



Temporal patterns of carbon flow from grassland vegetation to soil microorganisms measured using ^{13}C -labelling and signature fatty acids

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Abstract

Purpose We investigated how the C flow from plants to microorganisms varies throughout the year in a temperate grassland. Additionally, we investigated how the C flow relates to saprotrophic activity and vegetation changes.

Methods In situ stable isotope pulse labelling ($^{13}\text{CO}_2$) was employed to estimate the flow of recently plant-derived C to soil microorganisms by using signature fatty acids. Bacterial and fungal growth was estimated using radio-labelling in laboratory incubations.

Results The C flow from plants to arbuscular mycorrhizal (AM) fungi peaked during the warmer parts of the year, but saprotrophic microorganisms showed little temporal variation in C flow. Also saprotrophic fungi received considerable amounts of C from plants throughout the year. Bacterial and fungal growth showed temporal variation with a growth peak in August for both. This suggests a shift in the C source from mainly rhizosphere C in colder parts of the year, to older C-sources in warmer parts of the year (August).

Conclusion We conclude that AM fungi, saprotrophic fungi and bacteria differ in the amount of recently-fixed C they receive from plants throughout the year. Hence, temporal patterns need to be considered to understand ecosystem functioning. The studied plant community

included winter annuals, which potentially maintain a high C flow to saprotrophic fungi during the cold season.

Keywords Arbuscular mycorrhiza (AM) · Fatty acids · Rhizosphere · Season · PLFA, ^{13}C · Temporal patterns

Introduction

In temperate ecosystems, seasonality in temperature and light affects plants, microorganisms and their interactions. Both plants and microorganisms can be active during the winter season, despite low temperatures (Brooks et al. 1997). The winter season is often overlooked, which may limit our understanding of important ecosystem functions (Campbell et al. 2005; Kreyling 2010; Makoto et al. 2014; Kreyling et al. 2019).

Photosynthetic activity is low during the winter because of low temperature, but also because of reduced light availability and soil moisture (Flanagan and Johnson 2005). However, plants can still photosynthesise during the cold season, and photosynthesis has been measured at temperatures as low as 0°C (Larsen et al. 2007). Carbon (C) assimilation (Lekberg et al. 2013) and nitrogen (N) uptake (Andresen and Michelsen 2005; Groffman et al. 2001) in plants have been shown to occur in temperate ecosystems during the cold season. In addition, soil respiration can be substantial in winter climate, often under a snow cover (Brooks et al. 1997; Alm et al. 1999; Larsen et al. 2007).

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The feedback between plants and soil microorganisms are fundamental to the global C cycle, and these feedback mechanisms are partly influenced by temperature (Lloyd and Taylor 1994). Plants contribute with C sources to soil microorganisms both through litter fall and rhizodeposition, and microorganisms make nutrients available for the plants through decomposition. Plants and microorganisms may also compete for the same nutrients, and this competition varies seasonally (Bardgett et al. 2003). The plant community composition can influence the microbial community composition, not only through variation in litter quality, but also because different plants offer different niche opportunities for the microorganisms (Fry et al. 2016), and because plants preferentially allocate C to beneficial microorganisms (Bever et al. 2009).

An important and beneficial group of microorganisms in grasslands are arbuscular mycorrhizal (AM) fungi, which supply nutrients to plants in exchange for C as part of a close symbiotic relationship. There is a lack of knowledge about how AM fungi grow and function in cold climates, but plants have been shown to benefit from AM fungi even at low temperatures (Tibbett and Cairney 2007). Lekberg et al. (2013) found that AM fungi receive plant-derived C also during the colder months of the year in a temperate coastal grassland, and that similar amounts of C were allocated to AM fungi in late autumn as in summer.

Soil bacteria and saprotrophic fungi are also influenced by annual climate variation (Bardgett et al. 1999; Birgander et al. 2014) since temperature is an important regulator for heterotrophic respiration (Bárcena-Moreno et al. 2009; Birgander et al. 2013; Walksman and Gerretsen 1931). Still, soil respiration has been detected even at temperatures as low as $-5\text{ }^{\circ}\text{C}$ in a laboratory study (Clein and Schimle 1995). Saprotrophic fungi are suggested to be relatively more dominant compared to bacteria during winter (Lipson et al. 2002) and to have a greater potential to be active during the colder months (Birgander et al. 2014).

More than 90% of the total soil microbial biomass has been suggested to be inactive (Blagodatskaya and Kuzyakov 2013), and dormancy likely varies over the year (Jones and Lennon 2010). Therefore, biomass-based measurements may not always be relevant for studies of temporal variation of microorganism importance. The use of *in situ* pulse labelling with stable carbon isotopes, in combination with biomass signature compounds, enables tracing of C from the atmosphere to

the soil microorganisms via plants (Olsson and Johnson 2005).

