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Diurnal variation in mesophyll conductance and its influence on modelled water-use efficiency in a mature boreal *Pinus sylvestris* stand

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Abstract

Mesophyll conductance (g_m) is a critical variable for the use of stable carbon isotopes to infer photosynthetic water-use efficiency (WUE). Although g_m is similar in magnitude to stomatal conductance (g_s) , it has been measured less often, especially under field conditions and at high temporal resolution. We mounted an isotopic CO₂ analyser on a field photosynthetic gas exchange system to make continuous online measurements of gas exchange and photosynthetic ¹³C discrimination (Δ^{13} C) on mature *Pinus sylvestris* trees. This allowed the calculation of g_m , g_s , net photosynthesis (A_{net}), and WUE. These measurements highlighted the asynchronous diurnal behaviour of g_m and g_s . While g_s declined from around 10:00, A_{net} declined first after 12:00, and g_m remained near its maximum until 16:00. We suggest that high g_m played a role in supporting an extended A_{net} peak despite stomatal closure. Comparing three models to estimate WUE from Δ^{13} C, we found that a simple model, assuming constant net fractionation during carboxylation (27%e), predicted WUE well, but only for about 75% of the day. A more comprehensive model, accounting explicitly for g_m and the effects of daytime respiration and photorespiration, gave reliable estimates of WUE, even in the early morning hours when WUE was more variable. Considering constant, finite g_m or g_m/g_s yielded similar WUE estimates on the diurnal scale, while assuming infinite g_m led to overestimation of WUE. These results highlight the potential of high-resolution g_m measurements to improve modelling of A_{net} and WUE and demonstrate that such g_m data can be acquired, even under field conditions.

Keywords Carbon isotope ratio \cdot Cavity ring-down absorption spectrometry \cdot Online discrimination \cdot Vapour pressure deficit \cdot Photosynthesis

Abbreviations

A _{net}	Net photosynthesis rate (μ mol CO ₂ m ⁻² s ⁻¹)
$\Delta_{\rm e}$	Discrimination associated with respiration (%)
$\Delta_{\rm f}$	Discrimination associated with photorespiration
	(‰)

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Δ_{i}	Discrimination when $C_i = C_c$ (%)
Δ_0	Observed photosynthetic discrimination (%)
a _m	¹² C/ ¹³ C fractionation during dissolution and dif-
	fusion through water 1.8 (%)
$a_{\rm s}$	¹² C/ ¹³ C fractionation during diffusion through air
5	4.4 (‰)
b	¹² C/ ¹³ C fractionation during carboxylation 29
	(‰)
\bar{b}	¹² C/ ¹³ C net fractionation during carboxylation 27
	(‰)
$C_{\rm c}$	CO_2 concentration in the chloroplast
	$(\mu mol mol^{-1})$
C_{i}	CO_2 concentration in the intercellular space
	$(\mu mol mol^{-1})$
$C_{\rm in}$	CO_2 concentration in the air entering the cuvette
	$(\mu mol mol^{-1})$

C _{out}	CO_2 concentration in the cuvette (µmol mol ⁻¹)
Ε	Transpiration rate (mmol $H_2O m^{-2} s^{-1}$)
е	Fractionation during respiration (%)
f	Fractionation during photorespiration 16.2 (%)
$g_{ m m}$	Mesophyll conductance (mol $CO_2 m^{-2} s^{-1} bar^{-1}$)
$g_{\rm s}$	Stomatal conductance (mol $CO_2 m^{-2} s^{-1} bar^1$)
$R_{\rm d}$	Day time respiration rate (μ mol CO ₂ m ⁻² s ⁻¹)
$R_{\rm p}$	Photorespiration rate (μ mol CO ₂ m ⁻² s ⁻¹)
t	Ternary correction factor (–)
VPD	Vapour pressure deficit (kPa)
WUE	Water-use efficiency (mmol $CO_2 \text{ mol}^{-1} H_2O$)
WUE _G	WUE from gas exchange (mmol $CO_2 \text{ mol}^{-1}$
-	H ₂ O)
WUE_{Λ}	WUE calculated from Δ^{13} C (mmol CO ₂ mol ⁻¹
	H ₂ O)

 Γ^* CO₂ compensation point (µmol mol⁻¹)

Introduction

Mesophyll conductance (g_m) describes the ability of CO₂ to diffuse across the interior of the leaf. In plants with C3 photo synthesis, g_m is roughly similar in magnitude to stomatal conductance (g_s) , frequently accounting for about 40% of the decline in CO₂ concentration from the ambient atmosphere to the chloroplasts (C_c) (Flexas et al. 2008; Warren 2008a). As a consequence, it has an important place in leaf-level photosynthesis models (von Caemmerer 2000; Dewar et al. 2017), but has been so infrequently quantified that it is seldom included in earth-system models (Rogers et al. 2017). It also has a critical role in the inference of water-use efficiency (WUE) from stable carbon isotope composition (δ^{13} C) of plant tissues or, conversely, in the inference of $\delta^{13}C$ from gas exchange (Rogers et al. 2017). This role is caused by the decrease in CO₂ concentration at the enzyme rubisco, where δ^{13} C is determined, relative to the substomatal cavities, where WUE is determined. Mesophyll conductance provides a means to calculate this difference. If g_m could be accounted for, then δ^{13} C could provide independent tests of the WUE predictions of leaf (von Caemmerer 2000; Wei et al. 2014), canopy (Keenan et al. 2013), and earth-system models (Rogers et al. 2017).

