

First record of fossil anguines (Squamata; Anguinae) from the Oligocene and Miocene of Turkey

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Abstract Fossil anguine lizard specimens from several Turkish localities are described in this paper. The material comes from ten different localities, spanning a large geographic area consisting of both parts of the European Turkey and Anatolia, and ranging in age from the Oligocene to the Late Miocene. In certain cases, the generic determination was possible and, accordingly, members of *Ophisaurus* and *Anguis* were identified and described in detail. The specimens of *Anguis*, found in different, Middle and Late Miocene localities from Anatolia, represent two of only a few fossil occurrences of this taxon. Moreover, the material reported herein represents the oldest occurrences of anguine lizards, not only from Turkey, but from southeastern Europe and the Eastern Mediterranean basin as a whole. These rare records provide important information about the dispersal routes of anguines from Europe

to Asia and significantly enhance our understanding of their biogeography.

Keywords Anguimorpha · Cenozoic · Dispersion · Asia

1 Introduction

The clade Anguinae (Anguimorpha, Anguinae) comprises legless lizards and includes three extant taxa: *Pseudopus* (Southeast Europe to Central Asia), *Anguis* (Europe, Western Asia) and *Ophisaurus* (North America, Northern Africa and Southeast Asia). On the basis of the fossil finds as well as the mitochondrial DNA of the extant species it is believed that the ancestor of the Anguinae probably originated in North America. From there it would have migrated during the Paleocene and Eocene via the land bridge of Greenland and Spitsbergen from North America to Europe (Klembara 1981; Macey et al. 1999). The earliest unambiguous records of this clade in Europe come from the Eocene of Belgium, France, and Great Britain probably representing the genera *Ophisaurus* and *Anguis* (Augé 1990, 2005; Klembara and Green 2010). The oldest known certain record of *Pseudopus* is from the Early Miocene of Germany (Čerňanský et al. 2015), but it is not possible to exclude its presence already in the Late Oligocene of Germany (Čerňanský et al. 2016).

In Europe, the fossil history of the Anguinae clade is well documented from the Cenozoic era (Klembara and Rummel 2016 and references therein; Georgalis et al. 2016a). *Anguis* and *Pseudopus* persist in Europe and Western Asia to the present, but *Ophisaurus* is now absent from that region and has dispersed during the Cenozoic into North Africa (Blain et al. 2013), East Asia and North America (see discussion in Klembara and Rummel 2016; Vasilyan et al. 2017). One of the crucial geographical areas, as for the eastward dispersal

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of Anguinae, is the region of modern day Turkey. However, except for the boine snake *Bavarioboa* from the Late Oligocene or possibly the Early Miocene of eastern Turkey (Szyndlar and Hoşgör 2012), no other squamates have been recorded from Oligocene and Miocene deposits in Turkey. As for Anguinae, jaws, vertebrae and osteoderms of *Ophisaurus* and *Pseudopus* have been documented only from the Pliocene and Pleistocene of Turkey (Rage and Sen 1976; Venczel and Sen 1994; Zwick and Schleich 1994). *Pseudopus* was also reported from the Quaternary of Lebanon (Hooijer 1961) and Israel (Bar-Yosef and Tchernov 1966; Maul et al. 2010; Smith et al. 2015).

Here we describe for the first time, material pertaining to anguines from several Cenozoic Turkish localities of different age (Fig. 1), extending from the Oligocene (Kocayarma MP 23–MP 27, MP 25?) to the Late Miocene (Süleymanlı, MN 13). Although the material is fragmentary, it provides important information about the evolution of this clade in Anatolia and southeastern Europe, and supports the existence of geographical links between terrestrial squamate faunas between Europe and southwestern Asia from the Oligocene to the Early Miocene.

2 Materials and methods

Most fossils described herein are deposited in the Natural History Museum of Ege University in Izmir (Turkey) and are prefixed following the format “EUNMH PV-number”.