We investigated how different groups of microorganisms vary in their uptake of recently-derived plant C throughout the year, and how variation relates to microbial biomass, growth rates and vegetation composition. We hypothesized (1) that the plant-derived C in AM fungi would increase from the first time point that was in March (early spring) to the time points in June and August (representing the warmer season), and then decrease gradually again to October and then further during the cold season. We also expected (2) that a higher flow of C from plants to AM fungi would result in a higher AM fungal biomass. Incorporation of plant-derived C into bacteria and saprotrophic fungi was hypothesized (3) to show less temporal variation as they are less directly dependent on plants, and plant-derived C would thus peak in late August when plants are about to wilt and have less need for C for plant growth.

Materials and methods

Site description and sampling

Our study site was located in southern Sweden (N $55^{\circ} 42'$; E $14^{\circ} 10'$), in a temperate grassland with low grazing intensity (0.5 cow ha^{-1}) that had not received commercial fertilizer. The soil was classified as Eutric Cambisol according to the FAO system (Jones et al. 2005) and it was well drained. The vegetation consisted of mosses, grasses and low forbs. A $20 \times 20\text{ m}$ study plot was established in an area on flat ground that was visually judged to be as homogenous as possible and with the same vegetation type throughout (see Table S1 for species composition). Within this study plot we established 30 evenly distributed subplots (each $24 \times 43\text{ cm}$). We performed five pulse-labelling events between March 2011 and March 2012 (March 30, June 8, August 18 and October 24 in 2011, and March 16 in 2012). In the region of the study, in March it is early spring (but winter climate may still prevail), in June it is early summer, in August it is late summer and in October it is autumn. At each sampling time, six subplots were randomly assigned for labelling. No subplot was labelled more than once. Our study design includes replication within a relatively homogenous grassland plot, but it does not represent ecological replication across environmental gradients. Air temperature data

(mean annual temperature the week before sampling) was retrieved from Swedish Meteorological and Hydrological Institute for the weather station Skillinge (N 55° 49'; E 14° 31'). The warmest time-points sampled were June and August. In March 2011, the air temperature was lowest and was recorded as 4.1 °C the week before the labelling (Table 2). Vegetation analysis was performed on the same day as the pulse-labelling was performed, and soil and plant samples for isotopic analyses were collected 5 and 9 days (chase periods) after labelling. The 5-day soil samples were also used for analysing microbial biomasses and soil chemistry.

¹³CO₂ labelling in the field

Pulse labelling of the vegetation with ¹³CO₂ was carried out at five time points (see above). Transparent plastic hoods (35 × 20 × 12 cm) were placed over the entire vegetation at each of the six randomly assigned subplots, ensuring airtight conditions. A pulse of 50 ml 99.9% ¹³CO₂ gas (Cambridge Isotope Laboratories, Tewksbury, MA, USA) was injected into each hood using a gas syringe and then the hoods were sealed. The pulse period lasted for two hours and then the hoods were removed. Light intensity, estimated as photosynthetic photon flux density, was monitored during the pulse period, and labelling was performed at mid-day on days with a clear sky. Soil was collected with a soil corer (3 cm Ø) to 5 cm depth five days and nine days after labelling. Earlier studies in this area has shown that the A horizon is thin, and extending the sampling beyond 5 cm would also include a large part of relatively pure mineral sand. Two cores from each subplot (taken centrally at 20 cm distance) were combined into one composite sample (to reduce the number of samples) and brought back to the lab on ice for storage at -20 °C until further analysis.

Fatty acid analysis and ¹³C enrichment

Soil for analysis of fatty acid content was kept at -20 °C until all samples were collected, and they were analysed randomly for fatty acid content and stable isotopes in October–December 2012. Storage at -20° is considered appropriate for lipid analysis (Frostegård et al. 1991). Lipids were extracted according to Frostegård et al. (1993), using 5 g freeze-dried soil, extracted in a one-phase mixture of chloroform, methanol and citrate buffer, for two hours. Extracted lipids were fractionated into

neutral, intermediate and polar lipids by eluting with chloroform, acetone, and methanol, respectively, on pre-packed silica columns (van Aarle and Olsson 2003). The neutral and polar lipids (containing phospholipids) were subjected to mild alkaline methanolysis using fatty acid methyl ester 19:0 (for fatty acid taxonomy, see Tunlid and White 1992) as the internal standard and analysed on a gas chromatograph (Frostegård et al. 1993).

