



# Micro-meso-macroscale correlation mechanism of red-bed soft rocks failure within static water based on energy analysis

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## Abstract

The purpose of the present study is to fundamentally investigate the micro-meso-macroscale correlation mechanism of red-bed soft rocks failure within static water, based on energy analysis. Through a series of experimental studies, changes in the micro-meso-macroscale characteristics of the rocks were characterised by measuring ion concentrations in and pH of water, the pore structure and crack propagation. Based on the experimental measurements, theoretical modelling of micro-meso-macroscale energy conservation was carried out to determine the micro-meso-macroscale correlation mechanism of red-bed soft rocks failure. The results show that the microscale dominant effect of red-bed soft rocks is determined as chemical effects. The failure mechanism of red-bed soft rocks under static water saturation involves chemical, physical and mechanical dominant effects at the micro-meso-macroscale. It was demonstrated that the dominant chemical effect of clay minerals and water comprises a series of chemical energy changes in the microscale water–rock interface. This can induce the energy changes in both mesophysical and macromechanical levels. During the first month after the water saturation, there was a significant decrease in  $\text{Na}^+$  and  $\text{K}^+$  concentrations in aqueous solution, whereas there was little change in  $\text{Ca}^+$  and  $\text{Mg}^{2+}$  concentrations. The scanning electron microscope (SEM) images indicated an obvious change in microstructure of the red-bed soft rocks after 3 months of water saturation. A decrease in pore number and increase in total pore surface area could be seen after 6 months of water saturation. Furthermore, the theoretical results suggest that chemical effects at the microscale dominate the softening process of the rocks. During the water–rock interaction, nearly 80% of the energy dissipation results from the microscale change, and the meso-macroscale change accounts for the remaining 20%. As a uniform measure of material evolution, energy can correlate the micro-meso-macro dominant effect of red-bed soft rocks softening appropriately.

**Keywords** Energy conservation · Micro-meso-macroscale correlation · Pore structure · Red-bed soft rocks · Static water saturation · SEM imaging

## List of symbols

$\Delta H$	Enthalpy change
$Q_{\text{micro}}$	Microscale energy
$Q_{\text{meso}}$	Mesoscale energy
$Q_{\text{macro}}$	Macroscale energy
$Q_{\text{total}}$	Total scale energy

$Q_{\text{sum}}$	Total scale energy subtract overlapping energy
$Q_{\text{conservation}}$	Theoretical modelling of energy conservation under water–rock interaction
$Q_{\text{overlap}}$	Overlapping energy of meso and macro
$Q_1$	Adsorption energy
$Q_2$	Chemical energy
$n$	Molar amount of adsorbed ions in solution
$E_a$	Unit molar adsorption energy
$c_i$	The molar concentration of $i$ -th reaction
$V_{r/w}$	Volume of sample or volume of aqueous solution
$\Delta_r H_m^0$	Standard molar enthalpy change of $i$ -th reaction
$E_b$	Energy of added surface area after softening and disintegration of red-bed soft rocks

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$G_{1c}$	Energy per unit surface of red-bed soft rocks	$\Delta T_{\text{rock}}$	Temperature change of red-bed soft rocks
$A_{\text{surface}}$	Newly increased surface area of red-bed soft rocks	$\Delta T_{\text{water}}$	Temperature change of water
$A_{\text{particle}}$	Surface area of particles	$c_{\text{rock}}$	Specific heat capacity of silty mudstone
$Q_v$	Swelling energy	$c_{\text{water}}$	Specific heat capacity of water
$V_1$ and $V_2$	Volumes before and after swelling	$E_1$	Elastic modulus of red-bed soft rocks
$P$	Swelling stress	$\alpha$	Coefficient of thermal expansion
$V$	Volume of red-bed soft rocks	$\varepsilon$	Thermal expansion strain
$W_e$	Elastic property of soft rocks with volume $V_v$	$d$	Particle diameter of red-bed soft rocks
$\mu$	Poisson's ratio of red-bed soft rocks	$e$	Sum of porosity and volume change rate
$E$	Elastic modulus of red-bed soft rocks		
$W_1$	Fracture energy caused by crack		
$W_2$	Fracture energy caused by voids		
$W_f$	Energy dissipated due to initial crack slip		
$U_e$	Energy dissipated due to tensile crack propagation		
$W_V$	Fracture energy of voids when pore volume is $V_{\text{matrix}}$		
$2c$	Initial crack length		
$2l$	Tensile crack length		
$2w$	Distance between adjacent cracks		
$\theta$	Angle between initial crack surface and axial stress		
$\tau_f$	Shear force produced by the external load sliding on initial crack surface		
$\tau^*$	Shear force produced by the external load sliding on tensile crack surface		
$\nu$	Friction coefficient between crack surfaces		
$E$	Elastic modulus		
$k$	A constant, in plane stress state		
$\sigma_{\text{eq}}$	Equivalent stress		
$\varepsilon_{\text{eq}}^p$	Equivalent plastic strain		
$U^s$	Work on shear stress		
$U^f$	Work on friction		
$U^c$	Work for complete cutting on convex surface		
$U^n$	Work for normal load		
$U^v$	Work on viscous drag		
$V_{\text{fluid}}$	Fluid velocity		
$A$	Frictional surface area		
$l_i$	Frictional distance		
$U^c$	Work when the concave convex part is completely sheared during process of shearing		
$U_p$	Frictional energy of particles		
$F_s$	Frictional force of collision friction between particles		
$F_n$	Average impact force on the indirect contact surface of particles		
$\Delta t$	Time of particle collision and friction		
$V_{\text{water}}$	Volume of water in container		
$m_{\text{sample}}$	Quality of red-bed soft rocks sample		
$\Delta H_{\text{resultant}}$	Standard molar enthalpy of resultant		
$\Delta H_{\text{reactant}}$	Standard molar enthalpy of reactant		

## 1 Introduction

Although red-bed soft rocks demonstrate good mechanical properties, a series of physical, chemical and mechanical effects can occur if the rocks contact water, which may lead to their mechanical failure. To date, the study on softening behaviour of soft rocks mainly focuses on the external representation and the correlation of macromechanics and mesophysical effects, while the research work on microscale mechanism of the rocks is not cohesive [1, 15, 22, 24, 38, 46, 70, 71, 73].

Previous experimental studies on the softening processes of the soft rocks have mainly focussed on the decomposition and transformation of a single type of mineral [20, 21, 40, 55], while the change in mesoparticles and pores in soft rocks was investigated using scanning electron microscope (SEM) and high-definition digital imaging [36, 37, 54, 69]. Although a series of mechanical tests of soft rocks (e.g. consolidated undrained and drained triaxial testing) have been carried out [14, 17, 27], mechanical testing with consideration of water effect needs to be further investigated. The relationship between the macroscale mechanical behaviour and mesoscale structure of soft rocks was experimentally studied using the combination of scanning electron microscope, computed tomography scanning, digital image correlation and advanced imaging techniques [31, 53, 66]. However, changes at the microscale of soft rocks due to chemical reactions under the action of water are not fully understood.

The current theoretical models on studying the deformation and failure mechanism of soft rocks based on fractal and energy dissipation theories have limited capabilities to determine the relationship between the mechanics and structure at mesoscale [47, 49, 61]. The plastic deformation stress model, continuous damage model, FLAC3D rheological constitutive model and anisotropic damage model of soft rocks were developed for investigating the relationship between macroscale and mesoscale dominant effects during

the softening process [4, 8, 13, 23, 72]. A series of saturated soft rocks models were developed based on, for example, the evolutionary neural network, renormalisation group theory, damage fractional theory, free entropy theory and viscoelastic plastic creep theory [25, 28, 74], but the physical–chemical interaction was not studied. A range of contact models, crack propagation models and damage evolution models were established to study the energy dissipation of rock failure [6, 50, 57]. However, it narrowly focuses on the transformation and dissipation of damage energy under deformation and failure such as elastic–plastic properties and surface energy under macroscale stress–strain strength. In short, the study of soft rocks softening is based on external representation and the correlation of macromechanics and mesophysical effects. Literature on the microscale mechanism is not cohesive. The process of energy balance change under the action of water and rock is deeply undiscussed.

The purpose of this study is to investigate the micro-meso-macroscale correlation mechanism of red-bed soft rocks failure within static water, based on energy, by conducting a series of experimental studies in conjunction with theoretical modelling. First, the dominant factors (ion concentrations in and pH of water, the pore structure and crack propagation) that govern changes in the micro-meso-macroscale characteristics of red-bed soft rocks were experimentally studied. The micro-meso-macroscale energy conservation and failure mechanism were then theoretically investigated.

## 2 Materials and method

Red-bed soft rocks samples obtained from Southern China were used for the experimental testing. Figure 1 shows the methodology in this study for establishing the micro-meso-macroscale relationship of soft rocks under water–rock interaction.

### 2.1 Experimental studies

#### 2.1.1 Preparation of red-bed soft rock samples

Typical silty mudstone samples were collected from Southern China. The main clay mineral content in the rock samples was kaolinite, illite and quartz. There was a small amount of sericite, chlorite, muscovite, montmorillonite and iron. The size of cylindrical silty mudstone samples is  $\phi 50$  mm  $\times$  100 mm. The samples are tested by sound wave velocity, and the samples with large dispersion of sound wave velocity are eliminated. After extracting the defective samples, 15 samples were ultimately got (Fig. 2). The rock samples were placed into a plastic box (250 mm  $\times$  250 mm  $\times$  250 mm) containing water taken from the Dongjiang

River basin, Southern China. The tops of the samples were 10 cm below the water surface. To prevent excessive water evaporation, the plastic box was fully covered.