One reason for the relative paucity of g_m data is that it is more difficult to estimate than g_s . Stomatal conductance to CO₂ is easily estimated from humidity, temperature and transpiration measurements, which might come from leaflevel gas exchange or sap-flux data. Given the relative ease of making such measurements, high temporal resolution g_s data are available for many species and sites, and models of g_s based on theory and empirical data have converged (Medlyn et al. 2011) and been incorporated into global models (Prentice et al. 2014; Rogers et al. 2017). In contrast, measuring g_m requires simultaneous measurements of gas exchange and either chlorophyll fluorescence or photosynthetic discrimination against ¹³C (Δ^{13} C). Discrimination can be inferred from the δ^{13} C signature of photosynthesis products, e.g. leaf soluble sugars, phloem contents, plant tissues (e.g.: Hu et al. 2010; Ubierna and Marshall 2011), or directly from leaf CO₂ flux (e.g.: Evans et al. 1986; Warren et al. 2003; Bickford et al. 2009; Wingate et al. 2007, 2010; Maseyk et al. 2011; Campany et al. 2016). These methods are technically challenging, especially under field conditions, so that measurements are often made with low temporal resolution.

It has been difficult to model $g_{\rm m}$, because previous studies have found that g_m and g_s respond differently to changes in environmental conditions, suggesting that the two are not tightly coupled. Rapid responses of g_m have been described to several environmental variables (for a review see Flexas et al. 2008; Warren 2008a; Flexas et al. 2012). These variables include light intensity or quality (Flexas et al. 2007; Tholen et al. 2008; Hassiotou et al. 2009; Loreto et al. 2009; Campany et al. 2016), intercellular CO_2 concentration (C_i) (Flexas et al. 2007; Hassiotou et al. 2009; Vrábl et al. 2009; Bunce 2010; Douthe et al. 2011; Tazoe et al. 2011), and leaf temperature (Bernacchi et al. 2002; Yamori et al. 2006; Warren 2008b; Evans and von Caemmerer 2013). If $g_{\rm s}$ responded to other variables, or at different rates, then the ratio g_m/g_s would change. For example, it has been shown that $g_{\rm m}$ responds similarly, but more quickly, to variable $C_{\rm i}$ than g_s (Flexas et al. 2007). In addition, the g_m/g_s ratio was found to be temperature dependent in a study exploring the thermal acclimation of $g_{\rm m}$ in spinach (Yamori et al. 2006).

Vapour pressure deficit (VPD) is particularly interesting in this context, because g_s responds so strongly to it (Marshall and Waring 1984; Oren et al. 1999; Medlyn et al. 2011). In contrast, the response of g_m to VPD has not been extensively studied and the results so far are contradictory (Bongi and Loreto 1989; Warren 2008c, Loucos et al. 2017). Both temperature and VPD change dynamically under natural conditions, diurnally and seasonally, potentially influencing the g_s to g_m relationship. However, the magnitude and importance of this variability has yet to be explored.

Given constant g_s , an increase in g_m would increase wateruse efficiency (WUE) (Flexas et al. 2010; Galmés et al. 2011), which is defined as the ratio of net carbon assimilation (A_{net}) to water loss through transpiration (E). This would happen because an increase in g_m has no direct effect on transpiration, but it increases photosynthesis, resulting in an increase of the A_{net}/E ratio. Accounting for g_m is especially important when estimating WUE from δ^{13} C. For example, WUE is often inferred from historic tree-ring isotope data (Marshall and Monserud 1996; Seibt et al. 2008; Voelker et al. 2016). Such inferences require that some value for g_m be assumed. This assumption is often embedded as a constant, empirical adjustment in the relationship between C_i/C_a and isotopic discrimination (Farquhar et al. 1982), or extrapolated based on its correlation with g_s in models of WUE (Klein et al. 2015), although as noted above, the correlation with g_s is not always strong.

