} -O _{2yj}	29	3.37	1.69	0.76	2.42	2.56	1.83	0.86	0.62	0.85	0.54	0.61	0.72	7.44	0.80	0.80
TH08	S106-1	O _{3q} -O _{2yj}	34	3.45	1.79	0.72	2.13	2.57	1.77	0.68	0.49	0.80	0.75	0.63	0.83	9.80	0.75	0.80
TH09	S116-2	O _{3q} -O _{2yj}	37	3.69	1.84	0.76	2.48	2.64	1.73	0.81	0.55	0.82	0.89	0.63	0.67	7.75	0.65	0.88
TH10	S1181	O _{3q} -O _{2yj}	38	3.89	1.92	0.82	2.85	2.86	1.93	0.84	0.59	0.84	0.82	0.63	0.77	8.06	0.71	0.95
TH11	S106-3	O _{3q} -O _{2yj}	38	4.09	2.06	0.86	2.87	2.96	2.03	0.80	0.58	0.83	1.04	0.63	0.82	8.60	0.74	0.86
TH12	S106	O _{3q} -O _{2yj}	39	4.24	2.16	0.87	2.74	2.44	1.59	0.69	0.51	0.81	1.22	0.62	0.84	13.45	0.77	1.00
TH13	TK1109	O _{2yj}	40	4.28	2.13	0.90	3.07	3.20	2.17	0.82	0.59	0.84	1.00	0.63	0.88	7.74	0.70	0.73
TH14	TK1108	O _{2yj}	39	4.30	2.09	0.89	2.94	3.27	2.16	0.79	0.58	0.83	1.00	0.63	0.77	8.85	0.69	0.62
TH15	S117	O _{3l} -O _{2yj}	38	4.46	2.09	0.87	3.04	3.07	2.08	0.80	0.58	0.83	1.63	0.65	0.70	8.45	0.70	0.67
TH16	T901	O _{2yj}	34	4.84	2.23	0.95	2.82	3.34	2.05	1.03	0.69	0.88	0.52	0.60	0.91	8.86	0.91	1.14
TH17	S7203	O _{3q} -O _{2yj}	38	5.43	2.65	1.06	3.44	3.61	2.36	1.00	0.70	0.88	1.00	0.66	0.90	10.08	0.70	0.98
TH18	LN44	O	n.d.	5.57	2.77	1.12	3.33	3.65	2.36	1.04	0.72	0.89	1.50	0.66	0.70	9.67	0.72	n.d.
TH19	LN51	O	n.d.	6.12	3.27	1.44	3.40	4.33	2.79	1.20	0.82	0.93	1.94	0.66	0.83	10.21	0.45	n.d.
TH20	T208	O _{1l} -O _{2yq}	18	6.42	3.22	1.30	3.82	3.90	2.50	1.08	0.73	0.89	0.92	0.65	0.65	9.25	0.83	0.83
TH21	S7201	O _{3l} -O _{2yj}	42	6.60	3.51	1.52	4.21	4.99	2.99	1.61	1.02	1.01	0.49	0.62	0.76	11.89	0.95	1.00
TH22	S112-1	O _{3s} -O _{2yj}	39	6.64	3.32	1.33	4.20	4.22	2.67	0.99	0.72	0.89	1.86	0.68	0.63	9.92	0.66	0.74
TH23	T913	O _{3s} -O _{2yj}	38	7.41	3.05	1.34	3.88	4.08	2.69	1.10	0.75	0.90	4.56	0.70	0.87	8.68	0.71	n.d.
TH24	LN54	O	n.d.	7.49	3.11	1.36	4.13	4.53	2.81	1.17	0.78	0.91	2.33	0.70	0.71	9.09	0.69	n.d.
TH25	LG12	O _{1,2y}	n.d.	7.98	3.49	1.54	4.24	4.61	3.08	1.03	0.79	0.91	4.26	0.70	0.85	12.12	0.74	n.d.
TH26	S112	O _{2yj}	41	8.05	3.48	1.32	4.37	4.49	2.74	0.88	0.66	0.86	4.88	0.78	0.68	9.47	0.71	n.d.
TH27	LG13	O	n.d.	10.75	4.28	1.93	5.45	5.98	3.73	1.39	0.85	0.94	6.14	0.72	0.79	9.75	0.68	n.d.
TH28	TK925H	T _{2a} ¹	23	3.96	2.13	0.87	2.13	2.90	1.68	0.79	0.55	0.82	0.50	0.61	0.86	15.65	0.90	0.89
TH29	TK926H	T _{2a}	24	4.19	2.13	0.94	2.55	3.06	2.00	0.90	0.62	0.85	0.46	0.62	0.86	8.34	0.92	1.35
TH30	TK933H	T _{2a} ¹	54	4.27	2.24	0.95	2.40	3.05	1.89	0.92	0.60	0.84	0.48	0.61	0.89	10.64	0.85	1.08
TH31	TK910H	T _{2a} ¹	23	4.32	2.17	0.95	2.55	2.97	2.03	0.89	0.61	0.84	0.50	0.62	0.78	7.76	0.91	1.21
TH32	TK909H	T _{2a} ¹	22	4.40	2.15	0.93	2.57	3.11	2.10	0.92	0.63	0.85	0.45	0.63	0.87	8.64	0.93	0.95
TH33	LN51	T	n.d.	4.43	2.30	1.03	2.58	3.07	1.98	0.91	0.64	0.85	0.63	0.61	0.97	7.26	0.92	0.86
TH34	T901	T _{2a} ¹	34	4.51	2.28	0.98	2.82	3.34	2.05	1.03	0.69	0.88	0.51	0.63	0.81	8.86	0.92	0.98
TH35	TK932H	T _{2a} ¹	52	4.63	2.37	1.01	2.83	3.47	2.24	1.02	0.69	0.88	0.61	0.63	0.90	9.16	0.90	0.87
TH36	TK927H	T _{2a} ¹	46	4.73	2.17	0.99	2.88	3.10	2.12	1.09	0.70	0.88	0.51	0.62	0.86	7.83	0.88	1.18
TH37	TK919H	T _{2a} ¹	37	4.78	2.35	1.00	2.83	3.13	2.04	1.00	0.68	0.87	0.52	0.62	0.72	8.66	0.91	1.00
TH38	TK920H	T _{2a} ¹	34	4.82	2.37	0.98	2.87	3.57	2.06	1.08	0.66	0.86	0.50	0.62	0.93	11.30	0.97	0.86
TH39	S112-2	S _{1k}	38	4.93	2.42	0.92	2.66	3.38	2.01	0.83	0.50	0.80	1.10	0.71	0.87	7.66	0.78	n.d.
TH40	THN3	T _{2a} ³	42	5.54	2.97	1.16	3.23	4.14	2.41	0.99	0.72	0.89	0.86	0.65	0.72	6.28	0.87	1.09
TH41	THN8	T _{2a} ³	60	6.13	3.41	1.12	3.12	4.36	2.10	1.05	0.70	0.88	1.12	0.65	0.93	2.94	0.97	n.d.
TH42	THN9H	T _{2a} ³	54	6.40	3.34	1.17	3.29	3.88	2.03	1.04	0.70	0.88	1.00	0.66	0.79	3.78	0.82	0.73
TH43	LN44	T	n.d.	6.49	3.06	1.21	3.57	3.58	2.09	1.15	0.78	0.91	1.09	0.61	1.00	n.d.	0.93	0.86
TH44	THN2	T _{2a} ³	54	6.62	3.28	1.20	3.54	3.91	2.20	1.15	0.72	0.89	0.84	0.68	0.81	4.76	0.78	n.d.
TH45	THN5H	T _{2a} ³	57	7.10	3.44	1.14	3.41	4.17	2.11	1.13	0.70	0.88	1.04	0.68	0.84	3.84	0.84	n.d.