Other material is stored at the Palaeontological collections of the Faculty of Geosciences, University of Utrecht (The Netherlands) and are prefixed by the following formats: “UU KE number”, “UU KAR number”, “UU CD number”, “UU BAG number”, and “UU SLM number”. Photographs were taken with a Leica DVM5000 at the University of Tübingen, Germany and a Leica M205 C binocular microscope with axially mounted DFC 290 HD camera, LAS software (Leica Application Suite) version 4.1.0 (build 1264) at the Comenius University in Bratislava, Slovakia, and a Leica M205 microscope and the Leica application suite V 3.3.0 at the University of Torino, Italy. The dentary KAR 1102 was also photographed with scanning electron microscope (SEM—FEI XL30 Sirion FEG) at University of Fribourg, Switzerland). The image processing program ImageJ (Schneider et al. 2012) was used for measurements.

3 Geological settings

The fossil material described in this paper originates from ten different Late Oligocene to Late Miocene localities in Turkey (Table 1). Only the most relevant data concerning the geological settings of these localities are presented below.

Kavakdere and Kocayarma On the basis of micro-mammals, the age of both these faunas was determined by Ünay (1989) as Oligocene, within the range of MP 23–MP



Fig. 1 Map of Turkey with indication of the fossil localities with anguine specimens described herein

Table 1 List of taxa examined from Turkish Oligocene and Miocene localities

	<i>Ophisaurus</i>	<i>Anguis</i>	Anguinae indet.	Age
Kocayarma			+	MP 25
Kavakdere			+	MP 26/27
Kargi 2	+		+	MP 30–MN 1
Kargi 3			+	MN 1
Kılçak 3b			+	MN 1
Sabuncubeli			+	MN 3
Keseköy	+		+	MN 3
Çandır and Çandır HW	+		+	MN 6
Bağıcı	+	+	+	MN 7/8
Süleymanlı	+	+	+	MN 13

27, while she considered that the association of two rodents, a rather large-sized *Suevosciurus ehingensis* and *Edirnella*, is indicative of MP 25. This age is also consistent with that of radiometric age (33–32.2 Ma) that was sampled from the volcanic tuff overlying the Kavakdere fossiliferous zone (Ünay 1989).

Kargi The section is situated in a coal quarry located near the village of Dodurga. The deposits show a rhythmic alternation of white limestones and dark green clays. Three fossiliferous layers in the mine have been sampled for mammal fossils. The biostratigraphic correlations suggest that Kargi 2 lies at the Oligocene–Miocene transition (local zone A, MP 30–MN 1), and Kargi 3 is of Early Miocene age (local zone B, MN1) (Wessels et al. 2001; van den Hoek Ostende 2001; Ünay et al. 2003; de Bruijn et al. 2013; Kakali 2013).

Kılçak The sections from where fossils were collected are open lignite mines, located near the village of Kılçak. The whole area is completely disturbed and covered by landslides, therefore lithostratigraphic correlation of layers is difficult to be observed. Sediments are mainly grey-green clay, rich in mollusk debris, with several thin lignite horizons. The fossiliferous level Kılçak 3b is one of the layers sampled by a Dutch team during the early 1990s, with the other layers named Kılçak 0, Kılçak 0', and Kılçak 3a. On the basis of rodent assemblages, Kılçak 0 and 0' are considered to represent the biostratigraphically oldest levels while Kılçak 3b is the youngest (de Bruijn et al. 1993; de Bruijn and von Koenigswald 1994; Sen et al. 1998). There is a general agreement in the attribution of these assemblages to the local biozone B that is correlated to MN 1 (de Bruijn et al. 2013).

Sabuncubeli The fauna has been collected from the lower part of the Soma Formation, outcropped near the road between Izmir and Manisa. The fossiliferous bed consists of a lower part with fine-grained conglomerate lenses and an upper marly part which contains mollusk debris. The age of the micromammal assemblages was interpreted by de Bruijn et al. (2006) as pertaining to the

lower part of MN 3 (local zone C/D), being thus slightly older than Keseköy.