In this manuscript, we present continuous, simultaneous measurements of shoot-scale gas exchange and ¹³C discrimination in a 100-year-old Pinus sylvestris stand in northern Sweden. We use these simultaneous data streams to obtain hourly $g_{\rm m}$ estimates parallel to $g_{\rm s}$, $A_{\rm net}$ and E. We begin with a brief description of how the data were treated and evaluate the accuracy of our measurement system. We next explore the diurnal dynamics of g_s and g_m and their relationship to $A_{\rm net}$. Finally, we compare estimates of WUE derived from gas exchange (WUE_G) with estimates derived from photosynthetic discrimination (WUE $_{\Lambda}$). Three photosynthetic discrimination (Δ^{13} C) models were used to calculate WUE_{Λ}: a comprehensive model, a partial model and a simple model. Additionally, the comprehensive model was applied using three different assumptions for g_m values. We compare the different models and calculations and discuss their impact on WUE_{Λ} estimates.

Materials and methods

Description of the experimental site

The study was conducted in a~100-year-old, naturally regenerated, even-aged stand of Pinus sylvestris (Scots pine) at the Rosinedalsheden experimental forest in northern Sweden (64°10'N, 19°45'E, 153 m above see level), during the growing season of 2017. The Rosinedalsheden experiment includes an intensive fertilisation treatment (Lim et al. 2015), but the current study was conducted entirely on the unfertilised area. The photosynthetic season typically extends from mid-April to mid-November, buds burst at the end of May, and stem diameter-growth ceases in late August (Tarvainen et al. 2018). The June to August mean temperature was 12.4 ± 0.8 °C (mean \pm SD) and the mean monthly precipitation was 67.9 ± 8.6 mm (mean \pm SD), based on the 15-year (2003-2017) data measured at the Vindeln-Sunnansjönäs meteorological station (Swedish Meteorological and Hydrological Institute, www.smhi.se) approximately 5 km from the experimental site. The site has weakly podzolised fine sandy soil with a thin (2–5 cm) organic layer (Hasselquist et al. 2012). The leaf area index was 2.7 and the average tree height was 18.6 ± 2.3 m (mean \pm SD) in 2013 (Lim et al. 2015).

Experimental setup for continuous measurements of gas exchange

Shoot gas exchange (CO₂ and H₂O) was measured continuously on one 1-year-old upper canopy shoot on four trees. A 16-m tall scaffolding tower was used to reach the shoots and secure the equipment. The shoot-scale gas exchange was measured using a custom-built multichannel gas exchange system (GUS) (Wallin et al. 2001; Tarvainen et al. 2016), equipped with infrared gas analysers (IRGA, CIRAS-1, PP systems Hitchin Herts, U.K.) to measure CO₂ and H₂O partial pressure in the air from shoot cuvettes and reference channels. The 330 ml shoot cuvettes had a transparent acrylic plastic (Plexiglas) top for natural illumination. The cuvettes were temperature (T) controlled to track the ambient T and were equipped with a light sensor (PAR-1 M, PP systems, Hitchin, Herts, UK). The polyethylene tubing that connected the cuvettes to the IRGAs were insulated and heated with cables to avoid condensation. Nonetheless, morning condensation could occur in the cuvettes in connection with heavy rain events; we filtered those days out in the current analysis. The GUS cycled through the four shoot cuvettes and two non-cuvette lines once per hour, spending 7 min at each position, which were divided into 2 min of waiting time to allow instrument readings to stabilise and 5 min of measurement. We used the means from the 5-min measurement periods in the subsequent analyses, which yielded approximately one value/cuvette/hour throughout the 9 days. The non-cuvette lines were used for data quality assurance and for measurement of δ^{13} C of ambient air (see details in next chapter). The IRGAs were calibrated with 400 μ mol mol⁻¹ CO₂ gas at the beginning and at the end of the growing season. Additionally, every hour the IRGAs were zero calibrated and the system ran a cross-calibration protocol to match values in the sample and reference channels.

Continuous measurement of $\delta^{13}C$

The isotopic composition of the CO_2 in the air entering and leaving the cuvettes was analysed with a cavity ring-down spectrophotometer (CRDS; G2131-i, Picarro Inc., California, USA). The CRDS was connected to the same central line as the GUS, in parallel to the sample IRGA. We tested the instrument at varying CO₂ and H₂O vapour concentrations and found that the δ^{13} C values were dependent on both, with an asymptotic relationship of δ^{13} C to CO₂ concentration (Fig. S1) and a linear dependency to H_2O vapour concentration (Fig. S2). The continuous $\delta^{13}C$ readings were corrected to account for the CO₂ and H₂O concentration effects before the data were used in further analyses. The CRDS was factory-calibrated in 2017 and manually calibrated once per week, using two reference gases with known CO₂ concentrations (411 μ mol mol⁻¹, SD = 5.1; 1606 μ mol mol⁻¹, SD = 13.1) and δ^{13} C values $(-32.36\%_0, SD = 0.09; -4.14\%_0, SD = 0.06)$. The reference gases were analysed at the SLU Stable Isotope Laboratory (Umeå, Sweden) with GB-IRMS (Gasbench II—Isotope Ratio Mass Spectrometer, Thermo Fisher Scientific, Bremen, Germany), which was calibrated against IAEA-co-9 and NBS 19 standards. We found the weekly calibrations to be sufficient, because the reference δ^{13} C values were stable over the season (Fig. S3) and were offset from the reference gases by a constant 4.17% (SD = 0.1), after correction for CO₂ concentration. The CRDS recorded δ^{13} C values once per second during the 5-min calibration period, which were then combined into a mean for each calibration date and these means yielded standard deviations of 0.1% for δ^{13} C.