Note: definition of parameters same as Table 1

Eastern Nigeria (Africa) mainly occur in Eocene sandstones and originated from Mesozoic marine shales. They are characterized by relatively low Pr/Ph and high gammacerane/ C_{30} hopane ratios (Wan et al, 2014). In the cross plot of the Pr/Ph-ADBT/ADBDF (Fig. 2), most of the oil samples fall into the marine shale zone.

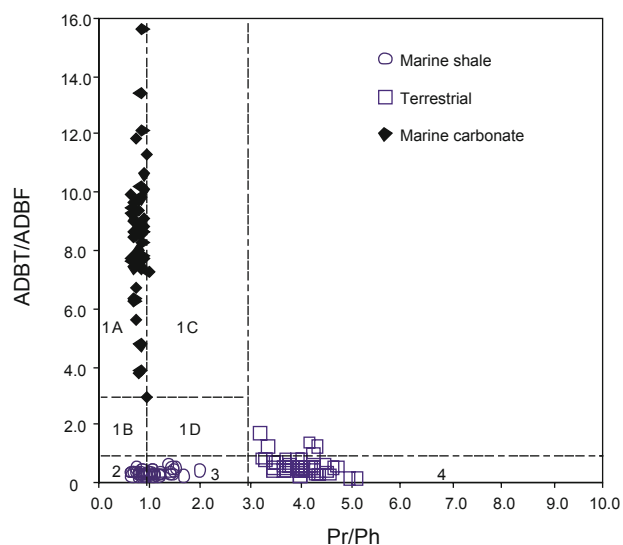


Fig. 2 Concentration ratios of ADBT/ADBDF (alkylated dibenzothiophene / alkylated dibenzofuran) plotted against Pr/Ph (pristane/phytane) for marine shale (○) from Termit Basin, Eastern Nigeria (Africa), terrestrial shale (□) from Beibuwan Basin, South China Sea, and marine carbonate (◆) from the Tarim Basin. Zones indicated by broken lines are used for interpreting depositional environments. Zone 1A: Marine carbonate; Zone 1B: Marine carbonate, marine marl and lacustrine sulfate-rich; Zones 1C and 1D: Mature mudstones and high-rank coals deposited in lacustrine swamp environments with possible marine incursions; Zone 2: Lacustrine sulfate-poor rocks of variable lithology; Zone 3: Marine and lacustrine shales; Zone 4: Fluvial/deltaic carbonaceous shales and coals (Radke et al, 2000)

2.2 Laboratory geochemical analyses

All oils and condensates were deasphalted using *n*-hexane and then fractionated by liquid chromatography using silica gel/alumina columns into saturated and aromatic hydrocarbon fractions using *n*-hexane with dichloromethane (50:50, v:v) and dichloromethane as respective eluents.

The GC-MS analyses of the aromatic hydrocarbon fractions and mixture solutions were performed on an Agilent 5975i GC-MS system equipped with an HP-5 MS (5%-phenyl-methylpolysiloxane) fused silica capillary column (60 m × 0.25 mm i.d., with a 0.25 μm film thickness). The GC operating conditions were as follows: the temperature was held initially at 80 °C for 1 min, ramped to 310 °C at 3 °C/min, and then kept isothermal for 16 min. Helium was used as the carrier gas. The injector temperature was set to 300 °C. The MS was operated in the electron impact (EI) mode with an ionization energy of 70 eV, and a scan range of 50–600 Da.