The bacterial biomass indicators were phospholipid fatty acids (PLFAs) i15:0, a15:0, i16:0, 16:1ω9, 16:1ω7, 10Me16:0, i17:0, cy17:0, 10Me17:0, 18:1ω7, 10Me18:0 and cy19:0 (Frostegård and Bååth 1996). PLFA 18:2ω6,9 was used as the indicator of ascomycetes and basidiomycetes and thereby mainly indicates saprotrophic fungi in grasslands (Frostegård and Bååth 1996). The neutral lipid fatty acid (NLFA) 16:1ω5 was used as the AM fungal biomass indicator. It has been shown to be more sensitive as an AM fungal signature and correlate better than PLFA 16:1ω5 with other indicators of AM fungal biomass, such as hyphal length density and root colonisation (Olsson et al. 1995; Olsson et al. 1998; Ven et al. 2020).

The flow of recently plant-fixed C to microorganisms was quantified by measurements of ¹³C in extracted PLFA and NLFA on a Delta IV Plus isotope ratio mass spectrometer (IRMS) coupled to a Trace GC Ultra gas chromatograph via the ConFlow IV interface (Thermo Fisher Scientific, Bremen, Germany) at the Stable Isotope Facility at the Department of Biology, Lund University, Sweden. The δ¹³C values were calibrated against the Pee Dee Belemnite (PDB) standard and corrected for the extra non-labelled methyl group from the methanolysis (eq. 1),

$$(nC * \delta^{13}C + Me) / (nC + 1) \quad (1)$$

where nC is the number of carbon in the fatty acid chain, and Me is the average δ¹³C in methanol. Thereafter samples were corrected for background ¹³C from non-labelled samples (eq. 2), giving the excess ¹³C in the fatty acids,

$$\delta^{13}C - \delta^{13}C_{\text{background}} \quad (2)$$

where δ¹³C_{background} is the measured δ¹³C in non-labelled samples (n = 5 and 8 for NLFA and PLFA, respectively). The non-labelled controls were the same as used in another study from the same grassland, and the exact values are published in Birgander et al. (2017).

Growth of bacteria and saprotrophic fungi

Soil was collected on the fifth day after pulse labelling, on each of the five labelling time points, and bacterial and fungal growth was analysed within 24 h of sampling. To standardize for the variation in soil temperature over the year, soil was kept on ice during handling and incubations for growth measurements were conducted at 14 °C.

Bacterial growth was estimated according to Bååth et al. (2001). Briefly, 1 g soil was mixed with 20 ml water and centrifuged at 1000×g. 1.5 ml of the supernatant (containing soil bacteria) was incubated with 2 µl radioactively labelled leucine ($[^3\text{H}]\text{Leu}$, 185 MBq ml⁻¹, 2 TBq mmol⁻¹, Perkin Elmer) and 2 µl of 200 mM unlabelled leucine, resulting in a concentration of 275 nM in the suspension. After 3 h of incubation at 14 °C, growth was terminated by addition of 75 µl 100% TCA. Samples were washed prior to measuring the incorporated radioactivity using liquid scintillation, which was then used as a proxy for bacterial growth.

Fungal growth was estimated using the acetate-ergosterol incorporation method, according to Bååth (2001), and modified by Rousk et al. (2009). Briefly, 1 g soil was incubated with 20 µl 1- $[^{14}\text{C}]$ acetic acid (sodium salt, 37 MBq ml⁻¹, 2.10 GBq mmol⁻¹, Perkin Elmer), 480 µl unlabelled 1 mM acetate and 1.5 ml water, resulting in a final concentration of 220 µM. After about 5 h of incubation at 14 °C, growth was terminated by adding 1 ml 5% formalin. Ergosterol was extracted and quantified with HPLC (UV detector, 282 nm), and collected using an automated fraction collector. Radioactivity incorporated into ergosterol was used as a proxy for fungal growth after measurement using liquid scintillation.

Soil analyses

Soil organic matter (SOM) content was determined gravimetrically as loss on ignition (LOI) on dried samples (5 h, 550 °C), assuming LOI = total soil organic matter content. Ammonium (NH₄⁺) and nitrate (NO₃⁻) were extracted from soil within 24 h after sampling, in 0.1 M BaCl₂ (1:5) and measured using flow injection analysis for soil samples from June 2011, August 2011 and March 2012, while data is missing for the other time points.