Calculation of leaf gas exchange parameters and mesophyll conductance

In this paper, we present data collected on nine sunny days during the summer (28th of June—2nd of July and 6th of July—9th of July), with daily minimum and maximum temperatures of 6.2 ± 0.5 °C and 24.4 ± 0.6 °C, respectively, and daily maximum irradiation of $1964 \pm 25 \ \mu mol \ m^{-2} \ s^{-1}$. Because of the high latitude and season, sunrise was typically around 02:15 and sunset was around 23:00. These days were chosen for high photosynthetic rates and lack of condensation in the cuvettes and tubing. We optimised the system setup to yield clear and consistent δ^{13} C values with the CRDS, using 5-min integrations at approximately 1-second intervals. Because each of these measurements contributed to the mean δ^{13} C value, it was appropriate to calculate the standard error of the mean from them. This yielded high precision, typically SE < 0.06‰.

 A_{net} , *E*, *g*_s, and *C*_i were calculated from the gas exchange data according to the model described by Farquhar et al. (1980). Boundary layer conductance has previously been found to be high (8.1 mol H₂O m⁻² s⁻¹) (Uddling and Wallin 2012) in our gas exchange cuvettes, therefore, we assumed boundary layer resistance to be insignificant. Needles from the shoots enclosed in the cuvettes were collected at the end of the study campaign to determine the projected leaf area using a flat bed scanner (Epson 1600) equipped for dual scanning, and WinSEEDLE Pro 5.1a (Regent Instruments, Canada) analysis software.

Mesophyll conductance (g_m) and C_c were estimated from the carbon isotope discrimination data collected by the CRDS. The g_m was calculated from the comprehensive Δ^{13} C model of Farquhar and Cernusak (2012) that includes ternary corrections. In particular, we used the formulation of Evans and von Caemmerer (2013) (see supplementary materials for details) that calculates g_m as

$$g_{\rm m} = \frac{\frac{1+t}{1-t} \left(b - a_{\rm m} - \frac{eR_{\rm d}}{A_{\rm net} + R_{\rm d}} \right) \frac{A_{\rm net}}{C_{\rm out}}}{\Delta_{\rm i} - \Delta_{\rm o} - \Delta_{\rm e} - \Delta_{\rm f}}$$
(1)

where *b*, $a_{\rm m}$ and *e* are the fractionation factors during carboxylation ($b = 29\%_{c}$), dissolution and diffusion through water ($a_{\rm m} = 1.8\%_{c}$) and respiration (*e*, see Eqn. S5), respectively. $R_{\rm d}$ is daytime respiration (Eqn. S1), and $C_{\rm out}^{-1}$ is the CO₂ concentration in the cuvette; $\Delta_{\rm i}$, $\Delta_{\rm o}$, $\Delta_{\rm e}$ and $\Delta_{\rm f}$ are, respectively, the discrimination when $C_{\rm i} = C_c$ (Eqn. S2), the observed discrimination during gas exchange (Eqn. S3 and S4), the discrimination associated with respiration (Eqn. S5 and S6) and with photorespiration (Eqn. S7). The term *t* is the ternary correction factor (Eqn. S8, Farquhar and Cernusak 2012). Note that $C_{\rm out}$ is lower than the atmospheric CO₂ concentration ($C_{\rm in}$), due to $A_{\rm net}$ within the cuvette. The CO₂ concentration at the site of carboxylation (C_c) was calculated from $g_{\rm m}$ through the following relationship:

$$A_{\rm net} = g_{\rm m} \left(C_{\rm i} - C_{\rm c} \right) \tag{2}$$

We evaluated how the magnitude of the net photosynthetic CO₂ drawdown, calculated as $C_{in}-C_{out}$, affected our estimates of C_c . This drawdown, together with instrument precision, determines the error associated with Δ^{13} C measurements, which ultimately determines the error in C_c and $g_{\rm m}$ estimates (for a discussion see Ubierna et al. 2018). The concentration drop is evaluated with the parameter $\zeta = C_{in}/\zeta$ $(C_{in}-C_{out})$. Pons et al. (2009) showed that the error associated with $g_{\rm m}$ estimates increased when ζ was large and the instrument precision was low. We likewise found that C_c became exponentially more variable as the CO₂ drawdown in the cuvette decreased below 20 μ mol CO₂ mol⁻¹ (Fig. 1a). Assuming an ambient CO₂ concentration of 400 ppm, a drawdown of 20 μ mol CO₂ mol⁻¹ corresponds to $\zeta = 20$ (=400/(400-380)). In this case, and with an instrument precision of 0.06% the error associated with Δ^{13} C measurements was 1.7% (= $\sqrt{2} \cdot \zeta$ · Precision). This low drawdown and large error mainly occurred during early mornings and evenings, so that it was not possible to acquire reasonable estimates before 04:00 and after 20:00 (Fig. 1b). For most of the day, the observed drawdown was substantially greater than 20 μ mol CO₂ mol⁻¹ which resulted in $\zeta < 10$ and associated errors in $\Delta_0 < 0.8\%$ (Fig. 1b).