2.3 Identification of isomers of methylated dibenzothiophenes

Two MDBT isomers MDBTs, i.e. 1-MDBT and 4-MDBT, were identified on *m/z* 198 mass chromatograms

(Connan et al, 1986). Peak identifications for MDBTs in this study were performed by comparison with published relative GC retention time and retention indices (Lee et al, 1979; Vassilaros et al, 1982). Three DMDBT isomers, i.e. 1,4-DMDBT, 2,4-DMDBT and 4,6-DMDBT were identified by co-injection of authentic standards (Budzinski et al, 1991; Chakhmakhchev et al, 1997). A recent study shows that the 1,6-DMDBT isomer may co-elute with 1,4-DMDBT (Li et al, 2012). However, 1,4-DMDBT and 1,6-DMDBT have similar thermodynamic stability (Richard, 2001). Thus, the co-elution of 1,6-DMDBT and 1,4- has no significant influence on the application of the 4,6-/1,4-DMDBT or 2,4-/1,4-DMDBT maturity parameters. By comparison of relative retention times with literature (Asif, 2010) and calculated retention indices (Schade and Andersson, 2006), other DMDBT isomers including 2,6-, 3,6-DMDBT, can be firmly identified in *m/z* 212 mass chromatograms (Fig. 1).

Five trimethyldibenzothiophenes (TMDBTs) have been tentatively detected on *m/z* 226 mass chromatograms (Chakhmakhchev et al, 1997). Their chemical structures, i.e. the specific positions of the three methyl groups, however, have not yet been assigned. Schade and Andersson (2006) specified a series of alkylated dibenzothiophenes through correlation of structure and gas chromatographic retention indices. By correlation with the retention indexes (I_s and I_c) proposed by Schade and Andersson (2006) and by comparison with co-elution of internal standards (Fig. 1) (2,4,7-TMDBT, 2,4,8-TMDBT, 4-E-6-MDBT were purchased from Chiron, Trondheim Norway; 1,3,7- and 1,4,7-DMDBT were purchased from Prof. Jan Andersson's Laboratory, Münster, Germany), a series of trimethyldibenzothiophene isomers were identified in the *m/z* 226 mass chromatograms in this study (Fig. 1).

3 Results and discussion

3.1 Maturity parameters related to dimethyldibenzothiophenes

MDR (4-/1-MDR ratio) proposed by Radke et al (1986) and Radke (1988), 4,6-/1,4-DMDBT and 2,4-/1,4-DMDBT ratios by Chakhmakhchev et al (1997), have been widely used in oil maturity assessments. The MDR has been calibrated against vitrinite reflectance (R_o %), and empirical formulas for estimating vitrinite reflectance (R_c %) have been introduced (Radke, 1988; Dzou et al, 1995; Luo et al, 2001). Two DMDBT based maturity parameters, i.e. 4,6-/1,4-DMDBT and 2,4-/1,4-DMDBT proposed by Chakhmakhchev and Suzuki (1995) and Chakhmakhchev et al (1997) also correlated well with other maturity indicators. All these DBT based maturity parameters are very useful for the characterization of very mature oils lacking biomarkers and work well over a wide range of catagenesis.

All these aforementioned maturity ratios have been calculated for 140 oil samples in this study to examine their interrelationship with each other (Tables 1, 2). For each data set, the 4,6-/1,4-DMDBT ratio correlates well with the MDR and 2,4-/1,4-DMDBT ratio (with correlation coefficients in the range of $R^2 = 0.68-0.93$) and have approximately the same trend (nearly parallel regression lines) (Fig. 3(a) and 3(b)), suggesting their similar behavior with increasing maturity.