Vegetation analysis

In connection to pulse labelling, the vegetation was surveyed in 1 m² plots. In four replicated plots, coverage of vascular plants, mosses and lichens was estimated. Identification to species was undertaken in most cases, however, some families where species were hard to distinguish were identified to family level (see supplementary Table 1). Plant species were classified as perennial, summer annuals or winter annuals if possible, according to classification by Grime et al. (1996). The mycorrhizal status of plants was classified according to Harley and Harley (1987). The functional classification was used to independently calculate the proportion non-mycorrhizal species as well as the proportion annual species.

Statistical analyses

Temporal variation of ¹³C enrichment was tested at day 9, since the enrichment was in nine out of 15 cases (Table 1) highest 9 days after pulse labelling. Variations of ¹³C enrichment, growth rates, and biomass between time points were tested with one-way ANOVA with date as group variable, and when significant, a Tukey post hoc test was used to identify which time points had differed significantly. To test for differences between sampling days, i.e. between sampling 5 and 9 days after pulse labelling, a paired t-test was used. All data were tested for equal variance, and ¹³C enrichments and growth rate measurements were log transformed to fulfil the assumptions for statistical analysis. A correspondence analysis was performed to assess vegetation changes. The first two axes were tested for temporal variation with repeated measures ANOVA. Data were square root transformed to obtain equal variance and time was used as within subject factor.

All statistical analyses were performed in SPSS (SPSS Statistics, IBM Corporation, New York, USA), Correspondence analysis (CA) of the vegetation composition were performed in CANOCO 4.5 (Ter Braak and Smilauer 2002), and graphs were constructed using KaleidaGraph (Synergy Software, Reading, USA).

Results

All microbial indicators showed ¹³C enrichment, showing that the in situ ¹³C labelling was successful. The

Table 1 The percent ground coverage was estimated at five time points over the course of one year. Values represent the mean of 4 replicates \pm SE. Plants were classified as annuals or perennials according to Grime et al. 1996 or Tyler et al. 2021, and mycorrhizal status was classified according to Harley and Harley (1987).

Ground cover (%)	Time of the year				
	March 30, 2011	June 8, 2011	August 18, 2011	October 24, 2011	March 16, 2012
Bare soil	13.8 \pm 6.2 a	7.8 \pm 2.8 a	5.8 \pm 1.5 a	4.5 \pm 2.0 a	31.5 \pm 1.25 b
Bryophytes	59.8 \pm 7.6 a	26.2 \pm 6.2 b	28.8 \pm 8.3 b	35.2 \pm 2.8 ab	29.0 \pm 5.4 b
Lichens	2.25 \pm 0.25	3.5 \pm 1.0	2.0 \pm 0.7	1.2 \pm 0.2	2.0 \pm 0.4
Annual herbs	16.8 \pm 2.7 ab	10.8 \pm 1.1 ab	7.5 \pm 1.7 a	18.2 \pm 3.2 b	9.5 \pm 1.8 ab
Perennial herbs	26.5 \pm 6.3 a	68.8 \pm 4.2 bc	90.0 \pm 6.5 c	56.2 \pm 10.2 b	45.8 \pm 3.2 ab
Proportion of vegetation cover (%)					
Annuals	41.2 \pm 10.4 a	13.7 \pm 1.6 b	7.9 \pm 1.8 b	26.0 \pm 5.6 ab	17.5 \pm 3.7 ab
Non-mycorrhizal species	17.0 \pm 1.6 ab	9.6 \pm 1.2 bc	7.9 \pm 1.4 c	11.8 \pm 3.1 abc	18.7 \pm 1.9 a

From this, the proportion of annuals and non-mycorrhizal (NM) vegetation cover was calculated independently (see Table S1 for classification of species in this study). Different letters indicate significant difference between time points ($p < 0.05$) tested with one-way ANOVA and Tukey post hoc test

enrichment in NLFA 16:1 ω 5, indicating AM fungi, varied throughout the year ($F_{4,17} = 8.38$, $p < 0.001$, 9 days after labelling) with ^{13}C enrichment in June and August twice as high as compared to March, in both 2011 and 2012 (Fig. 1a). When enrichment was highest, in June and August, there was a tendency for higher enrichment on day 9 as compared to day 5. The saprotrophic fungal indicator PLFA 18:2 ω 6,9 also received considerable amounts of plant-derived C; at some time-points receiving even higher enrichment than AM fungi and consistently more than bacterial PLFAs (Fig. 1b). There was no significant temporal variation in ^{13}C enrichment of PLFA 18:2 ω 6,9, measured 9 days after labelling, but ^{13}C enrichment tended to be highest in August. In August and October significantly more plant-derived C was incorporated into PLFA 18:2 ω 6,9 after 9 days as compared to 5 days after labelling ($t_3 = -7.47$, $p < 0.005$, and $t_2 = -5.02$, $p < 0.05$, respectively). Bacterial PLFAs showed ^{13}C enrichment throughout the season, but almost one magnitude lower enrichment compared to AM fungi and saprotrophic fungi (Fig. 1). Only little temporal variation in ^{13}C enrichment was seen for bacterial PLFAs, and no significant difference was found between day 5 and day 9.