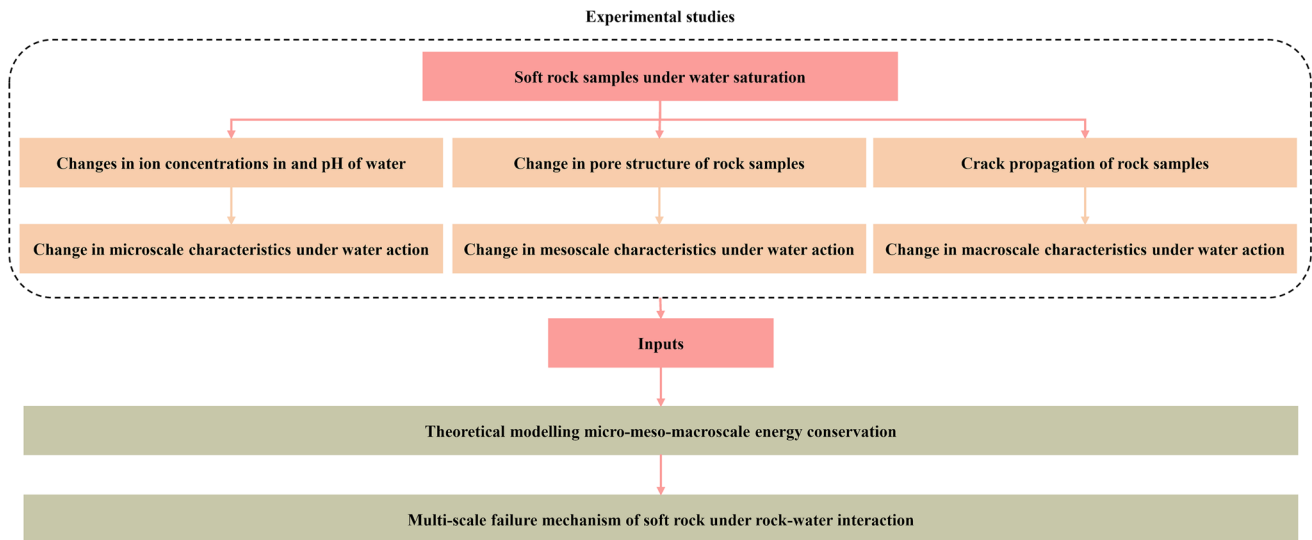
The test conditions for the samples were (I) natural condition, (II) water saturation for 1 month, (III) water saturation for 3 months, (IV) water saturation for 6 months and (V) water saturation for 12 months. To reduce the uncertainty of experimental measurements, three samples were used for each of the tests, as shown in Table 1. The samples were described as I (1,2,3), II (1,2,3), III (1,2,3), IV (1,2,3) and V (1,2,3).

#### 2.1.2 Measuring ion concentrations in and pH of water solution

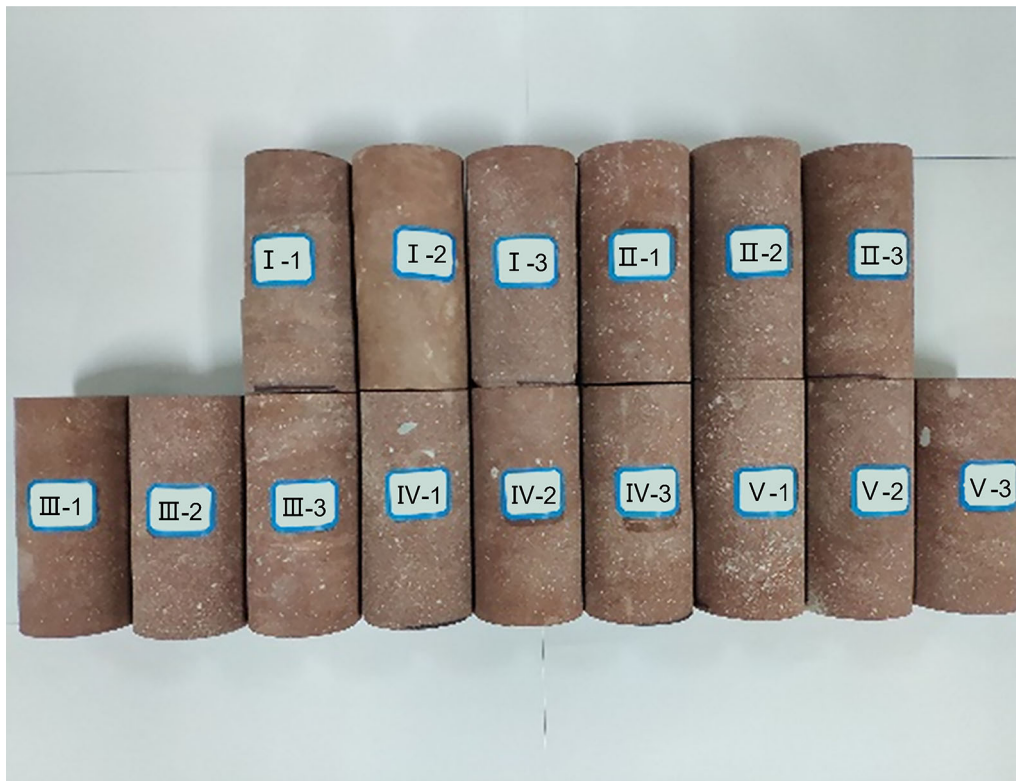
The time-dependent change in ion concentrations (e.g.  $\text{Na}^+$ ,  $\text{Ca}^{2+}$  etc.) in and pH of water solution for each test was measured using ion chromatograph (Model: Dionex ICS-900, Resolution: 0.05 mg/L) and pH reagent, respectively. Regularly take a small amount of rock blocks from the same point of soft rock samples. A small amount of rock blocks was ground into powder and passed through 0.5-mm sieve after naturally drying. The powder is packed in self-sealing bags and numbered. The time-dependent change in chemical components of rock samples under water saturation was studied using inductively coupled plasma emission spectrometer (Model: iCAP6500Duo, Resolution: 0.001 mg/L) in the Instrumental Analysis and Research Center, Sun Yat-sen University. Because of the massive difference in the magnitude of each chemical composition in soft rock, Fig. 5 is drawn by percentage as ordinate. To reduce the uncertainty of experimental measurements, the average value of the measurement data of 3 samples in each group is taken as a point of each test scene (Fig. 5 and Fig. 6). The purpose is to characterise the microbehaviour of red-bed soft rocks samples under static water action.

#### 2.1.3 Measuring pore structure of rock samples

Regularly take a small amount of rock blocks from the same point of soft rock samples. A small amount of rock blocks was fixed on the sample table after freeze drying. Gold was sprayed on the surface of the rock blocks to improve the quality and resolution of the image. Then the time-dependent change in pore structure of rock samples under water saturation was studied using SEM (Model: Gemini 500, Resolution: 0.6 nm) in the Instrumental Analysis and Research Center, Sun Yat-sen University. The obtained SEM images were processed using the SEM Image Processing Tool. The mesobehaviour of the red-bed soft rocks samples under water action can be quantitatively characterised by SEM.



**Fig. 1** Flow chart of experimental and theoretical methodology for investigating the failure mechanism of red-bed soft rocks under water–rock interaction



**Fig. 2** Cylindrical silty mudstone samples

### 2.1.4 Monitoring crack propagation of rock samples

The crack propagation of rock samples resulting from steady-state water–rock interaction was monitored using visual inspection (high-speed camera), so that the macrobehaviour of rock samples could be characterised.

### 2.2 Theoretical modelling

The experimentally measured change in micro-meso-macrocharacteristics of red-bed soft rocks samples under water action was theoretically analysed by modelling enthalpy change based on energy conservation with the aim

**Table 1** Test conditions for red-bed soft rock samples

Test number	Sample number	Test scenario
I	I-1	Natural condition
	I-2	
	I-3	
II	II-1	Water saturation for 1 month
	II-2	
	II-3	
III	III-1	Water saturation for 3 months
	III-2	
	III-3	
IV	IV-1	Water saturation for 6 months
	IV-2	
	IV-3	
V	V-1	Water saturation for 12 months
	V-2	
	V-3	

of establishing the relationship of micro-meso-macroscale changes in soft rocks characteristics under water action.

### 2.2.1 Modelling change in microscale characteristics of red-bed soft rocks under water action

Based on previous studies [66, 67, 70], the mesoscale dominant effect is a physical effect, and the macroscale dominant effect is a mechanical effect. On this basis, the microscale dominant effect can be determined.

The scale of red-bed soft rocks is divided mainly granting to the size of the selected objects in Physics (Fig. 3).

**Microscale:** it primarily studies the change process of functional groups, mineral grain structure, solution pH and ion concentration. The microeffects of rocks were analysed by microtests such as Fourier transform microinfrared spectrometer, ion chromatograph and pH reagent.

**Mesoscale:** it is reflected in mesodamage like dislocation friction between clay particles, volume expansion between particles, expansion and accumulation of pore cracks. It was observed by optical microscope or scanning electron microscope to study its mesoeffect.

**Macroscale:** it mainly includes mechanical damage, swelling, rock mass falling off and disintegration in the laboratory. The macroscale rock and soil mass are studied by indoor and field test equipment, and its deformation and mechanical properties are analysed.

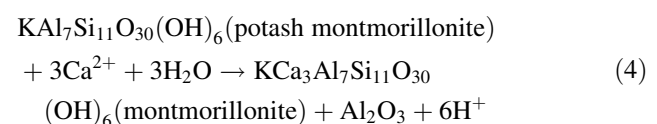
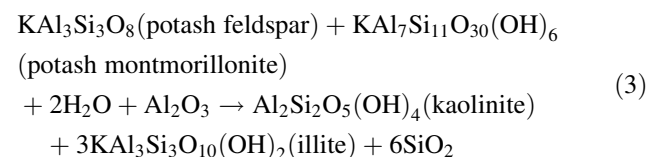
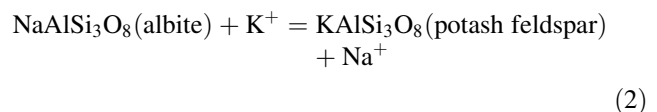
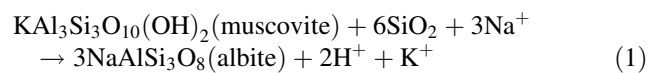
The mesoscale characteristics of the red-bed soft rocks can be described by physical characteristics of the rocks (e.g. pore structure), while the damage behaviour of the rocks (e.g. crack propagation) represents the macroscale

characteristics of the rocks. For the microscale characteristics of the red-bed soft rocks, the ion compositions in and pH of water resulting from water–rock interaction were investigated. The chemical behaviour at the microscale due to water–rock interaction could lead to the change in mechanical properties at macroscale and ultimately the failure of the rocks [56, 58, 64]. Therefore, it can be said that the chemical effect is the dominant effect at the microscale of red-bed soft rocks.

In conclusion, the microscale, mesoscale and macroscale catastrophic behaviour of red-bed soft rocks can be characterised by chemical effects, physical effects and mechanical effects, respectively.