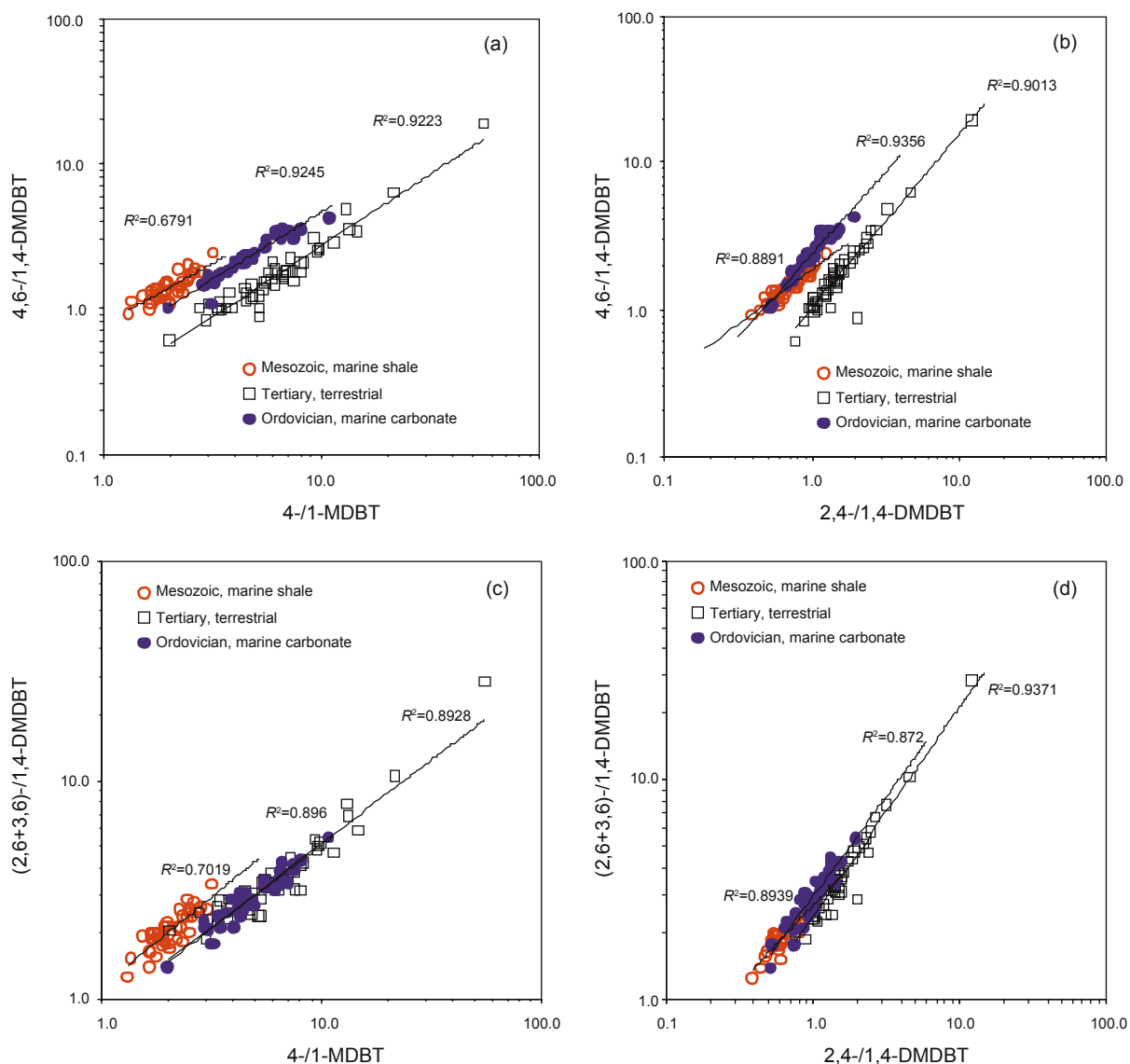


Fig. 3 Cross plots showing a good correlation between (a) 4,6-/1,4-DMDBT and 4-/1-MDBT ratios, (b) 4,6-/1,4-DMDBT and 2,4-/1,4-DMDBT ratios, (c) (2,6 + 3,6)-/1,4-DMDBT and 4-/1-MDBT ratios, and (d) (2,6 + 3,6)-/1,4-DMDBT and 2,4-/1,4-DMDBT ratios

In order to study the influence of organic facies on DBT based maturity parameters, some other maturity indicators must be chosen to serve as reference. Most of the sterane or terpane isomerization maturity parameters, such as C_{31} hopane $22S/(22S+22R)$, however, are not good candidates due to their low concentrations or having reached their isomerization equilibrium point in high mature oils or condensates (Peters et al, 2005). Some aromatic hydrocarbon based parameters, such as methylphenanthrene index (MPI-1) show a strong linear correlation with increasing vitrinite reflectance and are often related to R_o % for the principal phase of oil formation (Radke and Welte, 1981; Radke et al, 1982; Farrington et al, 1988; Radke, 1988; Radke et al, 1997). The empirical relationships derived from regression analysis (Radke and Welte, 1981; Radke, 1988) have been widely used in the calculation of equivalent vitrinite reflectance (R_c %), especially for high mature oils or carbonate sequences without vitrinite of good quality. In this study, the R_c (%) for all oils and condensates were calculated using the empirical relationship R_c (%) = 0.40

+ 0.6 × MPI-1 (for 0.65% to 1.35% R_m) (Radke and Welte, 1981; Radke, 1988).

Fig. 4 illustrates the relationship between MDR (4-/1-MDBT) and R_c (%). The MDR exhibits an exponential relationship with R_c (%). MDR increases slightly with maturity in the 0.7%–1.0% range and increases rapidly at advanced thermal stage (R_c = 1.0%–1.35%), approximately consistent with previous observations (Schou and Myhr, 1988; Dzou et al, 1995). It is important to note that the oils and condensates in this study originated from three different types of source rocks yet have nearly same MDR relationship with R_c (%). This suggests that lithologies and organic facies may have no significant influence on the 4-MDBT and 1-MDBT distributions in this maturity range. Although organic matter type and lithology of source rocks may affect the distributions of DBTs at early stages of thermal transformation (Radke et al, 1986; Connan et al, 1986; Huang and Pearson, 1999; Li et al, 2013b), the influence seems to reduce to a negligible level within the main oil window to post and over-mature stages.

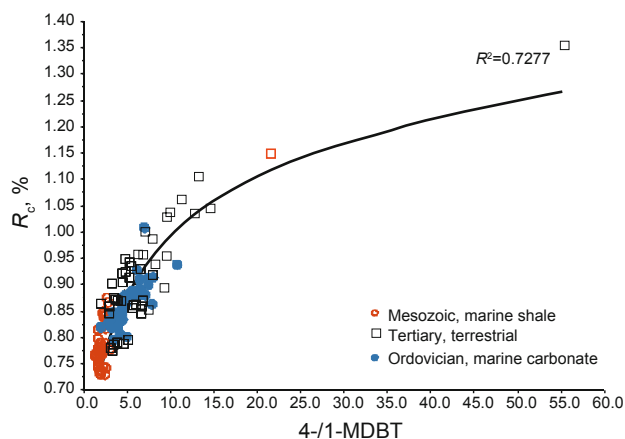


Fig. 4 Relationship of 4-/1-MDBT (MDR: 4-/1-methyldibenzothiophene) to vitrinite reflectance R_c (%) calculated by MPI-1 (methylphenanthrene index) for oils in this study. R_c (%) = $0.4 + 0.6 \times \text{MPI-1}$ for $\text{MPR} < 2.65$. $\text{MPR} = 2\text{-}/1\text{-methylphenanthrene}$

The ratio of 4,6-/1,4-DMDBT shows a good correlation with MDR for each data set (Fig. 3(a) and 3(b)) and also exhibits an exponential relationship with R_c (%) (Fig. 5). In this study, oils derived from marine carbonate and marine shale source rocks do not exhibit significant differences in the relationship between 4,6-/1,4-DMDBT and MDR or 2,4-/1,4-DMDBT (Fig. 3(a) and 3(b)). However, oils from terrestrial source rocks seem to have relatively lower 4,6-/1,4-DMDBT (Fig. 3(a) and 3(b), Fig. 5). Compared with oils from marine carbonate and shale source rocks, the oils from terrestrial Beibuwan Basin are systematically 0.5 lower (Fig. 5). This suggests that organic matter type and lithology of source rocks affect the 4,6-/1,4-DMDBT ratio even at advanced maturity stage. Thus, we cannot compare the maturity of oils originated from source rocks with different organic matter types and lithologies based solely on the 4,6-/1,4-DMDBT ratios. Calibration by other maturity parameters or investigation of DBT distribution in related source rocks seems to be necessary.