Bacterial biomass and saprotrophic fungal biomass, indicated by PLFAs, and AM fungal biomass, indicated by NLFA 16:1 ω 5, measured 9 days after labelling, did not change significantly over the year (Fig. 2a–c). Growth rates showed temporal variation for both saprotrophic fungi and bacteria ($F_{4,25} = 4.83$, $p = 0.005$, and $F_{4,25} = 10.84$, $p < 0.001$, respectively,

Fig. 3a, b). The saprotrophic fungal growth rate was lowest in March 2011 and peaked in August. The bacterial growth rate also peaked in August but was lowest in October.

Plant species composition differed between the different sampling times both along the first and the second axis from the CA (Fig. 4, $F_{1,8} = 22.3$, $p < 0.005$, $F_{1,5} = 12.6$, $p < 0.05$, respectively), but the two sampling times in March did not differ along PC 1. Species composition in March was more dominated by mosses, while grasses were more abundant in June and August, although the moss *Hypnum cupressiforme* was the most dominant species at all time points (Table S-1). In March 2011 and 2012, substantial parts of the soil were bare (Table 1). The plants classified as winter annuals were most common in March and October (i.e. spring and autumn; *Bromus hordeaceus*, *Cerastium semidecandrum*, *Erophilia verna*, and *Trifolium campestre*), while plants classified as being both winter and summer annual plants were found year-round (*Arenaria serpyllifolia*, *Erodium cicutarium* and *Geranium molle*). Several of the plants classified as winter annual plants are non-mycorrhizal, while most perennial plants are arbuscular mycorrhizal (Table S-1). The proportion of non-mycorrhizal species was highest in March 2012, followed by March 2011. The proportion of annual species was highest in March 2011, but was also high in October the same year (Table 1).

Both NH_4^+ and NO_3^- differed slightly over the course of the experiment ($F_2 = 4.69$, $F_4 = 4.57$, respectively, both $p < 0.05$), with less NH_4^+ in June compared

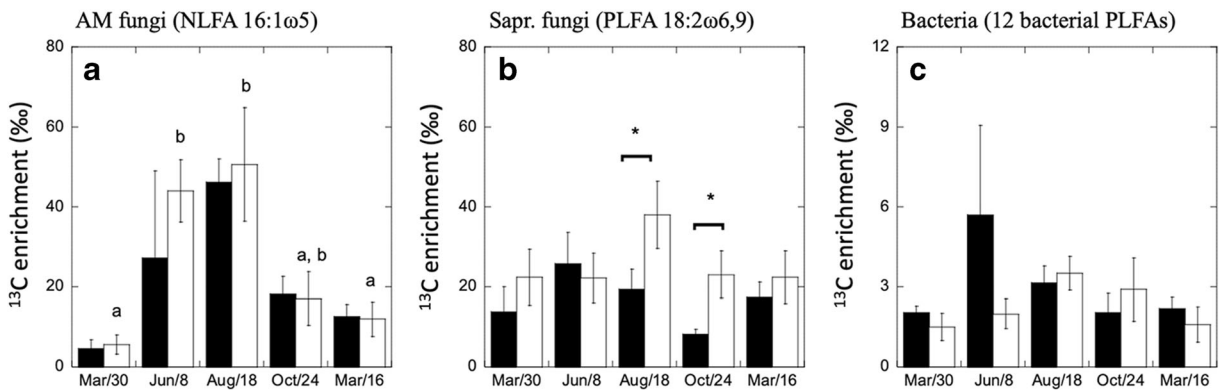


Fig. 1 Enrichment of plant-derived ^{13}C over the year, into (a) arbuscular mycorrhizal fungi, indicated by NLFA 16:1 ω 5, (b) saprotrophic fungi, indicated by PLFA 18:2 ω 6,9 (c) bacteria, indicated by the sum of 12 bacterial PLFAs. Black bars = 5 days after labelling, and open bars = 9 days after labelling. Different letters indicate significant difference ($p < 0.05$) in enrichment at

day 9, tested with one-way ANOVA and Tukey post hoc test. * = significant difference ($p < 0.05$) between sampling days (paired t-test). Values ($n = 3\text{--}6$; \pm SE) were log transformed prior to analyses. Measurements could not be achieved in some of the samples due to too low amounts for isotopic analysis

to March 2012 and more NO_3^- in October 2011 compared to March 2012 (Table 2). The peak in NH_4^+ occurred in March 2012, when the NO_3^- was lowest, while NO_3^- peaked in August. Soil moisture was lowest in June and highest in March 2011 ($F_4 = 28.15$, $p < 0.001$, Table 2).