As shown in Fig. 1, the present theoretical study on changes in the microscale characteristics of red-bed soft rocks under water saturation mainly focussed on the change in ion concentration and pH of water. Figure 4 shows that, at the water–rock interface, clay minerals, which are often negatively charged, attract cations in the aqueous solution, resulting in the formation of a hydration shell. The chemical reactions involving ion-exchange adsorption, dissolution and formation of soluble minerals lead to the production of water-soluble substances [38], resulting in the loss of element structure of the clay mineral layer.

Based on the previous studies and on time-dependent change in ion concentration in water during water–silty mudstone interaction under normal temperature and pressure conditions [5, 18, 33, 40, 44, 51, 52, 59], Considering the simplified calculation and the high concentration of  $\text{Na}^+$  and  $\text{Ca}^{2+}$  ions in the solution, feldspar only considers the transformation of albite and potash feldspar. And Fig. 5 and Eqs. (1)–(6) describe the chemical transformation of muscovite, feldspar, montmorillonite, kaolinite, chlorite and illite involving  $\text{Na}^+$ ,  $\text{K}^+$  and  $\text{Ca}^{2+}$  ions in the solution and synthesised albite, potash feldspar and anorthite.



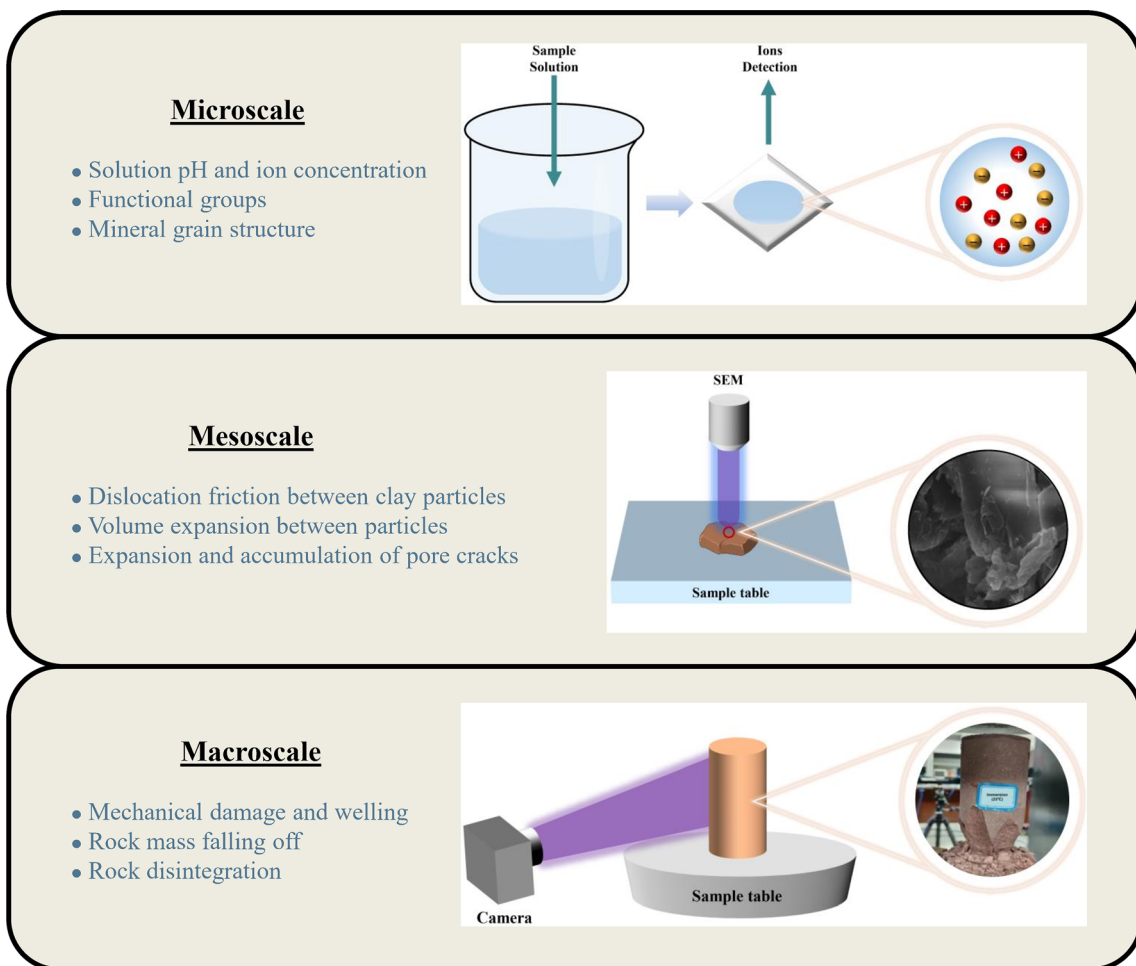
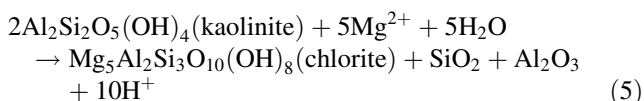


Fig. 3 The scale of red-bed soft rocks



Chemical reaction involves energy absorption and release, which can be described by the change in enthalpy ( $\Delta H$ ). There is a heat absorption of red-bed soft rocks from water when  $\Delta H > 0$ , while the release of heat from the rocks into water happens when  $\Delta H < 0$ . Thus,  $\Delta H$  can be used to characterise the microscale dissolution and formation of rock minerals. The microscale energy ( $Q_{\text{micro}}$ ) can be divided into adsorption energy ( $Q_1$ ) and chemical energy ( $Q_2$ ) [70]. That is,

$$Q_{\text{micro}} = Q_1 + Q_2 = n \cdot E_a + \sum_{i=1}^n c_i \cdot V_{r/w} \cdot \Delta_r H_{m_i}^0 \quad (7)$$

where  $n$  is the molar amount of adsorbed ions in solution,  $E_a$  is the unit molar adsorption energy,  $c_i$  is the molar concentration of  $i$ -th reaction,  $V_{r/w}$  is the volume of

aqueous solution, and  $\Delta_r H_{m_i}^0$  is the standard molar enthalpy change of the  $i$ -th reaction.

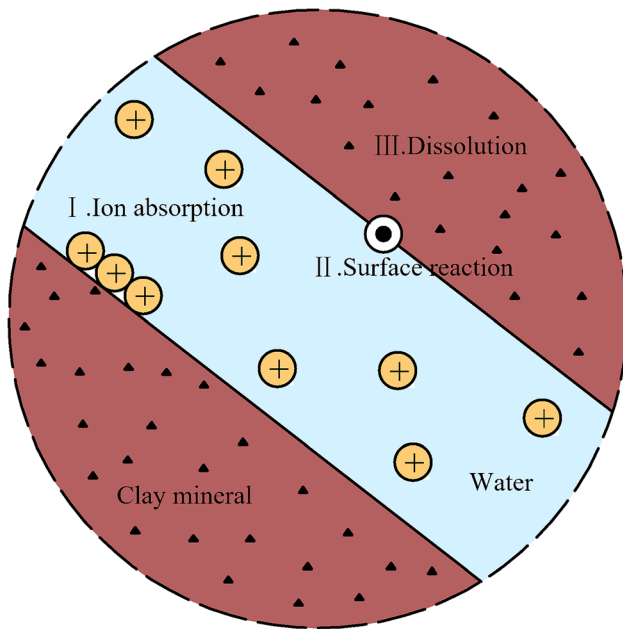
### 2.2.2 Modelling change in mesoscale characteristics of red-bed soft rocks under water action

During the water–rock interaction process, the chemical behaviour of the red-bed soft rocks at the microscale results in pore structure at the mesoscale involving change in potential energy, swelling energy and other types of energy [70]. It is assumed that potential energy and swelling energy are 2 major types of energy at the mesoscale.

The potential energy caused by the increase of specific surface area due to the accumulation of microcracks and pores can be expressed as [32]

$$E_b = G_{1c} \cdot A_{\text{surface}} \quad (8)$$

where  $E_b$  is the change in energy in rock surface due to softening and disintegration of red-bed soft rocks,  $G_{1c}$  is the energy per unit rock surface, and  $A_{\text{surface}}$  is the newly



**Fig. 4** Dissolution and diffusion diagram of clay minerals at water–rock interface

increased surface area of red-bed soft rocks. The swelling energy of mesoscale clay minerals can be expressed as [45]

$$Q_v = - \int_{V_1}^{V_2} P dV \tag{9}$$

where  $Q_v$  is the swelling energy,  $V_1$  and  $V_2$  are the volumes before and after swelling, respectively,  $P$  is the swelling stress, and  $V$  is the volume of red-bed soft rocks.

### 2.2.3 Modelling change in macroscale characteristics of red-bed soft rocks under water action

In this study, the study of macroscale energy of red-bed soft rocks mainly focuses on strain energy and frictional energy. Strain energy can be divided into elastic potential energy and plastic potential energy.

The elastic potential energy can be expressed as [60]

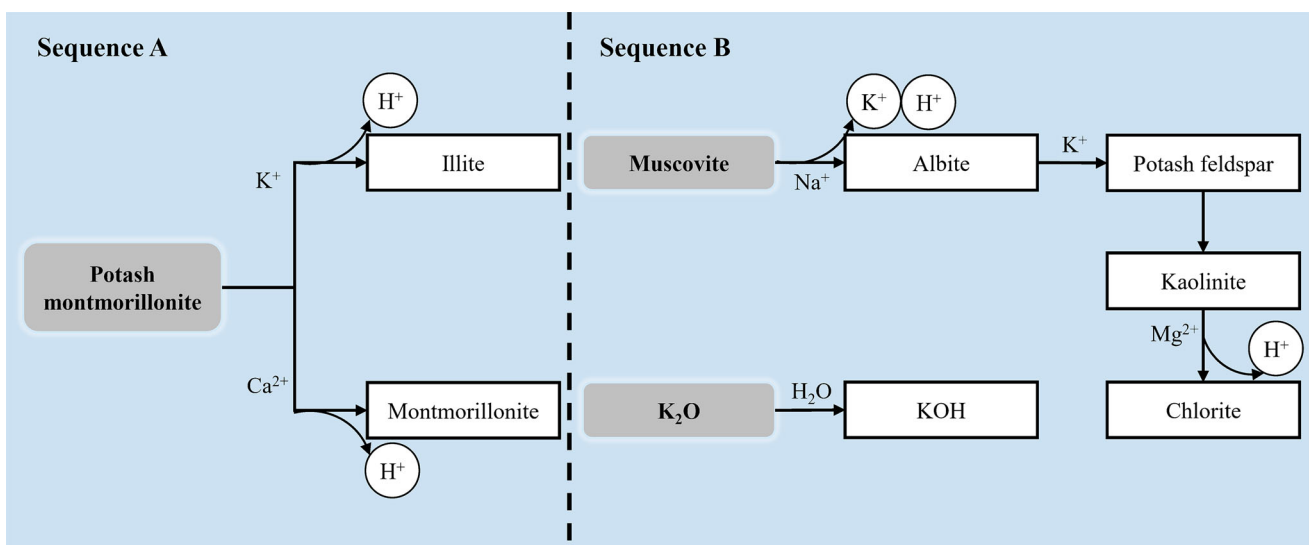
$$W_e = [\sigma_1^2 + \sigma_2^2 + \sigma_3^2 - 2\mu(\sigma_1\sigma_2 + \sigma_1\sigma_3 + \sigma_2\sigma_3)] / (2E) \cdot V_v \tag{10}$$

where  $W_e$  is the elastic property of red-bed soft rocks with a volume of  $V_v$ ,  $\mu$  is the Poisson’s ratio of red-bed soft rocks, and  $E$  is the elastic modulus of red-bed soft rocks.