Comparison of models to estimate water-use efficiency from $\Delta^{13}C$

Instantaneous WUE can be derived from $\Delta^{13}C$ (WUE_{Δ}) or from gas exchange (WUE_G) measurements as (Seibt et al. 2008; Hu et al. 2010; Wang et al. 2014; Klein et al. 2015; Guerrieri et al. 2016):

¹ In calculations based on on-line discrimination, we consider the CO₂ concentration and the isotopic composition inside the cuvette (noted with the subscript 'out') as 'ambient' because they represent the immediate environment of the shoot or leaf. If the models presented here (Eqs. 1, 3, 4, 5) are to be applied elsewhere, C_{out} can be replaced with C_a and ${}^{13}C_{\text{out}}$ can be replaced with ${}^{13}C_a$.



Fig. 1 a Standard deviation of the CO₂ concentration in the chloroplast (C_c) in relation to CO₂ drawdown in the cuvette, defined as the difference between the atmospheric concentration (C_{in}), and the concentration inside the cuvette (C_{out}). The SD of C_c was estimated for every 2 µmol mol⁻¹ change in the CO₂ drawdown. The figure is based

on non-filtered data. **b** The diurnal time course of the CO_2 drawdown. The whiskers of the box-plots extend to 1.5 times the interquartile range. The grey area marks the time of day excluded from the analysis. The figure is based on non-filtered data

$$WUE_{t} = \begin{cases} a) \text{ from } \Delta^{13}C \to WUE_{\Delta} = \frac{(C_{out} - C_{i}) * VPD}{1.6} \\ b) \text{ from gas exchange } \to WUE_{G} = \frac{A_{net}}{E} \end{cases}$$
(3)

where C_i is solved from a theoretical Δ^{13} C model.

We considered three models for Δ^{13} C, which resulted in three estimations of WUE_{Δ}. First we estimated C_i from the simple model by Farquhar et al. (1982), as

$$C_{\rm i} = C_{\rm out} * \frac{\Delta_{\rm o} - a_{\rm s}}{\bar{b} - a_{\rm s}} \tag{4}$$

where \bar{b} is taken as 27‰, a standard value for C₃ plants, that was derived empirically from relationships between δ^{13} C of leaf bulk material and C_i/C_a values (Farquhar et al. 1982; Cernusak et al.2013; Ubierna and Farquhar 2014). This model does not account specifically for the dependency of Δ^{13} C on g_m , R_d or photorespiration (R_p); instead it includes these effects empirically within \bar{b} , which is often sufficient in practice (Cernusak et al. 2013; Bloomfield et al. 2019). Second, we estimated C_i from a model proposed by Seibt et al. (2008), subsequently referred to as the partial model, as

$$C_{\rm i} = C_{\rm out} * \frac{\Delta_{\rm o} - a_{\rm s} + (b - a_{\rm m}) \frac{g_{\rm s}}{1.6 * g_{\rm m}} + f \frac{\Gamma^*}{C_{\rm out}}}{b - a_{\rm s} + (b - a_{\rm m}) \frac{g_{\rm s}}{1.6 * g_{\rm m}}}$$
(5)

where Γ^* is the CO₂ compensation point, derived from an Arrhenius function (Bernacchi et al. 2001; Medlyn et al. 2002 Eq. 12). This model accounts explicitly for the effect of $g_{\rm m}$ and $R_{\rm p}$ and assumes negligible effect of $R_{\rm d}$. Finally, from the comprehensive model of Farquhar and Cernusak (2012) $C_{\rm i}$ can be solved as (see supplemental materials in Cernusak et al. 2018)

$$C_{\rm i} = \frac{-\mathrm{II} \pm \sqrt{\mathrm{II}^2 - 4\mathrm{I} * \mathrm{III}}}{2\mathrm{I}} \tag{6}$$

where equations I, II and III are given in the supplementary materials (Eqn S9–S11). This model accounts explicitly for $g_{\rm m}$, $R_{\rm p}$ and $R_{\rm d}$, and it includes a correction for the ternary effect.