The DBT maturity parameters are based on the theoretical differences in thermodynamic stability of various methyl

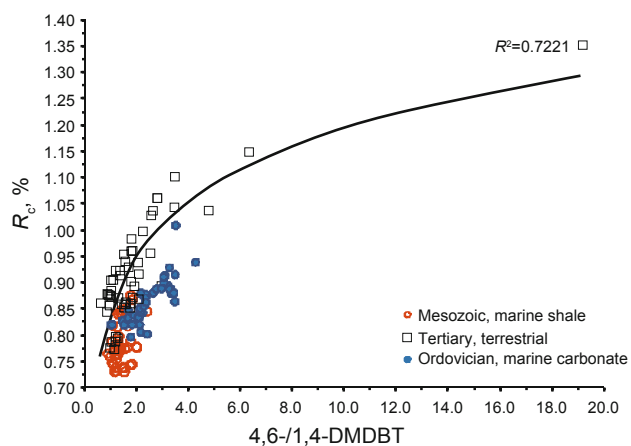


Fig. 5 Relationship of 4,6-/1,4-DMDBT (4,6-/1,4-dimethyldibenzothiophene) to vitrinite reflectance R_c (%) calculated by MPI-1 (methylphenanthrene index) for oils in this study

DBT isomers. The enthalpy of formation (ΔH) is a commonly used thermodynamic parameter, and the lower the value, the greater the stability of that isomer (Budzinski et al, 1993). For example, the enthalpy of formation for 4-MDBT is 23,755 cal/mol, which is 5,979 cal/mol lower than that of 1-MDBT, and the value for 4,6-DMDBT is 14,355 cal/mol, which is also 5,979 cal/mol lower than that of 1,4-DMDBT (Richard, 2001). The enthalpy of formation for 2,4-DMDBT is 14,156 cal/mol, which is a little lower than that of 4,6-DMDBT. The isomers 2,6-, 3,6-DMDBT have the same ΔH values as that of 2,4-DMDBT (Richard, 2001), suggesting their similar thermodynamic stability. Thus, 2,6-/1,4- or 3,6-/1,4-DMDBT ratios are also potential maturity indicators similar to 2,4-/1,4-DMDBT. However, these two isomers cannot be completely separated in routine GC-MS analysis (Figs. 2 and 6). Thus, we propose a new maturity parameter $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ based on the relative abundance of DMDBT isomers.

Oil samples originated from typical terrestrial, marine carbonate and marine shale source rocks demonstrate that the newly proposed $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ ratio correlates well with the MDR and 2,4-/1,4-DMDBT (Fig. 3(c) and 3(d)) (with correlation coefficients in the range of 0.70 to 0.90, and 0.87 to 0.94, respectively), suggesting its similar behavior with increasing maturity. Like MDR, the $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ ratio for each data set also has similar trends with that of MDR or 2,4-/1,4-DMDBT (Fig. 3(c) and 3(d)). In contrast with 4,6-/1,4-DMDBT, $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ seems to have minor differences among oils with different origins, which can be clearly demonstrated in the relationship between $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ and R_c (%) (Fig. 7). Like 4,6-/1,4-DMDBT, $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ also has an exponential relationship with R_c (%). However, unlike 4,6-/1,4-DMDBT, oils with different origins seem to exhibit a similar trend, suggesting that lithology and organic facies seem to have minor influences on the $(2,6 + 3,6)\text{-}/1,4\text{-DMDBT}$ ratio for high mature oils. We can also infer that the organic facies and lithology may have different influences on different DBT based maturity parameters.

3.2 Maturity parameters related to trimethyldibenzothiophene

One maturity parameter TMDBT index based on trimethyldibenzothiophene (TMDBT) was previously proposed by Chakhmakhchev and Suzuki (1995). The ratio of two isomers in the m/z 226 mass chromatogram increases with maturity and has a good correlation with MDR, suggesting its similar behavior with increasing maturity. However, this empirical index has not been widely used for the unknown structures of these TMDBT isomers. They are simply named peak 3 and peak 5 (Fig. 1), respectively by Chakhmakhchev et al (1997). The progress on GC retention behavior of DBT on different stationary phase columns (Mössner et al, 1999; Schade and Andersson, 2006) provides a relatively reliable way to identify more isomers of trimethyldibenzothiophene. A series of TMDBT isomers were identified in oils, coal and sedimentary rock extracts (Li et al, 2012) by comparing with retention indices reported in literature. In this study, some isomers were identified by comparing with the separation experiments of TMDBT isomers (Mössner et al, 1999), the

