

CO₂ COOLING IN TERRESTRIAL PLANET THERMOSPHERES

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The comparative approach to planetary problems is becoming increasingly fruitful as new information from various planetary atmospheres is assimilated. As an example, it is clear that the important problem of CO₂ cooling in the Earth's lower thermosphere is closely tied to the thermospheric heat budgets of Venus and Mars.

The thermospheres of Earth, Venus, and Mars are cooled by a combination of molecular/eddy conduction and infrared (IR) cooling processes. Molecular conduction cools the upper thermosphere by transferring the EUV or auroral heat down-gradient toward the mesopause, where it is effectively radiated to space by an IR active cooling agent. The height and temperature of the mesopause are controlled by the effectiveness of the IR cooling processes. For all three planets, this IR cooling results primarily from CO₂ emission at 15- μ m. Atomic oxygen collisions are known to be especially effective in exciting CO₂ vibrational states, resulting in enhanced CO₂ 15- μ m emissions and cooling at thermospheric heights where non-Local Thermodynamic Equilibrium (NLTE) prevails. The effectiveness of this enhancement process depends upon the relative O densities at a given pressure level plus the collisional energy transfer rate coefficient (CO₂-O) specified. The importance of this mechanism for enhancing CO₂ cooling on Earth and Venus has been debated since 1970. For Venus, a longstanding dayside heat budget problem can only be resolved by understanding the relative role of CO₂ cooling [1,2,3]; the Venus atomic oxygen density is well measured over the solar cycle. For the Earth, the lower thermosphere heat budget is difficult to understand without a thorough characterization of the variability of atomic oxygen and CO₂ cooling in that region.

Progress has been made recently that improves our understanding of CO₂ cooling processes in the Venus, Earth, and Mars lower thermospheres. First, the corresponding CO₂-O relaxation rate has been measured in the laboratory at room temperature for the first time ($k = 1.2 \times 10^{-12} \text{ cm}^3/\text{sec}$). Also, derived values of this relaxation rate are becoming better constrained (1.6×10^{-12}) and are based primarily upon analyses of Earth CO₂ radiance and absorption measurements (see Table 1 and references therein). We see that these recently measured and derived values are not entirely consistent. Furthermore, the temperature dependence of this rate over 200-400 K has yet to be measured. This rate must be consistently incorporated into calculations of all three terrestrial planet thermospheres.

Independent one dimensional (1-D) model studies of the Venus dayside heat budget [2,3] were recently conducted which confirm the need for a fast ($2-4 \times 10^{-12}$ cm³/sec) relaxation rate [2,3]. Also, the newly simulated Mars dayside heat budget [3] is consistent with CO₂ cooling derived from a fast ($2-3 \times 10^{-12}$) relaxation rate. For Venus, this strong cooling serves as an effective thermostat that gives rise to a small variation of thermospheric temperatures over the solar cycle, just as observed. Conversely, CO₂ cooling does not appear to be dominant in the day-side heat budget of the Mars thermosphere, yielding larger observed temperature variations. For the Earth, this strong cooling implies that the lower thermosphere does not require significant eddy diffusion or heat conduction, consistent with independent estimates of a weak mesopause eddy coefficient. Our ability to predict the clearly visible thermospheric "CO₂ greenhouse" cooling response into the 21st Century is directly dependent upon the proper incorporation of this CO₂-O relaxation rate in current modeling schemes [4]. This comparative approach provides the broadest range of conditions under which a common CO₂-O relaxation rate should provide consistent results.

Future measurements are needed that address the laboratory CO₂-O rate over 200-400 K. Also, the discrepancy between laboratory and atmospheric derived values of this rate must be resolved. In particular, what role does hot oxygen play in exciting CO₂ vibrational states, thereby enhancing CO₂ 15- μ m cooling?

TABLE 1. Recent Estimates of CO₂-O Relaxation Rates

Value (10 ⁻¹² CM ³ /SEC)	Temp. (K)	Technique	Reference
6±3	300	CO ₂ Radiance (SPIRE Rocket)	[8]
5-6	200-400	CO ₂ Radiance (SPIRE Rocket)	[10]
1.5±0.5	300	Laboratory	[9]
1.2±0.2	306	Laboratory	[6]
1.5 - 6	300	CO ₂ Absorption (ATMOS)	[7]
3-6	200-400	CO ₂ Absorption (ATMOS)	[5]
2-4	300	Venus GM Model	[2,3]
2-3	300	Mars GM Model	[3]

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