To avoid circularity, we divided our data set into two parts. We used 4 days' data to estimate a mean value of g_m (0.29 mol CO₂ m⁻² s⁻¹ bar⁻¹). Subsequently, this mean g_m was used on the remaining 5 days' data to estimate C_i from Δ^{13} C (Eqs. 4, 5 and 6) and WUE (Eq. 3). Note that the 4-day mean g_m was slightly different from the mean for all 9 days together, which was 0.31 mol CO₂ m⁻² s⁻¹ bar⁻¹. For the comprehensive model, besides using a constant mean g_m , we also performed the calculations using either a constant g_m/g_s (2.9) or infinite g_m and compared these estimates as well to WUE_G.

Data analysis

A statistical filter was applied to the dataset to discard outliers in C_i-C_c and g_m . Any data point outside the range of mean ± 3 SD was considered an outlier and removed. This filter removed 4.5% of the data. Despite the filtering, some few negative conductances remained. Although they are not theoretically possible, we retained them in the analysis because they represent the tails of the statistical distributions and they influenced the means. The sole exception was when we analysed the dependency of net photosynthetic rates on the conductances. In this one analysis, the negative values were deleted. The four cuvettes were treated as biological replicates, from which the mean hourly values were used for further analysis. Regression analysis was used to evaluate diurnal patterns. This included linear, polynomial, and nonlinear regression, as deemed appropriate. Correlations were treated as significant for $p \le 0.05$. All variability is given as standard error, unless stated otherwise. All statistical analyses were performed using the base package of R (version 3.3.2).

Results

Diurnal trends of g_s and g_m

We evaluated the diurnal trends in stomatal and mesophyll conductance, and in their ratio. Mean g_s was 0.115 (SE=0.002) mol CO₂ m⁻² s⁻¹ bar⁻¹. As expected, g_s showed a significant diurnal pattern (F=79.17, p <0.001), with peak values between 09:00 and 10:00, and decreased thereafter (Fig. 2a). We found a mean g_m value of 0.31 (SE=0.02) mol CO₂ m⁻² s⁻¹ bar⁻¹. Furthermore g_m also had a significant diurnal pattern (F=13.52, p <0.001; Fig. 2b) with relatively stable mean values between 08:00 and 16:00 and lower values in the early morning and towards the evening. The mean for the unitless ratio g_m/g_s was 2.67 (SE=0.3); with a weak, but significant diurnal pattern (F=3.9, p=0.02).

The relationship of A_{net} to g_s and g_m

 A_{net} followed a typical diurnal pattern, with highest rates between 08:00 and 12:00, with a mean of 16.2 (SE=0.26) µmol CO₂ m⁻² s⁻¹, and gradually declining rates in the afternoon (Fig. 2c). We found a significant asymptotic relationship between A_{net} and g_s (p < 0.001, $R^2 = 0.53$, Fig. 3a). Similarly there was a significant asymptotic relationship between A_{net} and g_m (p < 0.001, $R^2 = 0.28$, Fig. 3b).

Relationship of g_s and g_m to VPD

The average hourly VPD varied from 0.26 kPa to 1.82 kPa during the day, with sharp increase during the mornings until about 12:00 (Fig S4). We found a significant linear relationship between g_s and VPD (F=37.6, p < 0.001, $R^2=0.26$) (Fig. 4a) but no relationship between g_m and VPD (F=1.12, p=0.3, $R^2=0.001$) (Fig. 4b).

Contrasting estimates of WUE from $\Delta^{13}C$

We compared the performance of the three models to estimate WUE from Δ^{13} C (WUE_{Δ}) against direct measurements of WUE as A_{net}/E by the gas exchange system (WUE_G). Theoretically WUE_{Δ} and WUE_G should be identical with an ideal fit, where slope m = 1 and intercept a = 0. Using



Fig. 2 Diurnal variation in **a** stomatal conductance (g_s) , **b** mesophyll conductance (g_m) , and **c** net photosynthesis (A_{net}) . The points represent the cuvette means (n=4) for each hour and day. The blue line is the second order polynomial fit to the data and the shaded grey area is the standard error of the fit

the simple model resulted in a poor fit to our data (Fig. 5a). Further analysis revealed that this model could not predict WUE_G in the early hours (04:00–08:00), but fit the data well between 08:00 and 20:00 (a = -0.4, m = 1.0, $R^2 = 0.69$, Fig S5). The partial model consistently overestimated WUE_G by ca. 15% with no diurnal pattern (Fig. 5b). The comprehensive model matched the data well on average (Fig. 5c), but had a slight tendency to overestimate WUE_G in the low range and underestimate it in the high range. Furthermore, it introduced more variability into the estimates compared



Fig. 3 Relationship between net photosynthesis (A_{nel}) and **a** stomatal conductance (g_s), and **b** mesophyll conductance (g_m). The points represent the cuvette means (n = 4) for each hour and day. The blue line represents the asymptotic fit to the data