The plastic potential energy, which consists of crack-breaking energy and void-breaking energy, can be described as [12]

$$W_p = W_1 + W_2 = W_f + 2U_e + W_v = 2c\tau_f \frac{(k+1)(1+\mu)}{2E \sin \theta} \cdot \left[ \frac{2c\tau^* \sin \theta}{\sqrt{w \sin \frac{\pi(l+l^*)}{w}}} - \sigma_2 \sqrt{2w \tan \frac{\pi l}{2w}} + \sigma_2 \sqrt{\frac{\pi l}{2}} \right] \cdot \sqrt{2\pi(l+l^{**})} + \frac{(k+1)(1+\mu)}{E} \cdot \left[ \frac{4c^2\tau^{*2} \sin^2 \theta}{\pi} \ln \left[ \frac{\tan \frac{\pi(l+l^*)}{2w}}{\tan \frac{\pi l}{2w}} \right] - 4\sigma_2^2 \frac{w^2}{\pi} \ln \left( \cos \frac{\pi l}{2w} \right) - 8c\tau^* \sigma_2 \sin \theta \frac{w}{\pi} \cdot \ln \left[ \tan \frac{\pi}{4} \left( 1 + \frac{l}{w} \right) \right] \right] + V_{\text{matrix}} \cdot \sigma_{\text{eq}} \cdot \epsilon_{\text{eq}}^p \tag{11}$$

where  $W_1$  is the fracture energy caused by cracking,  $W_2$  is the fracture energy induced by the development of voids,



**Fig. 5** Schematic diagram showing chemical reactions at water–rock interface

$W_f$  is the energy dissipated due to initial slippage of cracks,  $U_e$  is the energy dissipated due to tensile crack propagation,  $W_v$  is the fracture energy of void development, and the pore volume is  $V_{\text{matrix}}$ . The initial crack length is  $2c$  with a tensile crack length of  $2l$  ( $l^* \approx 0.27c$  and  $l^{**} = 0.083c$ ). The distance between adjacent cracks is  $2w$ , and the angle between the initial crack surface and the axial stress is  $\theta$ .  $\tau_f$  is the shear force produced by the external load sliding on the initial crack surface ( $\tau_f = \frac{1}{2}\mu[(\sigma_1 + \sigma_2) - (\sigma_1 - \sigma_2)\cos 2\theta]$ ).  $\tau^*$  is the shear force produced by the external load sliding on the tensile crack surface

$$(\tau^* = \frac{1}{2}(\sigma_1 - \sigma_2)\sin 2\theta - \frac{1}{2}v[(\sigma_1 + \sigma_2) - (\sigma_1 - \sigma_2)\cos 2\theta]).$$

$v$  is the friction coefficient between crack surfaces,  $E$  is the elastic modulus, and  $\mu$  is the Poisson’s ratio of the material.  $k = (3 - \mu)/(1 + \mu)$  is a constant in plane stress state.  $\sigma_{\text{eq}}$  is the equivalent stress, and  $\varepsilon_{\text{eq}}^p$  is the equivalent plastic strain. Since there is little change in elastic energy, it is reasonable to assume that  $\varepsilon_{\text{eq}}^p = \varepsilon_{\text{eq}}$ .

The macroscale frictional energy of red-bed soft rocks consists of the heat generated by the friction of structural plane and that generated by friction of particles. The adiation energy, kinetic energy and other types of energy are ignored in this study [16, 29, 48, 62, 63]. The frictional energy of the structural plane can be expressed as [2, 3, 30]

$$U^s = U^f + U^c + U^n + U^h \tag{12}$$

where  $U^s$  is work done by shear stress.  $U^f = \sum^n \tau_f A l_i$  is work done by friction, where  $A$  is the frictional surface area and  $l_i$  is the frictional distance.  $U^c$  is work done by complete cutting on convex surface and is ignored in this study by assuming a flat structure plan without concave convex fluctuation, in order to simplify the calculation.  $U^n$  is work done by normal load. For a flat surface, there is no difference in height between the undulating surface and the frictional surface (i.e.  $h = 0$ ). Thus,  $U^n = \sigma_n, h = 0$ .  $U^h$  is the work done by resistance between red-bed soft rocks and fluid. Since  $V_{\text{fluid}} = 0$ ,  $U^h$  can also be ignored.

The frictional energy of red-bed soft rocks particles ( $U_p$ ) can be expressed as [65]

$$U_p = \frac{F_s \cdot F_n \cdot \Delta t^2}{m} \tag{13}$$

where  $F_s = \tau_f$  is the frictional force of collision friction between particles and  $F_n = mg$  is the average impact force on the indirect contact surface of particles.  $\Delta t$  is the time of particle collision and friction (0.100 s), and  $m$  is the particle mass.

### 2.2.4 Modelling micro-meso-macroscale energy conservation of red-bed soft rocks under water action

The overall water–rock interaction can be modelled based on conservation of energy. According to the above analysis, the different types of energy caused by the crack propagation and change in pore structure can be expressed as

$$Q_{\text{meso}} + Q_{\text{macro}} = E_b + Q_v + W_e + W_p + E_s + U_s + U_p \tag{14}$$

where  $Q_{\text{meso}}$  is the change of energy due to mesoscale physical effects and  $Q_{\text{macro}}$  is the change of energy due to macroscale mechanical effects. Finally, the relationship between microscale adsorption energy and chemical energy is

$$Q_{\text{total}} = Q_{\text{micro}} + Q_{\text{meso}} + Q_{\text{macro}} = Q_1 + Q_2 + E_b + Q_v + W_e + W_p + U_s + U_p \tag{15}$$

where  $Q_1$  is the adsorption energy,  $Q_2$  is the chemical energy,  $E_b$  is the potential energy caused by the increase of surface area resulting from red-bed soft rocks softening and disintegrating,  $Q_v$  is the swelling energy of red-bed soft rocks,  $W_e$  is the elastic potential energy of red-bed soft rocks,  $W_p$  is the plastic potential energy of red-bed soft rocks,  $U_s$  is the structural plane frictional energy of red-bed soft rocks, and  $U_p$  is the frictional energy of mudstone particles.

At the microscale, ion adsorption and chemical reaction occurring at the water–rock interface lead to the swelling of fracture pores at the mesoscale and, ultimately, energy change resulting from the increase of specific surface area. Therefore, there is no overlap between microscale energy and mesoscale energy. For macroscale energy, the strain energy in the stress–strain curve overlaps with the potential energy in the mesoscale. That is,

$$Q_{\text{overlap}} = E_b \cap (W_e + W_p) = W_e + W_p \tag{16}$$

Thus, the energy conservation under water–rock interaction can be expressed as

$$Q_{\text{conservation}} = Q_{\text{micro}} + Q_{\text{meso}} + Q_{\text{macro}} - Q_{\text{overlap}} + Q_{\text{water}} = Q_1 + Q_2 + E_b + Q_v + W_e + W_p + U_s + U_p = \Delta C \tag{17}$$

where  $Q_{\text{conservation}}$  is the energy conservation under water–rock interaction, and  $Q_{\text{overlap}}$  is the overlapping energy at the mesoscale and macroscale.  $\Delta C$  is a constant ( $\Delta C = 0$  if there is no energy exchange with the outside).