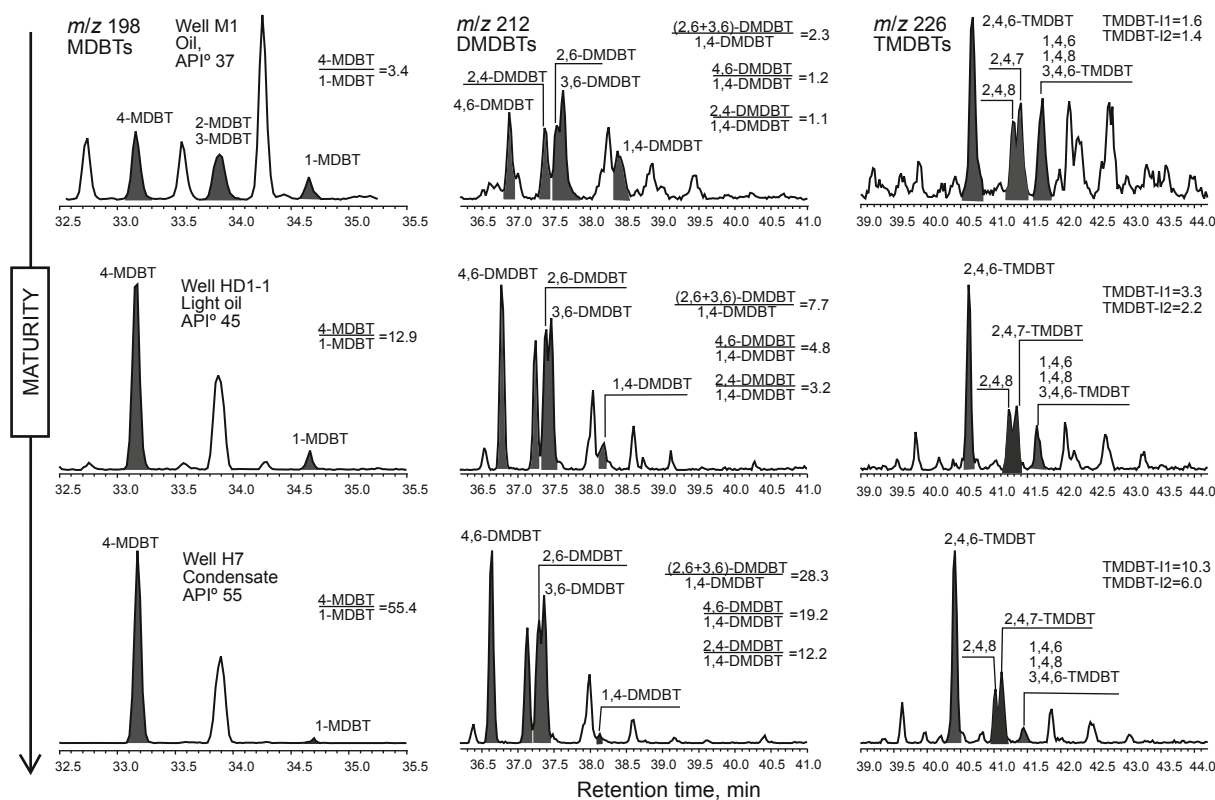


Fig. 6 Maturity effect on the distribution of methyl-, dimethyl-, and trimethyldibenzothiophenes (m/z 198, 212, and 226). Oil samples are from Beibuwan Basin, South China Sea. Note that 1,6-DMDBT isomer may co-elute with 1,4-DMDBT. Due to their similar thermodynamic stabilities, there is no significant influence on the application of 2,4-/1,4-DMDBT and 4,6-/1,4-DMDBT as maturity indicators

retention indices (Schade and Andersson, 2006) and co-elution of some internal standards. For example, the peak 3 in the m/z 226 mass chromatogram by Chakhmakhchev et al (1997) is assigned as 2,4,6-TMDBT, and peak 5, as 1,4,6-, 1,4,8- and 3,4,6-TMDBT (Fig. 1). Richard (2001) did not provide the thermodynamic parameters for TMDBT isomers in his study. However, the thermodynamic stability for some isomers can be reasonably inferred. Budzinski et al (1991) reported that the 1-position is less stable than the 3-position which is less

stable than the 2-position, and the positions 4- and 6- are more stable. Thus the 4,6-, 2,6- and 3,6-DMDBT isomers are more stable and are the main isomers in mid-mature to high-mature oils. According to this rule, the 2,4,6-TMDBT may be the most stable isomer among all TMDBT isomers. Our observations also suggest that peak 3, i.e. 2,4,6-TMDBT for all oils in this study is the most abundant isomer (Figs. 1 and 6). Therefore, the empirical TMDBT index proposed by Chakhmakhchev and Suzuki (1995) and Chakhmakhchev et al (1997) is reliable on a thermodynamic basis.

The TMDBT index (TMDBT-II in our study) here defined as $2,4,6-(1,4,6 + 1,4,8 + 3,4,6)$ -TMDBT for 140 oil samples from three basins is listed in Tables 1 and 2. It also exhibits a good correlation with MDR and 2,4-/1,4-DMDBT ratio with coefficients in the range of 0.76 to 0.91 for MDR and 0.75 to 0.93 for 2,4-/1,4-DMDBT (Fig. 8(a) and 8(b)). Unlike 4,6-/1,4-DMDBT or $(2,6 + 3,6)$ -/1,4-DMDBT, oils from terrestrial and marine shale have similar trends with MDR or 2,4-/1,4-DMDBT, and they are relatively lower than those of marine carbonate sourced oil. Thus, we cannot compare the maturity of oils with different origin by TMDBT-II without any calibration. The distributions of methylated PAHs in oils and sedimentary organic matter are controlled by complicated chemical processes, such as methylation, demethylation, and disproportionation reactions (methyl shifting) (Alexander et al, 1985; Strachan et al, 1988; van Aarssen et al, 1999; Bastow et al, 2000). Thus, further study is needed to investigate the formation and transformation of polycyclic aromatic sulfur hydrocarbons with the increasing thermal