Fig. 4 Response of **a** stomatal (g_s) , and **b** mesophyll (g_m) conductance to vapour pressure deficit (*VPD*). The points represent the cuvette means (n=4) for each hour and day. The blue line represents the regression fit to the data and the shaded grey area is the standard error of the fit

to the partial model, with a residual standard error (SEr) of 2.7 mmol CO₂ mol H₂O⁻¹ ($R^2 = 0.61$) compared to 2.1 mmol CO₂ mol H₂O⁻¹ ($R^2 = 0.78$). In the comprehensive model, representing g_m as a constant ratio to g_s overestimated WUE_G by ca. 9% compared to observations (a = 0.7, m = 1.05, Fig. S6a). Assuming infinite g_m resulted in a poor fit to observed data (a = 0.7, m = 1.5, Fig. S6b) and an overestimation of WUE_G by 49%.

Discussion

Here, we report the first g_m estimates for mature *Pinus* sylvestris, one of the most widespread coniferous species of the northern hemisphere. Our mean value (0.31, $SE = 0.02 \text{ mol } \text{m}^{-2} \text{ s}^{-1} \text{ bar}^{-1}$) is somewhat higher than previously reported for other Pinus species (Flexas et al. 2008), but it is within the range of values reported for various conifers (Table 1) (De Lucia et al. 2003; Flexas et al. 2008; Bown et al. 2009; Bickford et al. 2010; Han 2011; Maseyk et al. 2011; Ubierna and Marshall 2011; Veromann-Jürgenson et al. 2017). Some of this variation may be due to differences in the methods used (Flexas et al. 2008). In particular, the "variable J method", which uses simultaneous measurements of gas exchange and chlorophyll fluorescence to infer g_m , generally yields lower $g_{\rm m}$ values than do isotopic techniques. If we exclude the "variable J" estimates from the list in Table 1, then our estimate of g_m matches the other values for conifers quite well. Furthermore, our estimates of δ^{13} C of A_{net} were in the range reported previously (e.g. Wingate et al. 2010) (Fig. S7), and are the first based on measurements using



Fig. 5 Water-use efficiency calculated from Δ^{13} C data (WUE_{Δ}) using **a** the simple model **b** the partial model, and **c** comprehensive model presented here. The results are compared to water-use efficiency calculated from continuous gas exchange data (WUE_G). In model **b** and **c** we assume a constant vale for $g_{\rm m}$ (0.29 mol CO₂ m⁻² s⁻¹ bar⁻¹).

The points represent the cuvette means (n=4) for each hour and day. The blue line is the linear fit to the data, the shaded grey area is the standard error of the fit, and m is the slope of the fit. The red line represents the theoretical 1:1 fit for comparison

 Table 1
 Mesophyll conductance values reported for various coniferous species

Measurement method	Species	$g_{\rm m} {\rm mol} {\rm m}^{-2} {\rm s}^{-1} {\rm bar}^{-1} * *$	References
Instantaneous Δ^{13} C and gas exchange (CRDS)	Pinus sylvestris	0.33 ± 0.03	Current paper
Instantaneous Δ^{13} C and gas exchange	Pinus halepensis	0.2–0.4	Maseyk et al. (2011)
Instantaneous Δ^{13} C and gas exchange (TDL)	Juniperus monosperma	0.04-4.0	Bickford et al. (2009)
Instantaneous Δ^{13} C and gas exchange	Pseudotsuga menziesii	0.14-0.20	Warren et al. (2003)
Phloem δ^{13} C	Pseudotsuga menziesii	0.22 ± 0.11	Ubierna and Marshall (2011)
Phloem $\delta^{13}C$	Thuja plicata	0.09 ± 0.04	Ubierna and Marshall (2011)
Phloem δ^{13} C	Larix occidentalis	0.41 ± 0.21	Ubierna and Marshall (2011)
Variable J method	Pinus radiata	0.153 ± 0.002	De Lucia et al. (2003)
Variable J method	Pinus sylvestris	0.124 ± 0.009	Veromann-Jürgensson et al. (2017)
Variable J method	Pinus radiata	0.16-0.2	Bown et al. (2009)
Variable J method	Pinus densiflora	0.107-0.250	Han (2011)
Various methods	Pinus spp.*	0.04–0.17	Flexas et al. (2008)

*Pinus lambertiana, Pinus monticola, Pinus pinaster, Pinus radiata

**If published values were presented as mol $m^{-2} s^{-1}$, then we assumed barometric pressure of 1 bar and converted the units to mol $m^{-2} s^{-1} bar^{-1}$

the CRDS (cavity ring-down spectroscopy) technology. The good agreement encourages us to suggest that this method, which is less expensive than most alternatives, produces reliable δ^{13} C measurements and is suitable for field applications. The CRDS was steady under the variable conditions of field-measurements, exemplified by the fact that we did not see any drift in the δ^{13} C values of the reference gases during several weeks of continuous measurements (Fig. S3). Nevertheless, it was crucial to correct the data for the CRDS' sensitivity to changing CO₂ and H₂O vapour concentrations. The sensitivity of isotope measurements to CO₂ concentration is a known phenomenon and is commonly corrected for in other laser technologies, like in lead alloy tunable diode lasers (Ubierna et al. 2018). Furthermore, drying the gas before isotope analysis will avoid having to correct for H_2O concentrations dependency, which is likely to improve measurement accuracy.