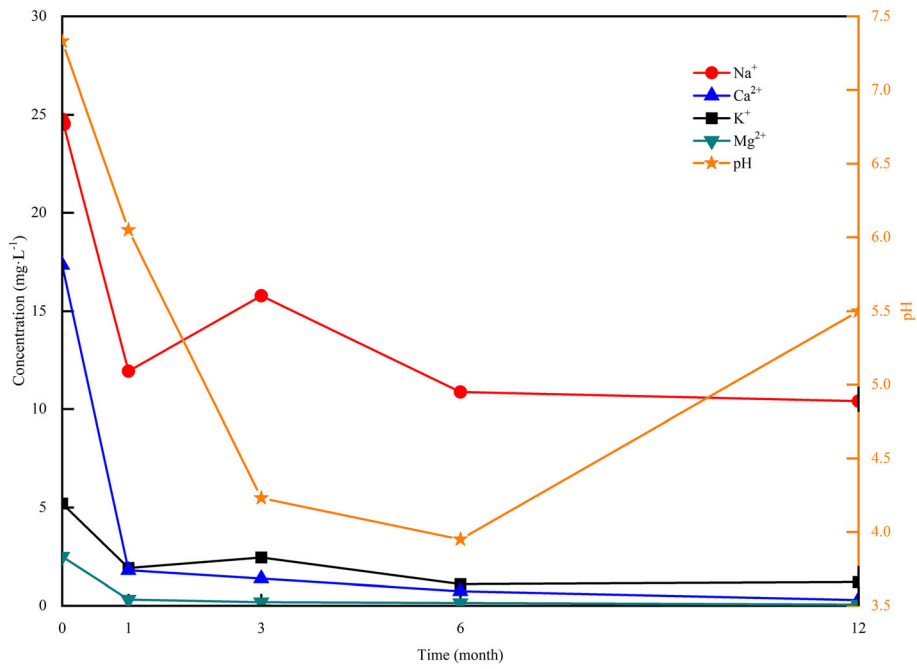


Fig. 6 Time-dependent cation concentration and pH in aqueous solution

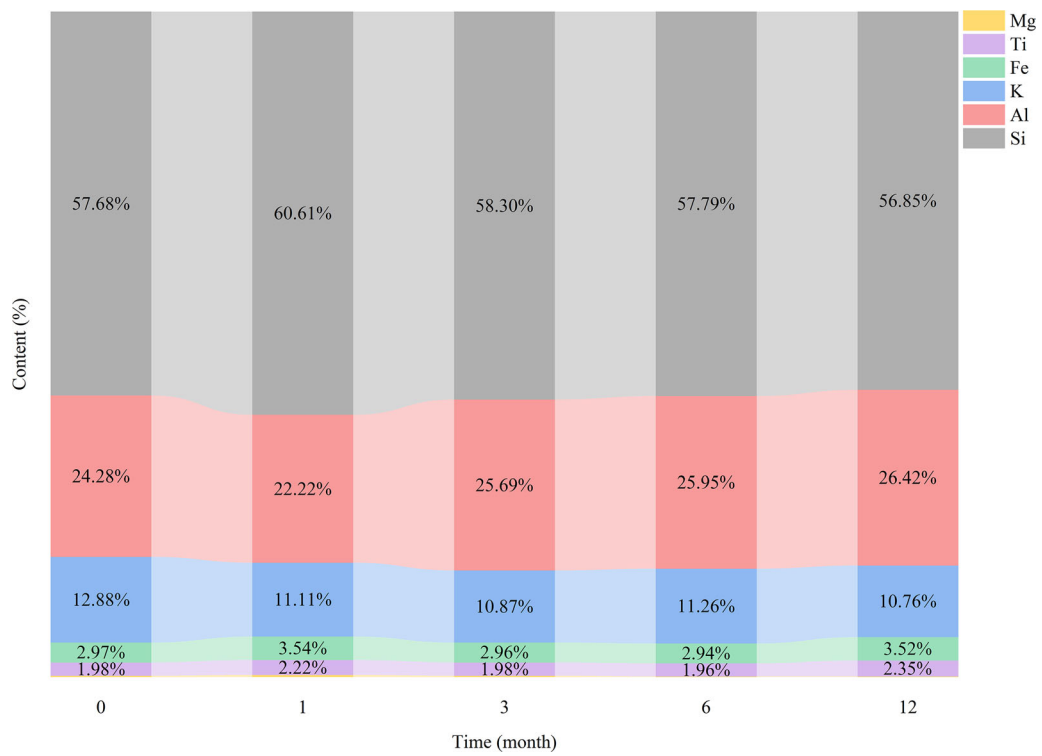
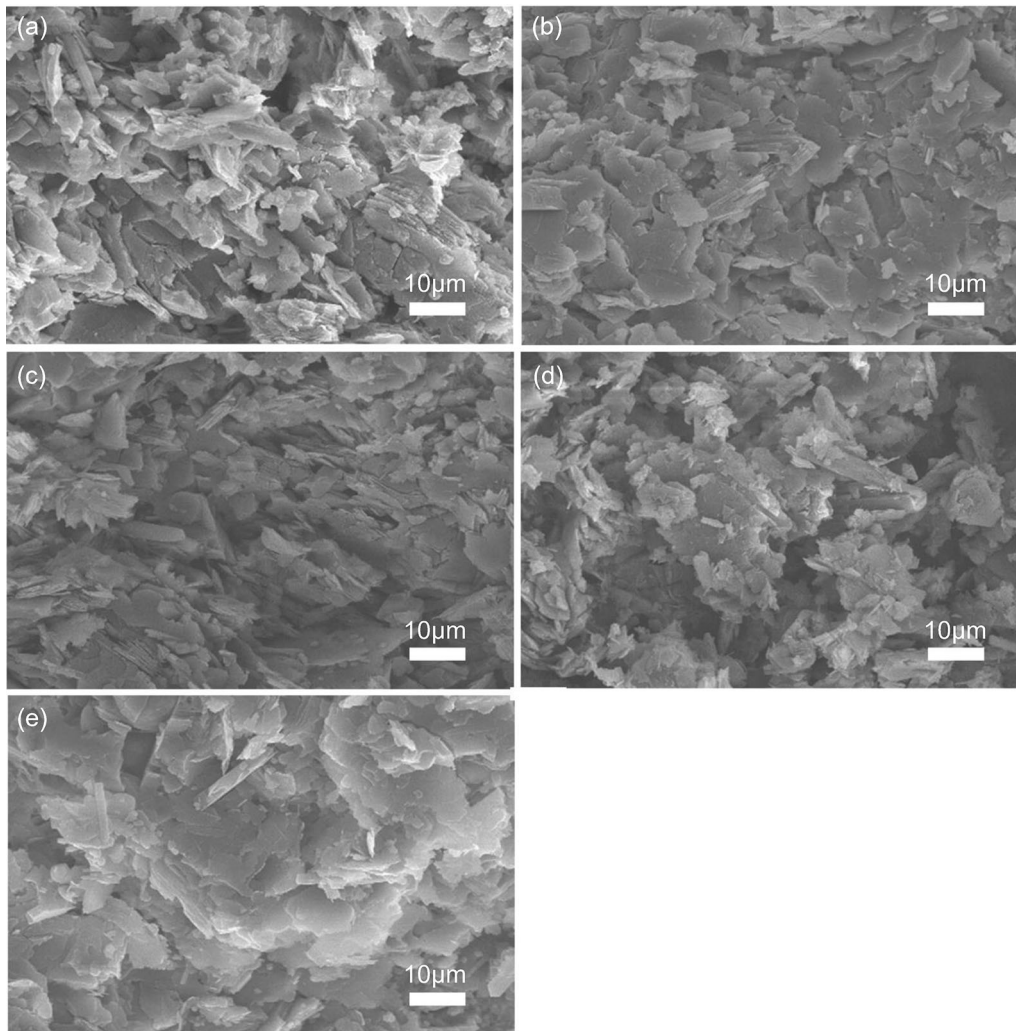


Fig. 7 Time-dependent change in chemical composition of red-bed soft rocks samples during water saturation



**Fig. 8** SEM image of microstructure of red-bed soft rocks samples over a magnification range from  $3,000\times$  down to  $1\times$ . **a** Before water saturation; **b** after 1 month of water saturation; **c** after 3 months of water saturation; **d** after 6 months of water saturation; **e** after 12 months of water saturation

### 3 Results and discussion

#### 3.1 Time-dependent change in microscale chemical composition and pore structure in soft rock samples

Figure 6 shows that the concentrations of  $\text{Na}^+$  and  $\text{K}^+$  decreased exponentially with time and that the decrease was more significant during the first month after water saturation began, whereas there was little change in concentrations of  $\text{Ca}^+$  and  $\text{Mg}^{2+}$ . Initially, concentrations of  $\text{Na}^+$  and  $\text{K}^+$  in aqueous solution were higher than that within the rock. This resulted in the migration of  $\text{Na}^+$  and  $\text{K}^+$  into the rock due to chemical potential, leading to the increase in concentration of  $\text{H}^+$  in aqueous solution (i.e. decrease in pH). The rate of decrease in concentration of  $\text{H}^+$  then gradually reduces until the dynamic balance of

exchange of metal ions between aqueous solution and rock is reached. The further increase in pH in aqueous solution is due to the formation of the electric double layer at the water–rock interface [26]. Figure 7 shows that there is little change in chemical composition in the red-bed soft rocks during water saturation process because a relatively small proportion of the rocks were involved in the chemical reaction.

In the process of water saturation of red-bed soft rocks, the elements of aqueous solution change and, accordingly, the element contents of rock sample should also change. It can be seen from Fig. 7 that the element contents of the rock sample do have some changes, but they are very small relative to the change of metal ion concentration in the aqueous solution. This is because the proportion of the reaction part in the total content of the soft rocks is

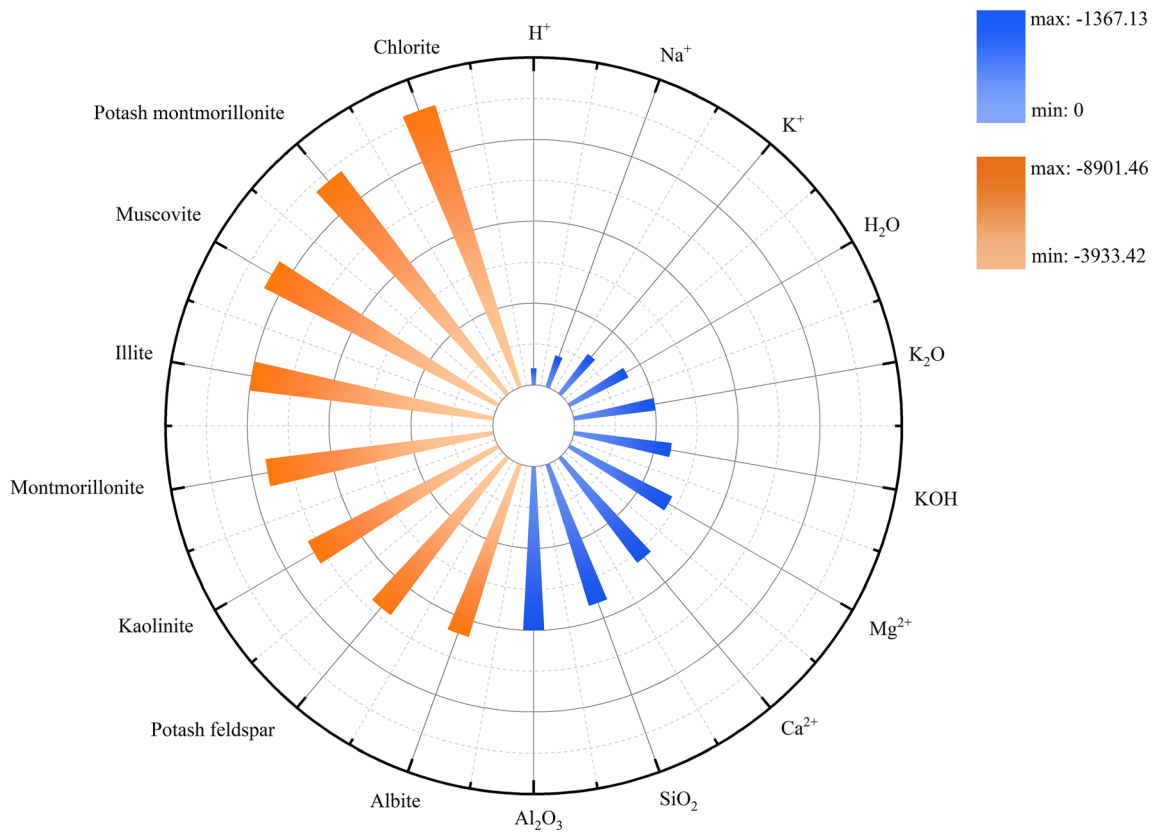


Fig. 9 The standard molar reaction enthalpy of reactants and resultant  $\Delta_r H_m^0$

Table 2 Standard molar enthalpy change of related chemical reactions

Reaction sequence	Standard molar enthalpy of formation
	$\Delta_r H_m^0$ (kJ/mol)
Muscovite → albite	109.075
Potash feldspar + potash montmorillonite → kaolinite + illite	-11,332.900
Potash montmorillonite → montmorillonite	1179.656
Kaolinite → chlorite	497.540
K <sub>2</sub> O → KOH	-193.004
Kaolinite → illite	68.170

relatively small. For the red-bed soft rocks as a whole, the percentage of element change is small.