Discussion

We found temporal difference between AM fungi, saprotrophic fungi and bacteria differ in how much recently-fixed C they receive from plants. There was a clear difference between AM fungi and saprotrophic fungi, in the turnover of recently-fixed C in the studied semi-natural temperate grassland, and between C flow

from plants to microbes and microbial saprotrophic activity. This emphasise the importance of considering the whole year when studying or modelling effects of global warming on ecosystem C cycling (Kreyling et al. 2019). The microbial community in the semi-natural grassland in this study was strongly influenced by nutrient limitation (Birgander et al. 2014). Plants and saprotrophic microorganisms are known to compete for N in several types of ecosystems (Kaye and Hart 1997; Kuzyakov and Xu 2013), but temporal niche separation in the demand of N for plants and microbes have been suggested to reduce the competition (Hodge et al. 2000; Kuzyakov and Xu 2013). Lowest concentration of inorganic N was found in June, when plant demand was likely to be high. This in addition to the seasonal vegetation changes we observed, highlight the

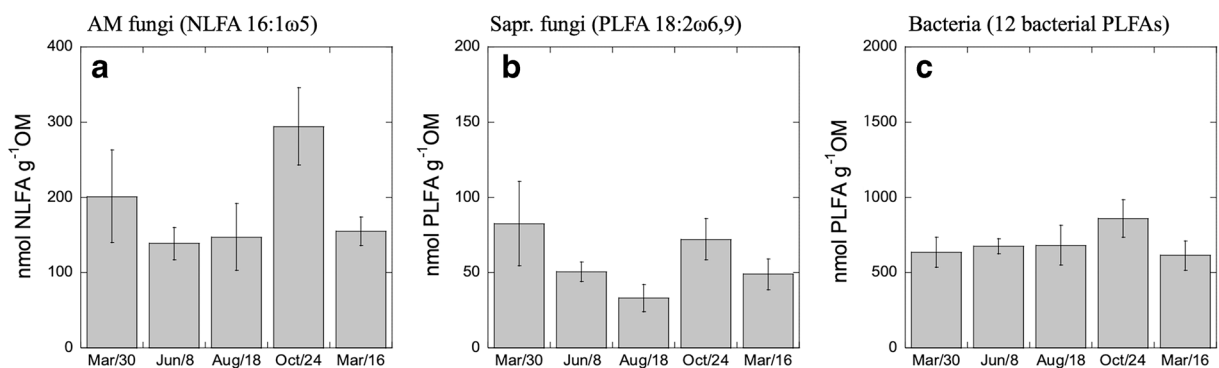


Fig. 2 Biomass of (a) arbuscular mycorrhizal fungi, indicated by NLFA 16:1 ω 5, (b) saprotrophic fungi indicated by PLFA 18:2 ω 6 and (c) bacterial, indicated by the mean of 12 bacterial PLFAs, from March 2011 to March 2012. None of the biomass values

differed significantly over the year. Values represent mean of 6 replicates \pm SE for (b) and (c), but only 3–6 replicates in (a) due to too low amounts

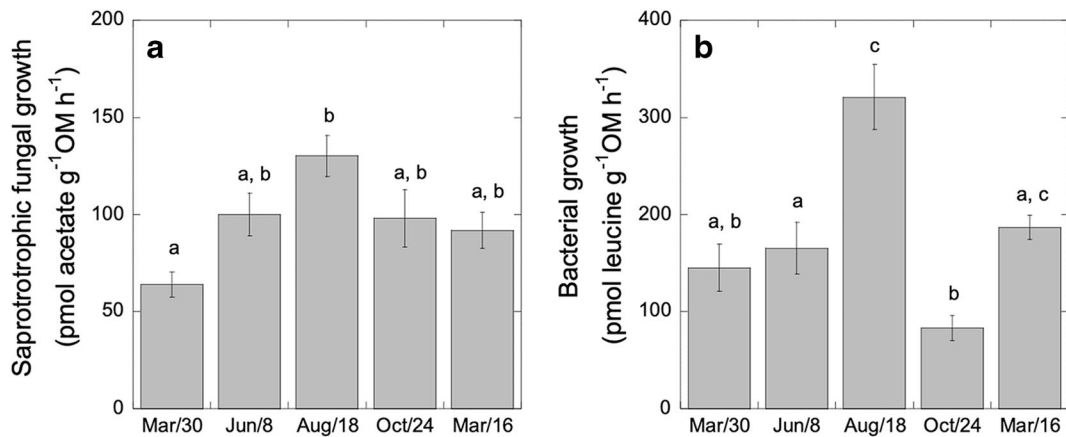


Fig. 3 Temporal variation in growth for (a) saprotrophic fungi and for (b) bacteria, from March 2011 to March 2012. Different letters indicate significant difference in growth at the different times ($p < 0.05$), investigated with ANOVA and Tukey post hoc