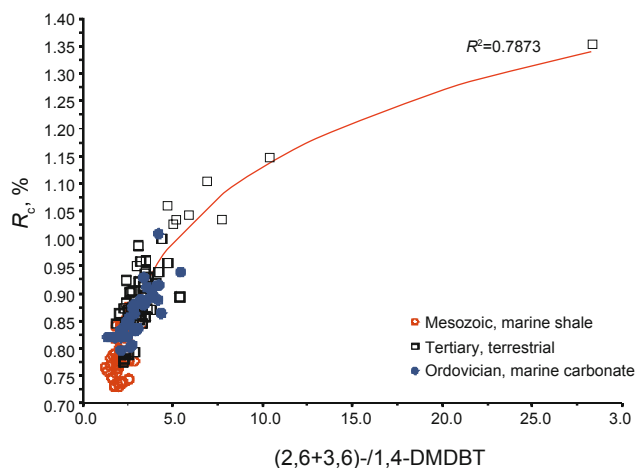


Fig. 7 Relationship of $(2,6 + 3,6)$ -/1,4-DMDBT ($(2,6 + 3,6)$ -/1,4-dimethyldibenzothiophene) to vitrinite reflectance R_c (%) calculated by MPI-1 (methylphenanthrene index) for oils in this study

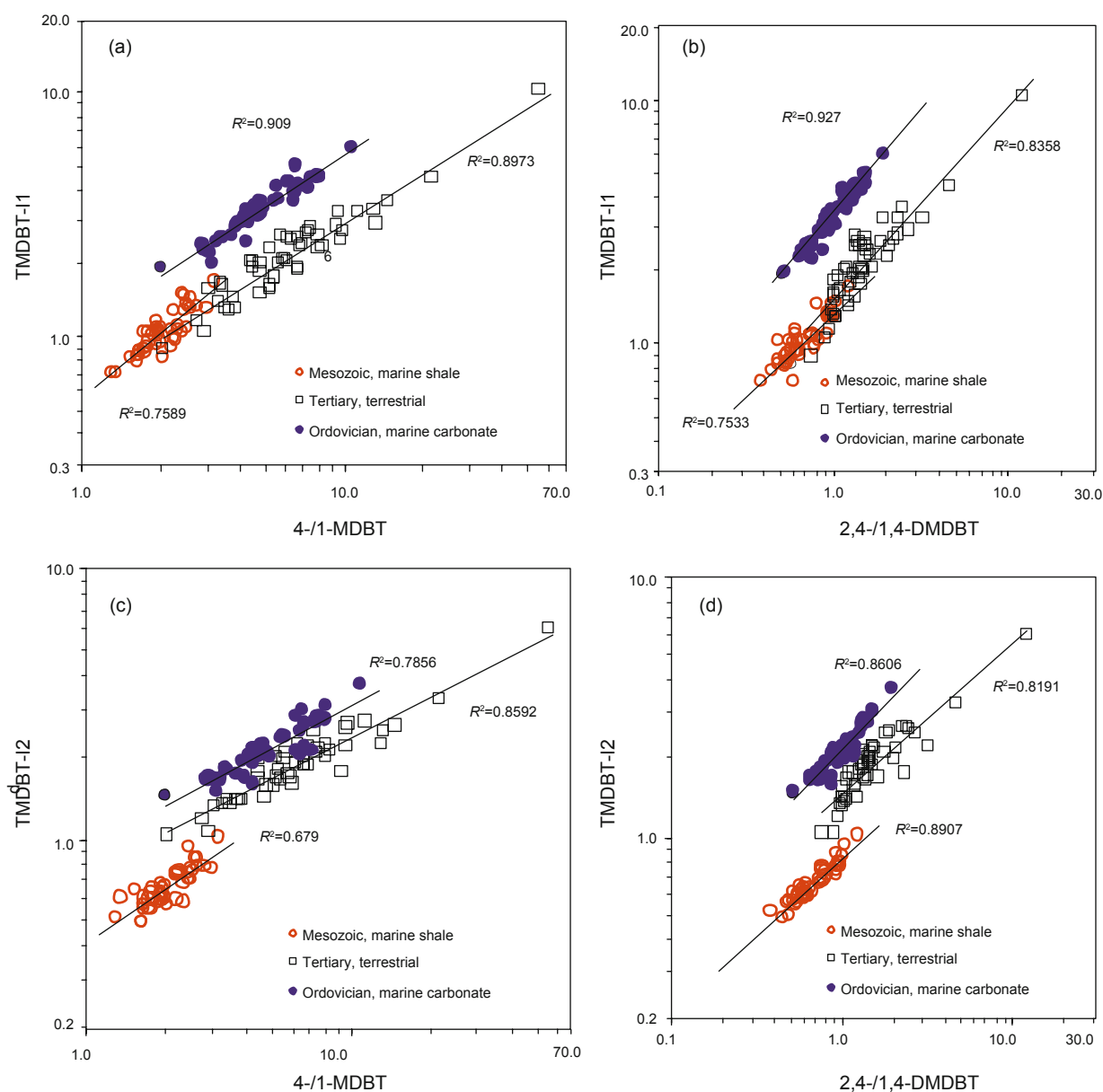


Fig. 8 Cross plots showing a good correlation between (a) TMDBT-I1 and 4-/1-MDBT ratio, (b) TMDBT-I1 and 2,4-/1,4-DMDBT ratio, (c) TMDBT-I2 and 4-/1-MDBT ratio, and (d) TMDBT-I2 and 2,4-/1,4-DMDBT ratio

stress, especially during diagenesis.

Other TMDBT related ratios applicable for maturity assessments can be proposed on the basis of TMDBT thermal stability. This study proposes a new TMDBT maturity parameter TMDBT-2, which is defined as (2,4,8 + 2,4,7)-/(1,4,6 + 1,4,8 + 3,4,6)-TMDBT ratio (Tables 1 and 2, Fig. 6). Data of 140 oil samples in our study demonstrate that it correlates well with other DBT maturity parameters, such as MDR or 2,4-/1,4-DMDBT (Fig. 8(c) and 8(d)) with correlation coefficients in the range of 0.68 to 0.86, 0.82 to 0.89, respectively. Our study together with previous studies (e.g. Chakhmakhchev et al, 1997) shows that all thermodynamic stability based DBT maturity parameters correlate well with each other, suggesting their similar behavior with increasing maturity. Like TMDBT-I1 and 4,6-/1,4-DMDBT ratio, TMDBT-I2 is also related to the organic facies or/and depositional environment (Fig. 8(c)

and 8(d)). Thus, the same TMDBT-I2 values for oils from different basins may not indicate similar thermal maturity.