Figure 8 shows the time-dependent change in microstructure of red-bed soft rocks before and after water saturation, quantified using SEM imaging. First, the porosity of the rock was characterised by processing SEM images through binary and the fractal box dimension. It shows that the pore structure of the rocks remained unchanged after 1 month of water saturation, although

there was a certain increase in porosity and pore expansion. After 3 months of water saturation, the connection of small pores resulted in the formation of large pores, and the shape of pores gradually changed from strip and irregular polygon to oblong and oval. After 6 months of water saturation, the total number of pores in the rock decreased, but the total area of pores increased. The shape of pores became more rounded, and the pore structure became looser. Obvious pore connection was seen after 12 months of water saturation, with significant change in the morphology of the particles due to chemical reaction. The pore shape changes flocculent and petal.

At the microscale, a series of chemical reactions took place between the minerals within the rocks and water entering fractures on the rock surface through seepage. The formation of water-soluble substances results in the loss of components of the clay mineral layer and, ultimately, changes in chemical energy. While the chemical reactions between rock minerals and aqueous solution could accelerate the damage of fracture pores, the development of fractures in rock could in return promote the chemical effects. Thus, there is a close interrelation between micro-meso-macroscale chemical, physical and mechanical effects on red-bed soft rocks.

**Table 3** Micro-meso-macroscale energy in red-bed soft rock samples

Rock scale	Category	Magnitude
Microscale	Adsorption energy	1.195 kJ
	Chemical energy	−44.384 kJ
	Other energy	0
Mesoscale	Potential energy	−6.390 kJ
	Swelling energy	−0.033 kJ
	Other energy	0
Macroscale	Strain energy	0.134 kJ
	Frictional energy	−0.840 kJ
	Other energy	0

### 3.2 Micro-meso-macroscale energy conservation during water–rock interaction

It is known that  $V_{\text{water}} = 1.250 \times 10^{-2} \text{ m}^3$ ,  $m_{\text{sample}} = 1.795 \text{ kg}$ ,  $E_{\text{sample}} = 1960 \text{ MPa}$  and Poisson’s ratio of red-bed soft rocks  $\mu = 0.280$  [66, 67]. The process of ion adsorption can be divided into surface adsorption and chemical adsorption, with an average surface adsorption of 16.732 kJ/mol and chemical adsorption of 83.720 kJ/mol [35]. Figure 9 shows the standard molar reaction enthalpy of reactants and resultant  $\Delta_r H_m^0$  [42].

Basing on the thermodynamic parameters of reaction-related minerals and relevant thermodynamic methods [7, 40, 42], the standard molar enthalpy of formation  $\Delta_r H_m^0$  can be obtained by using the data of standard molar enthalpy of reactants and resultants ( $\Delta_r H_m^0 = \Delta H_{\text{resultant}} - \Delta H_{\text{reactant}}$ ) as shown in Table 2.

Based on the coefficients for chemical equations and the change of ion concentration before and after the process of water saturation, a polynomial equation was established with the parameters  $a, b, c, d, e$  and  $f$ , which were  $0.208 \text{ mol/m}^3$ ,  $0.193 \text{ mol/m}^3$ ,  $0.142 \text{ mol/m}^3$ ,  $0.020 \text{ mol/m}^3$ ,  $0.219 \text{ mol/m}^3$ ,  $0.375 \text{ mol/m}^3$ , respectively. Therefore, the total amount of adsorbed ions in the solution ( $n$ ) was 0.162 mol. Substituting the values of parameters into Eq. (7) gives

$$\begin{aligned} Q_{\text{micro}} &= Q_1 + Q_2 = n \cdot E_a + \sum_{i=1}^n c_i \cdot V_{r/w} \cdot \Delta_r H_{m_i}^0 \\ &= n \cdot E_a + c_1 \cdot V_{\text{water}} \cdot \Delta_r H_{m_1}^0 \\ &\quad + c_2 \cdot V_{\text{water}} \cdot \Delta_r H_{m_2}^0 + c_3 \cdot V_{\text{water}} \cdot \Delta_r H_{m_3}^0 \\ &\quad + c_4 \cdot V_{\text{water}} \cdot \Delta_r H_{m_4}^0 \\ &\quad + c_5 \cdot V_{\text{sample}} \cdot \Delta_r H_{m_5}^0 + c_6 \cdot V_{\text{water}} \\ &\quad \cdot \Delta_r H_{m_6}^0 = -23.161 \text{ kJ} \end{aligned}$$

Assuming specific heat capacity  $c_{\text{rock}}$  of silty mudstone is  $0.4375 \text{ kJ/(kg}\cdot\text{°C)}$  [11], the temperature change of red-

bed soft rocks due to the exothermic reactions at micro-scale can be described as  $\Delta T_{\text{rock}} = Q_{\text{micro}}/(c_{\text{rock}} \cdot m_{\text{sample}}) = 54.756 \text{ °C}$ . The above is the temperature change of soft rock mass. For the temperature change of water, it can be described as  $\Delta T_{\text{water}} = Q_{\text{micro}}/(c_{\text{water}} \cdot m_{\text{water}}) = 0.441 \text{ °C}$ . It is consistent with the results of almost no change in the temperature of aqueous solution measured by the thermometer before and after the test. The thermal stresses  $\sigma_{\text{ft}} = E_1 \cdot \varepsilon = E_1 \cdot \alpha \cdot \Delta T_{\text{rock}} = 643,930.600 \text{ Pa}$ , where  $E_1$  is the elastic modulus of red-bed soft rocks,  $\alpha$  is the coefficient of thermal expansion,  $\alpha = 6 \times 10^{-6}/\text{°C}$ ,  $\varepsilon$  is the thermal expansion strain, and  $\Delta T$  is the temperature difference of red-bed soft rocks.

Previous studies [9] showed that the initial crack length ( $2c$ ) of red-bed soft rocks is correlated with the particle diameter of red-bed soft rocks. The average particle size of silty mudstone is about 1.500 mm [68], and the distance between adjacent cracks is about 6 mm [19]. A friction coefficient of crack surface  $\nu = 3$ , angle between the initial crack and the axial stress  $\theta = 45^\circ$ ,  $l = \sqrt{2c} = 1.225 \text{ mm}$ ,  $P = \sigma_1 = \sigma_2 = \sigma_{\text{ft}} = 643,930.600 \text{ Pa}$  and  $\sigma_3 = \sigma_{\text{static}} = 2450 \text{ Pa}$  gives

$$F_s = \tau_f = \frac{1}{2} \mu [(\sigma_1 + \sigma_2) - (\sigma_1 - \sigma_2) \cos 2\theta] = 90.493 \text{ kPa};$$

$$\begin{aligned} k &= (3 - \mu)/(1 + \mu) = 2.125, \quad l^{**} = 0.083c \\ &= 0.06225 \text{ mm}; \end{aligned}$$

$$\begin{aligned} \tau^* &= \frac{1}{2} (\sigma_1 - \sigma_2) \sin 2\theta - \frac{1}{2} \nu [(\sigma_1 + \sigma_2) - (\sigma_1 - \sigma_2) \cos 2\theta] \\ &= 223.783 \text{ kPa}; \end{aligned}$$

$$\begin{aligned} \sigma_{\text{eq}} &= \frac{\sqrt{2}}{2} \sqrt{(\sigma_1 - \sigma_2)^2 + (\sigma_2 - \sigma_3)^2 + (\sigma_3 - \sigma_1)^2} \\ &= 643.931 \text{ kPa}; \end{aligned}$$

$$\begin{aligned} \varepsilon_{\text{eq}}^p &= \varepsilon_{\text{eq}} = \frac{\sqrt{2}}{3} \sqrt{(\varepsilon_1^p - \varepsilon_2^p)^2 + (\varepsilon_2^p - \varepsilon_3^p)^2 + (\varepsilon_3^p - \varepsilon_1^p)^2} \\ &= 2.622 \times 10^{-5} \end{aligned}$$

The energy change per unit surface can be assumed as  $3.6 \text{ kJ/m}^2$  [10, 43]. At the mesolevel, the rock is composed of many particles of various sizes and cements filled between particles. To simplify the problem, the natural morphology of red-bed soft rocks is replaced by  $n$  spheres with an average particle size of 1.500 mm [34, 39, 41]. Granting to the principle of volume reciprocity, the number of particles ( $n$ ) can be assumed to be 479,900–480,100 and the average value is 480 000. In addition, based on the porosity and volume changes, it can be assumed that  $A_{\text{surface}} = n \cdot A_{\text{particle}} \cdot e = 1.775 \text{ m}^2$  [67], where  $A_{\text{particle}}$  is the surface area of particles,  $e$  is the sum of porosity and volume change rate. Further,  $U^s = U^f + U^n$ , where  $U^f =$