test. Values were log transformed in statistical analysis to fulfil criteria of equal variance. Values represents mean of 6 replicates \pm SE

importance of considering the whole year when studying plant-microbe interactions and how these interactions influence climate change effects (Birgander et al. 2017; Kreyling 2010; Kreyling et al. 2019). Still, we should keep in mind that the present study was replicated within a rather homogenous part of a semi-natural grassland, and further studies that replicate across environmental gradients would significantly contribute to the knowledge in this field.

Microbial biomass did not change significantly over the year for any of the microbial groups. Hence, our hypothesis, that a high C flow from plants to AM fungi

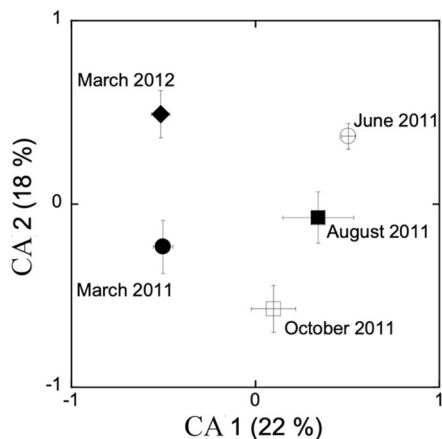


Fig. 4 Correspondence analysis (CA) of the plant species composition along the first two axes, at five time points (March 2011 closed circle, June 2011 open circle, August 2011 closed square, October 2011 open square, and March 2012 closed diamond). Sum of all eigenvalues = 1.149. Values presented are mean of 4 \pm SE

would be followed by an increase in biomass, was not supported. However, there was a tendency for increased AM fungal biomass (storage lipids) in late autumn, possibly due to increased C input to AM fungi during the summer. Most likely, high biomass during the winter, particularly that of AM fungi, consists of a large body of less active biomass diluting the response in biomass estimates (Blagodatskaya and Kuzyakov 2013). The C flow from plants to soil microorganisms has previously been shown not to correlate with temperature (Birgander et al. 2017). Lack of correlation between biomass and activity (e.g. growth) for saprotrophic fungi and bacteria could indicate that changes in activity relates more to changes in turnover rates and does not therefore lead to changes in biomass (Rousk and Bååth 2011). In addition, soil freezing may influence the activity of fungivores such as microarthropods and nematodes in the soil (Sulkava and Huhta 2003). The chosen signature compound for AM fungi, NLFA 16:1 ω 5, is part of the storage lipids and may not be related directly to biomass. However, earlier studies have shown that this compound is more sensitive as an indicator of AM fungal mycelium (Olsson et al. 1995) and it therefore can be expected to have a higher “signal to noise” ratio in environmental samples. Furthermore, since the NLFA 16:1 ω 5 constitutes a major energy storage and energy transfer component in AM fungal mycelium (Van Aarle and Olsson 2003), it is thus not surprising that its content correlate well with the amount of viable mycelium (Ven et al. 2020) and root colonisation, although this correlation

Table 2 Variation in environmental factors. Different letters indicate significant difference between time points ($p < 0.05$), tested with one-way ANOVA and Tukey post hoc test. For soil moisture

and soil organic matter, a mean of 24 samples \pm SE is presented. For NH_4^+ and NO_3^- , the mean of 6 replicates \pm SE is presented

	March 30	June 8	August 18	October 24	March 16
Mean air temperature the week before ($^{\circ}\text{C}$)	4.1	15.1	15.9	7.6	5.3
Soil moisture (mg g^{-1} soil)	157 \pm 9 a	54 \pm 8 b	106 \pm 5 c	117 \pm 5 c	117 \pm 7 c
Organic matter (mg g^{-1} soil)	62 \pm 5 a	31 \pm 12 b	37 \pm 2 b	36 \pm 3 b	45 \pm 3 a,b
NH_4^+ ($\mu\text{g g}^{-1}$ soil)		2.1 \pm 0.2 a	3.2 \pm 0.4 a,b		3.5 \pm 0.4 b
NO_3^- ($\mu\text{g g}^{-1}$ soil)		0.42 \pm 0.06 a,b	0.84 \pm 0.23 a		0.25 \pm 0.06 b

differs between AM fungal species (Van Aarle and Olsson 2003).