We observed an asynchronous reduction in g_m and g_s over the day (Fig. 2a, b). This may happen because g_s is under strong control by ambient *VPD* (Fig. 4a), whereas we found no correlation between g_m and *VPD* (Fig. 4b). The response of g_m to *VPD* has only been investigated in few studies, and with contrasting results. While Bongi and Loreto (1989) and Loucos et al. (2017) found a significant negative correlation between g_m and *VPD*, a study looking at the effect of air humidity and soil moisture (Warren 2008c) on g_s and g_m found a strong correlation of both conductances to soil moisture, while *VPD* only affected g_s and not g_m . All studies involved different species, and higher *VPD* ranges (1–3 kPa and 1–2 kPa, respectively) compared to our study (0.23–1.82 kPa). These discrepancies highlight the need for further investigation including a wider range of *VPD* conditions for *Pinus sylvestris*.

We fitted an asymptotic relationship of A_{net} to g_s and g_m . The asymptotic response agrees with theoretical expectations of CO₂ saturation at high conductances. In a diurnal context, Anet was maintained at high rates until midday, despite declining g_s from mid-morning (Fig. 2). We suggest that high g_m helped to maintain A_{net} during the late morning, enabling high $C_{\rm c}$ and compensating for the decline of g_s . This diurnal asynchrony between g_s and g_m is qualitatively similar to the observations by Theroux-Rancourt et al. (2014) on hybrid poplar cuttings exposed to soil drying over 12 days. They suggested based on daily measurements of g_s and g_m that a delayed g_m response reduced the decline in photosynthesis and enhanced WUE during the beginning of the drought treatment. Our finding suggests, that even within a diurnal context, the asynchronous response of g_m and g_s to environmental conditions has significant influence on A_{net} and presumably WUE.

We compared three models to estimate WUE from δ^{13} C. We found that the simple model can estimate WUE well for most of the day. This model uses \bar{b} , an empirical value that accounts for the drop of concentration between $C_{\rm i}$ and $C_{\rm c}$, the different fractionations occurring during photosynthetic discrimination, as well as possible postphotosynthetic discrimination. Many studies have shown that \bar{b} works well as an approximation (e.g. Farquhar et al. 1982; Seibt et al. 2008; Bloomfield et al. 2019). However, it performed poorly during the early morning, when WUE (Fig. S8), and especially δ^{13} C of photosynthesis (Fig. S7) were more variable. This meant that it produced unreliable WUE estimates for 25% of the photosynthetically active period of the day. Our analysis clearly shows that a more complete model that accounts explicitly for the effects of photorespiration, or both photorespiration and daytime respiration, performs better under such variable conditions and provides more accurate estimates of WUE. The relatively high variability in the estimates highlight the need to further refine some of the model assumptions.

We have shown that it is critical to account for g_m in the estimation of WUE from Δ^{13} C. This point has been made before (Seibt et al. 2008; Klein et al. 2015), but is still often neglected. Our data is yet an other example of WUE being overestimated if g_m is assumed to be infinite, and we show that assuming constant g_m/g_s or constant g_m both yield better estimates than infinite g_m . Estimating WUE from g_m/g_s had the further advantage of accounting for some of the diurnal change in g_m , resulting in a slope closer to 1 than when g_m was assumed constant. Nevertheless, this approach does not take into account the diurnality of g_m/g_s itself, and neglects the fact that g_m is much less strongly correlated with VPD than g_s .

The current study presents the first estimate of g_m for mature Pinus sylvestris trees, one of the most wide-ranging tree species in the world. Those estimates were derived with a CRDS/gas exchange system, which presents opportunities for simplifying the measurement of online Δ^{13} C discrimination. The measurements were made continuously and in the field over several sunny days in the summer. The high temporal resolution of our data allowed us to evaluate diurnal trends in conductance in relation to A_{net} , and test different models to estimate WUE form Δ^{13} C. Our analysis revealed that the simple model to account for ¹³C discrimination worked well, but only under stable conditions, and that the comprehensive model has the potential to account for variable conditions and provide reliable estimates of C_i and WUE. We highlight the need for further work under a broader range of environmental conditions, and including seasonal phenology. Our g_m estimate provides a means of improving inferences of WUE from Δ^{13} C and our continuous measurements provide a path forward to improve the modelling of $g_{\rm m}$ in the future.

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Compliance with ethical standards

Conflict of interest The authors have no conflict of interest to declare.

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