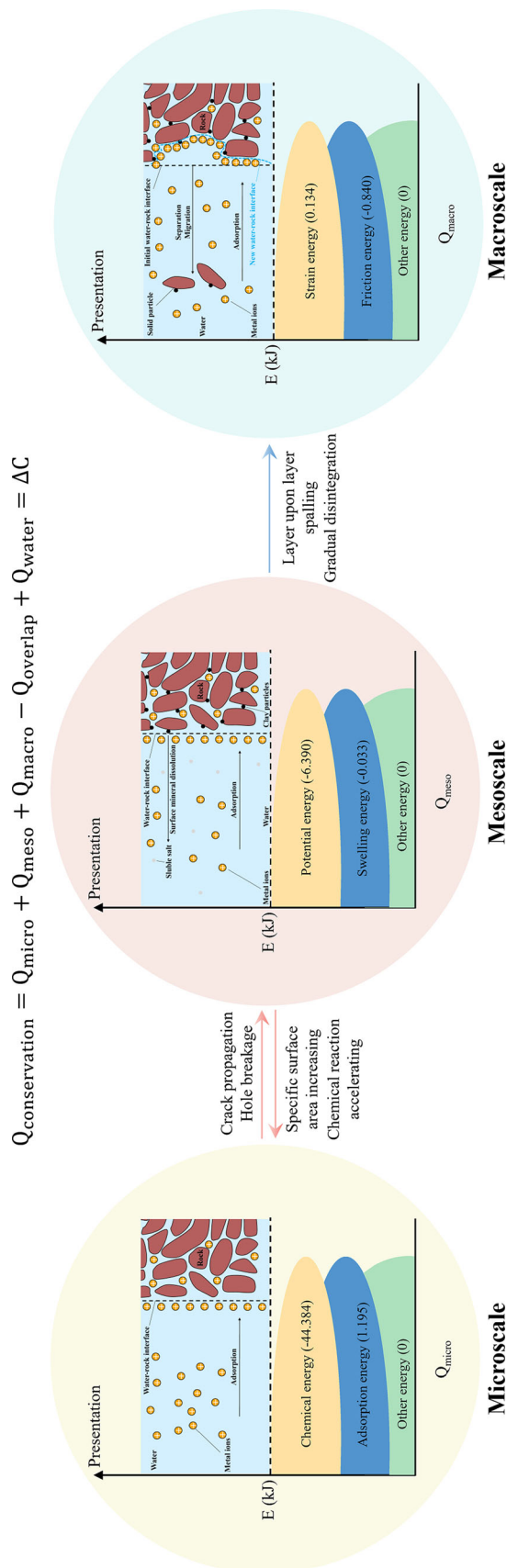


Fig. 10 Micro-meso-macroscale energy conservation during water–soft rocks interaction

$\sum_i^n \tau_f A l_i$  and  $A$  is the frictional surface area. To simplify the calculation,  $A = 0.008 \text{ m}^2$  is calculated as twice the horizontal area of the sample, and the frictional distance  $l_i = 0.050 \text{ m}$ . The average impact force on the indirect contact surface of particles is  $F_n = m_{\text{particles}} \cdot g$ , where  $m_{\text{particles}}$  is the mass of particles. The time of particle collision and friction ( $\Delta t$ ) is calculated as  $0.100 \text{ s}$ . Thus, the micro-meso-macroscale energy can be calculated using Eqs. (7)–(15), and the results are shown in Table 3.

The adsorption of ions and chemical reactions at the microscale at the water–rock interface could lead to the increase of swelling in fracture pores at the mesoscale and an increase in change of energy in particular regions of the rock surface. While there is no overlap in microscale and mesoscale energy, the strain energy in the stress–strain curve overlaps with the potential energy at the mesoscale. Figure 10 shows the conservation of energy during water–rock interaction.

We can analyse and summarise the micro-meso-macrocorrelation mechanism of red-bed soft rocks failure within static water by energy. As shown in Figs. 10 and 11, the surface of the red-bed soft rocks specimen has a large number of microfractures and microporosity, which provides space for the movement of aqueous solution seepage. In the initial stage of soft rock encountering water, the aqueous solution enters the initial fracture and initial pore space on the surface of soft rock through seepage. At this time, the integrity of the soft rock is basically intact, and there is an obvious initial water–rock interface between the soft rock and the aqueous solution. The microscopic effect is dominated by the migration and adsorption of cations from aqueous solutions to the water–rock interface, and almost no chemical effect of clay mineral dissolution occurs. It is mainly accompanied by the adsorption energy of ions with the water absorption and swelling energy of clay minerals, etc.

As the concentration of aggregated cations within the water–rock interface becomes higher, their chemical reaction with the soft rock results in the loss of the original microstructure of the clay mineral layer. This process generates soluble salts released into aqueous solution and generates chemical thermal. The mesoscale is reflected by the crack-pore damage and the increase of surface energy. In turn, the mesoscale behaviour promotes the seepage and microchemical effects of the aqueous solution. When the clay particles continue to dissolve and fall off, it eventually leads to weakening and separation of the bond between the skeletal particles in the soft rock. This results in the formation of a new water–rock interface. At the new interface, the aqueous solution continues the above process of

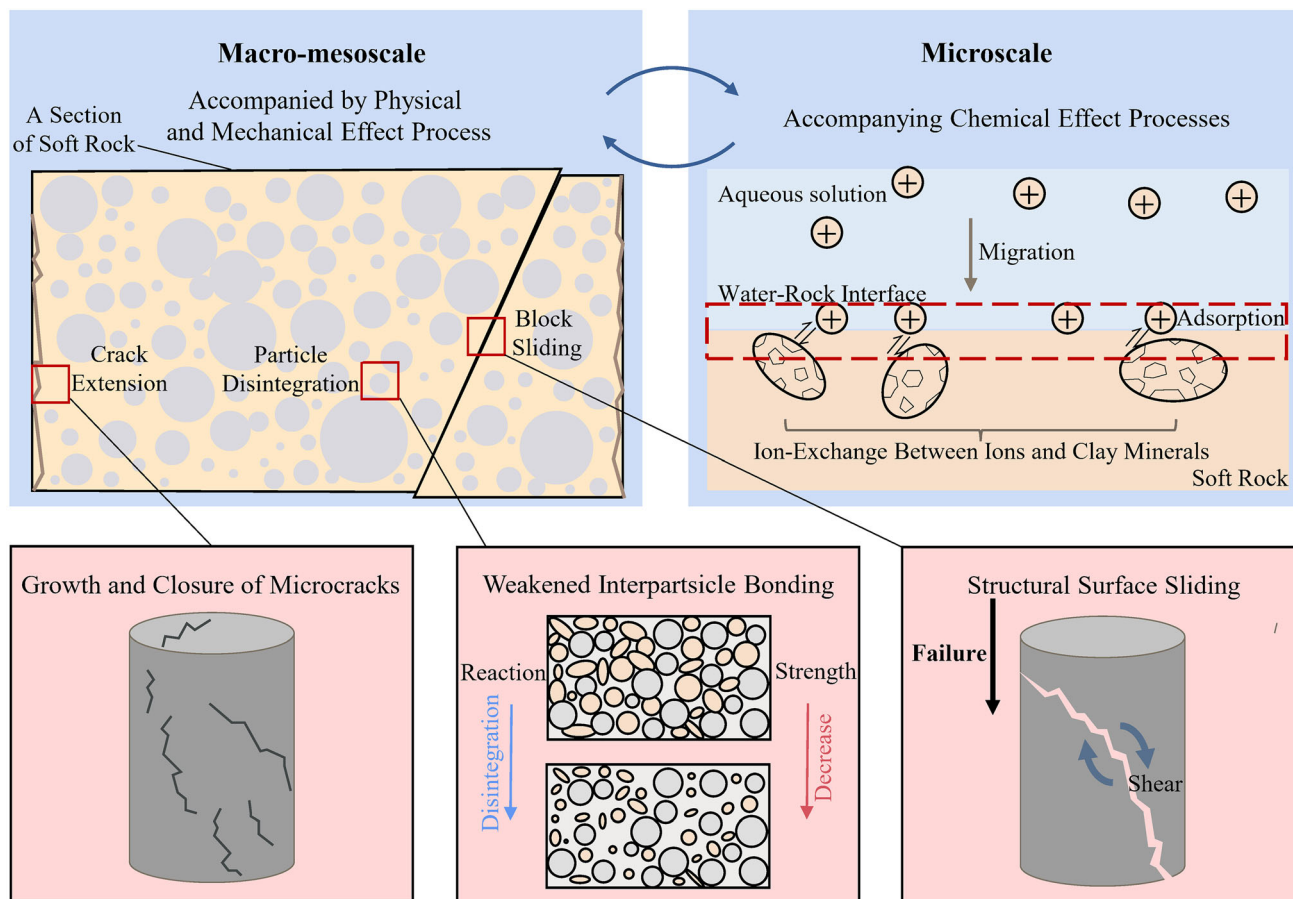


Fig. 11 The micro-meso-macrocorrelation mechanism of red-bed soft rocks failure within static water

seepage, dissolution and shedding of clay particles, leading to the formation of more new water–rock interfaces.

The cascade of the previous processes ultimately causes the disintegration and dislodgement of the fragments of the macroscale soft rock masses by increasing loosening. This macroscopic process is accompanied by the strain energy and the friction energy until the entire soft rock specimen is completely destroyed. The temperature in aqueous solution increases when the overall red-bed soft rocks reaction is exothermic.

As shown in Fig. 12, during the water–rock interaction, nearly 80% of the energy dissipation results from the microscale change, and meso-macroscale change accounts for the remaining 20%. Therefore, the chemical effect of the interaction between red-bed soft rocks and water is microscopic and long term. The impact on red-bed soft rock is, furthermore, the most fundamental. Thus, the chemical effects at microscale dominate the softening of red-bed soft rocks. Around 90% of the energy is generated from the chemical reaction between ions and clay minerals at the microscale, while the remaining 10% comes from ion adsorption. It can be mentioned that microscale water–rock interaction redistributes chemical elements between soft

rock and water. Chemical action is accompanied by the formation of new minerals, which will demolish the initial internal structure of the rock. At the mesoscale, the dominant potential energy (98%) leads to the damage of pores and an increase in surface area of red-bed soft rocks. As mentioned above, the change of rock microstructure under the action of water physics and chemistry constitutes the fundamental reason for the change of rock macromechanical properties. Almost all of the energy is accumulated in the red-bed soft rocks, but some of the energy releases into the aqueous solution. As shown in Fig. 13, the disintegration of the mesoparticles could appear when the energy accumulation of the mesoparticles reaches the critical point resulting in disintegration of the particles. While the damage of soft rocks fragments at macroscale results in energy dissipation, the remaining energy could be stored in the surrounding undamaged red-bed soft rocks.