The maturity parameters based on the distributions of DBT work well over a relatively wide range of maturity. They are effective maturity indicators for oils with moderately mature to high mature petroleum, especially at an advanced stage of thermal transformation (Chakhmakhchev et al, 1997). They are usually present in significant concentrations and can be easily identified and measured in light oils and condensates depleted in biomarkers. However, with increasing maturity, the thermodynamically unstable isomers such as 1-MDBT may be absent or present in extremely low concentrations below the detection limit in GC-MS analyses. In our study, the 1-MDBT and 1,4-DMDBT isomers for condensate from Well H7 in the Beibuwan Basin are present in extremely low concentration (Fig. 6), which may result in inaccurate values of the MDR and other 1,4-DMDBT related maturity

parameters.

The MPI-1 of this condensate is as high as 1.59 and the MPR (2-/1-methylphenanthrene) gets to 3.77 (Fig. 9). According to Radke and Welte (1981) and Radke et al (1986), when the R_c is higher than 1.35% or $MPR > 2.65$, the MPI-1 shows a reverse trend with the increasing maturity and the R_c (%) could be calculated using another empirical formula R_c (%) = $2.3 - 0.6 \times MPI-1$. Here we found that the R_c (%) values calculated by both equations are nearly the same (1.35%). It may just correspond to the inflection point for MPI against R_c relationship (Fig. 9). This coincides with the upper level of maturity that can be evaluated by DBT ($R_o \approx 1.35\% - 1.5\%$ by Chakhmakhchev et al, 1997, and $R_o = 1.35\%$ by Radke and Welte, 1981).

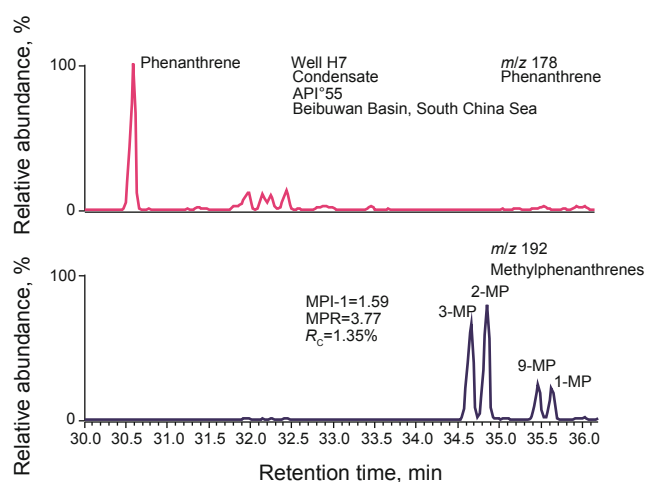


Fig. 9 The m/z 178 and 192 mass chromatograms for the condensate from Well H7 in the Beibuwan Basin, South China Sea show the distribution of phenanthrene and methylphenanthrenes.

Note: 3-MP: 3-methylphenanthrene; $MPI-1 = 1.5 \times (\text{Phenanthrene} + 2\text{-MP} + 3\text{-MP}) / (\text{Phenanthrene} + 9\text{-MP} + 1\text{-MP})$; $MPR = 2\text{-MP} / 1\text{-MP}$; R_c (%) = $2.3 - 0.6 \times MPI-1$ (Radke and Welte, 1981)

The TMDBT-based maturity parameter TMDBT-I2 proposed in this study and previous TMDBT-I1 appear to be more reliable for over-mature light oils and condensates. The thermodynamically unstable 1,4,6-, 1,4,8- and 3,4,6-TMDBT isomers have relatively higher concentration (Fig. 2) at high maturity.

Up to now, there is little knowledge about the formation of DBTs during diagenesis. Like most of the PAHs in oil and ancient sedimentary organic matter, DBTs do not have certain biological precursors. They were mainly produced by reaction with other compounds, e.g. biphenyl and sulfur (Asif et al, 2009). Understanding the formation and evolution processes of DBTs during diagenesis will be very important to the application of maturity parameters based on DBT in low mature oils.

4 Conclusions

The structures of a series of trimethyldibenzothiophenes (TMDBTs) have been firmly assigned by comparison with reported retention indices and GC-MS separation experiments in this study. A previously proposed TMDBT index that is

based on the relative abundances of two TMDBT isomers (peak 3 to peak 5) with unknown structures is here identified as four isomers and defined as 2,4,6-/(1,4,6 + 1,4,8 + 3,4,6)-TMDBT. Another maturity parameter (2,4,7 + 2,4,8)-/(1,4,6 + 1,4,8 + 3,4,6)-TMDBT based on TMDBT is proposed in light of the same thermodynamic principles. Both these indices (TMDBT-I1 and TMDBT-I2) correlate well with the MDR and 2,4-/1,4-DMDBT, displaying their high potential for maturity evaluation, especially at elevated maturity stages.

A new maturity parameter (2,6 + 3,6)-/1,4-DMDBT relative to dimethyldibenzothiophene is also proposed based on their same thermodynamic stability as that of 2,4-DMDBT isomers. This ratio exhibits a good correlation with MDR and 2,4-/1,4-DMDBT, suggesting its similar behavior with increasing maturity. Unlike 4,6-/1,4-DMDBT ratio, oils with different origins seem to exhibit a similar trend, suggesting that lithology and organic facies seem to have minor influence on the (2,6 + 3,6)-/1,4-DMDBT ratio for high mature oils.

DBTs distribution indicators would be widely applied in maturity assessments in oil and source rocks since they appear to work well over a relatively high range of maturity, especially for over-mature light oils and condensates lacking biomarkers.

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