C allocation to AM fungi showed temporal variation, which was in accordance with our hypothesis. The highest ^{13}C enrichment in AM fungi was seen in June and August, the two warmest months in our study, representing summer. Plants can allocate C to AM fungi when the symbiosis is particularly beneficial (Bever et al. 2009). AM fungi have therefore been suggested to benefit from low nutrient conditions (Anderson et al. 1984), potentially due to an increased dependence of, and preferential C allocation from nutrient limited plants. In contrast to AM fungi, saprotrophic fungi showed little temporal variation in ^{13}C enrichment. The C flow, however, was in the same range as the C flow to AM fungi, and the enrichment was about six times higher in the PLFA signature for saprotrophic fungi than average bacterial PLFAs. It may seem strange that saprotrophic fungi acquire such a large proportion of the plant-derived C during the whole year, but saprotrophic fungi have been suggested to thrive during the colder parts of the year (Bardgett et al. 1999; Lipson et al. 2002).

The high enrichment of recently-fixed C in saprotrophic fungi contrasted with the assumption that bacteria dominate saprotrophic activity in the rhizosphere (Wardle et al. 2004), but is in line with recent studies indicating that saprotrophic fungi are more important than bacteria in the turnover of rhizosphere C (Birgander et al. 2017; Pausch et al. 2016; Streit et al. 2014). The substantial amount of plant-derived C in saprotrophic fungi indicates a close connection between saprotrophic fungi and plants. However, the PLFA indicator we used for saprotrophic fungi (PLFA 18:2 ω 6,9) is common for all ascomycetes and basidiomycetes, which also include some root endophytes. Dark septate endophytes, common in forbs and grasses

in most terrestrial ecosystems (Read and Haselwandter 1981) may therefore also contribute to enrichment in PLFA 18:2 ω 6,9. Little is known about these endophytic fungi and their contribution to turnover of rhizosphere C, and studies on their functional role, including studies on the saprotrophic and symbiotic activities in this group, would be very informative.

Neither bacteria nor saprotrophic fungi showed any temporal variation in ^{13}C enrichment, as judged from the chosen signature PLFAs. Overall, the ^{13}C enrichment in bacteria was lower than the enrichment in AM and saprotrophic fungi. Both bacterial and saprotrophic fungal growth rates varied throughout the year with highest growth in August, which is in accordance with previous studies (Birgander et al. 2014). We found little temporal variation in the C-flow from plants to microorganisms, but a clear temporal variation in their growth rates. This indicates that other C sources, such as litter and soil organic matter, are tentatively more important drivers of soil respiration during the warmer season.

The C flow from plants to different groups of microorganisms did not differ due to length of the chase period – except for saprotrophic fungi in August and October. From a previous labelling study at a cold time of the year, we know that the peak in ^{13}C enrichment occurs before 5 days after labelling for all groups (Birgander et al. 2017). As discussed above, seasonality influences the activity of plants and thus probably the timing of C flow from plants to microorganisms as well (Kuzyakova and Gavrichkova 2010). For example, five days after a pulse in summer does not correspond to five days after a pulse in spring, and this obscures differences between sampling days.

We conclude that the flow of plant-derived C to soil microorganisms in temperate grassland may show significant temporal variation. The C allocation to AM fungi peaks in the warmer part of the year (August in

our study), while during the colder parts of the year (March and October in our study), the highest proportion of the plant-derived C flows to saprotrophic fungi. Winter annual plants, which are also active during the colder parts of the year, could be one explanation for the continuous C flow to saprotrophic fungi. There is a clear temporal difference between AM fungi and saprotrophic fungi, while bacteria seem to be a less influential player in the turnover of rhizosphere C in the studied semi-natural temperate grassland. Further studies are needed to show if similar variation is found across ecological gradients in grasslands, and other types of vegetation. The differences in C flow from plants to different groups of microorganisms in the rhizosphere and their responses to seasonal climate variation are important in predicting their response to climate change. Earlier studies have shown that warmer winters may stimulate mycorrhizal C-flow in grasslands (Birgander et al. 2017) and may favour more productive plant species (Keyling et al. 2019). Warmer winters can thus be expected to fundamentally influence grassland plant communities, and therefore have implications for conservation biology as well as ecosystem C sequestration.

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Declarations

Ethics approval and consent to participate We have no potential conflicts of interest.

Consent for publication All authors have agreed on submission of this manuscript to Plant and Soil for publication.

Competing interests We have no competing interest to declare.

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