Figure 14 shows that, in the early stages of water–rock interaction, water moves into rock fractures through seepage, while metal ions accumulate at the water–rock interface, resulting in ion adsorption and chemical reaction at the interface. This leads to the formation of soluble salt and mineral colloid that sheds the clay particles in the red-bed

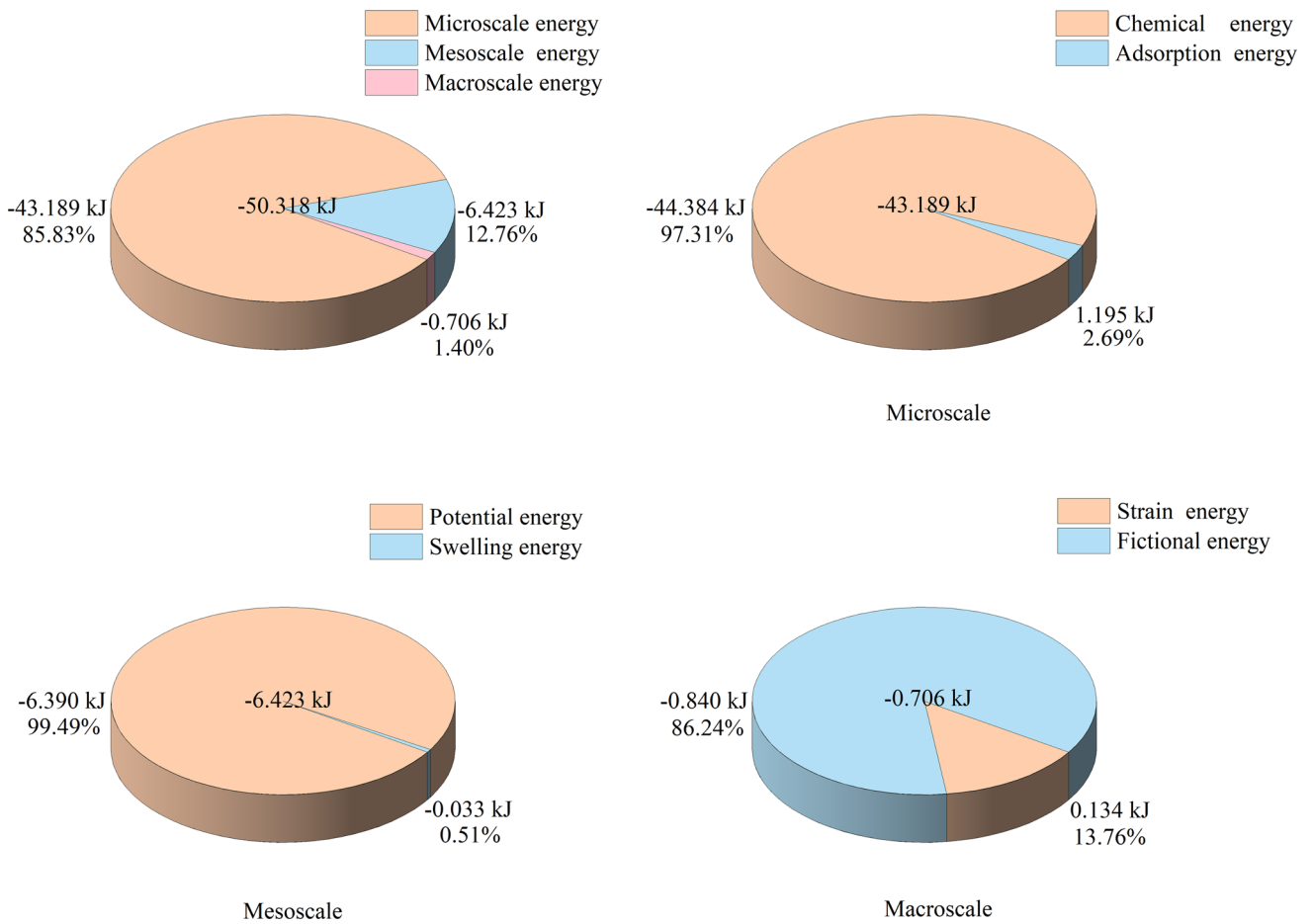


Fig. 12 Distribution of micro-meso-macroscale energy of red-bed soft rocks samples after 12 months of water–rock interaction

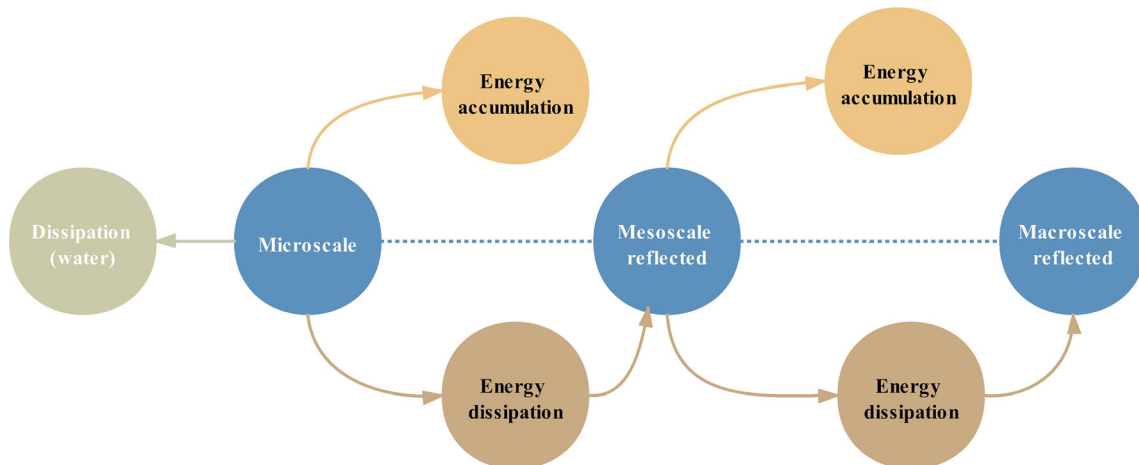
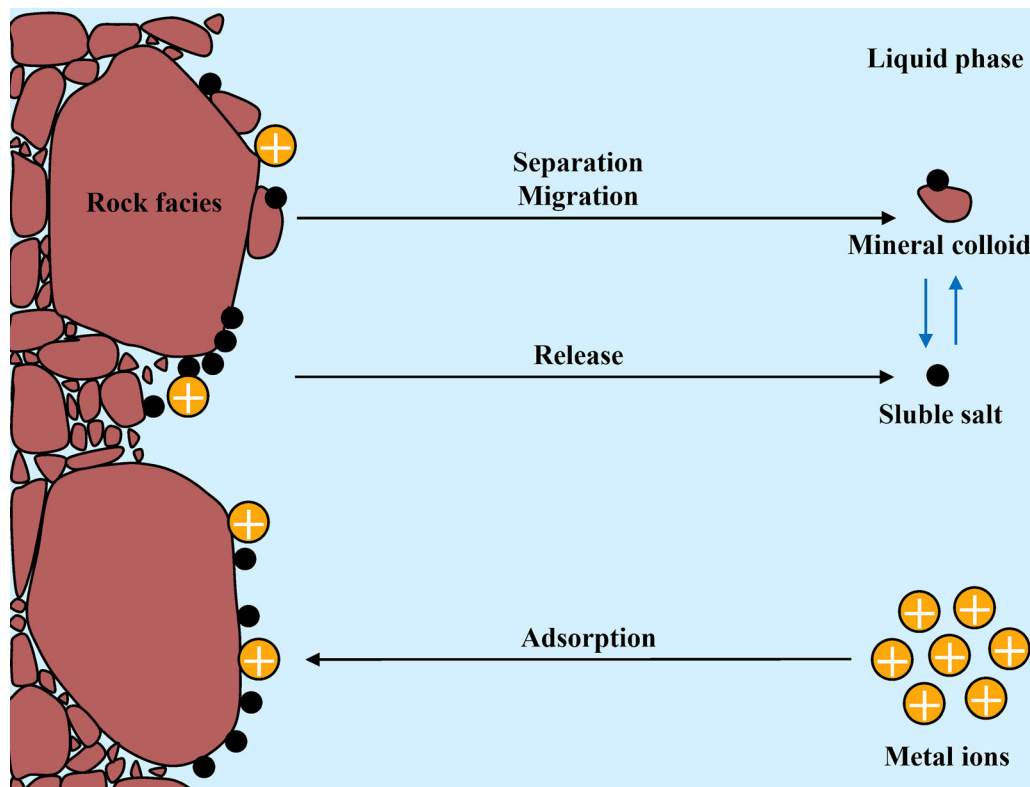


Fig. 13 Energy dissipation and accumulation at micro-meso-macroscale of red-bed soft rocks samples during water–rock interaction

soft rocks from releasing into water. Red-bed soft rocks gradually disintegrate until it was destroyed finally.

### 4 Conclusion

In this study, a series of experimental studies were carried out to investigate the micro-meso-macroscale correlation of red-bed soft rocks during the water–rock interaction.



**Fig. 14** Micro-meso-macroscale failure mechanism of red-bed soft rock samples during water–rock interaction

The multiscale energy conservation was also theoretically studied to understand the failure mechanism of the rocks. The following are some major conclusions:

- The microscale dominant effect of red-bed soft rocks is determined as chemical effects.
- The failure of red-bed soft rocks under water–rock interaction results from chemical, physical and mechanical dominant effects at the micro-meso-macroscale.
- The concentrations of  $\text{Na}^+$  and  $\text{K}^+$  in water decrease exponentially with time, and the decrease is more significant during the first month of the water saturation begins, whereas there is little change in concentrations of  $\text{Ca}^+$  and  $\text{Mg}^{2+}$  in water.
- The transport of  $\text{Na}^+$  and  $\text{K}^+$  from water into rock leads to a decrease in pH in aqueous solution.
- SEM images show that change in microstructure of the red-bed soft rocks becomes obvious after 3 months of water saturation. After 6 months of water saturation, the total number of pores in the rock decreases, but the total area of pores increases.
- While there is no overlap in microscale and mesoscale energy, the strain energy in the stress–strain curve overlaps with the potential energy at the mesoscale.

- During the water–rock interaction, nearly 80% of the energy dissipation results from the microscale change, and meso-macroscale change accounts for the remaining 20%. Thus, the chemical effects at microscale dominate the softening of red-bed soft rocks.
- Around 90% of the energy generates from the chemical reaction between ions and clay minerals at microscale, while the remaining 10% is from ion adsorption.
- As a uniform measure of material evolution, energy can correlate the micro-meso-macrodominant effect of red-bed soft rock softening appropriately, which provides a new idea for revealing the catastrophic mechanism of red-bed soft rock.

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### Declarations

**Conflict of interest** The authors can confirm there are no conflicts of interest associated with this work.

**Data availability statement** The data used to support the findings of this study are available from the corresponding author upon request.



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