Keseköy This locality is a coal quarry near the town of Kizilcahamam. The section predominantly consists of green–brown, partly laminated clays, intercalated with several coal layers. It contains an assemblage of small mammals that is attributed to the local zone D, being correlated to MN 3 (Ünay et al. 2003; de Bruijn et al. 2013).

Çandır and Çandır HW These localities are represented by several fossiliferous horizons of green-grey silty clays and paleosols, located near Çandır. The correlation of the Çandır assemblages to an appropriate MN zonation is complicated. Çandır was initially biostratigraphically correlated to MN 6 (de Bruijn et al. 1992) and was subsequently reassigned to the MN5 (de Bruijn et al. 2003). As such, the age of Çandır is still under debate and is thus dated either as MN5 or MN6 (Casanovas-Vilar et al. 2011; Mayda et al. 2015). Using both biochronological correlations MN 5 and MN 6, the magnetostratigraphic studies (Krijgsman 2003) correlated Çandır to the base of the chron C5ACn, 14 Ma and C5ABn, 13.6 Ma respectively. In any case, the overall faunal resemblance between Paşalar and Çandır allows a confident correlation of Çandır with the local zone F (de Bruijn et al. 2013). The locality Çandır HW represents a test sample from a temporal outcrop near the original locality. The small mammal assemblage of Çandır HW consists of few isolated teeth of the rodents *Megacricetodon* sp., *Democricetodon* sp., *Crice-todon candirensis*, *Pliospalax* sp., and *Spermophilinus bredai*. The assemblage cannot be precisely biostratigraphically interpreted, though it suggests the same age as the original Çandır locality, correlating with the local zone F.

Bağıcı This locality has been mostly known in the literature under its former name “Zivra”. According to Saraç (1994) the fauna is characterized by its rich and diverse micromammal assemblage. Large mammals have also been reported by Sickenberg et al. (1975). Both small and large

mammals correspond stratigraphically to the local zone H (MN 7/8) (Saraç 2003; de Bruijn et al. 2013).

Süleymanlı Situated near the village of Süleymanlı, this locality consists of brown silty clays, rich in vertebrates and gastropod remains. The fauna is not yet described in detail, however, the age of the assemblage has been determined as Late Miocene (MN 13, local biozone L; see de Bruijn et al. 2013).

4 Systematic palaeontology

Squamata OPPEL, 1811

Anguimorpha FÜRBRINGER, 1900

Anguidae GRAY, 1825

Anguinae GRAY, 1825

Ophisaurus DAUDIN, 1803

4.1 *Ophisaurus* sp. (Fig. 2a–i)

Material, locality and horizon Right dentary UU KAR2 1102, Kargi 2, Oligocene/Miocene, MP 30/MN 1; dorsal vertebra UU KAR2 1101, Kargi 2, Oligocene/Miocene, MP 30/MN 1; dorsal vertebra UU KE 5101, Keseköy, Lower Miocene, MN 3; dorsal vertebra UU CD 5101, Çandır, Middle Miocene MN 6; dorsal vertebra UU BAG 5104, Bağıcı, late Middle Miocene, MN 7/8; dorsal vertebra EUNMH PV-14000, Süleymanlı, Upper Miocene, MN 13.

4.1.1 Description

Right dentary Specimen UU KAR 1102 represents a small fragment of the right dentary, preserving two tooth positions (Fig. 2a). Only one tooth is almost complete, with its apex, however, being damaged. The second, more anteriorly located tooth is broken off, preserving only its base. The subdental shelf (sensu Rage and Augé 2010) is thin. The preserved portion of the shelf has a straight course. Meckel's groove is fully open.

Dentition The dentition is pleurodont, with teeth being large and conical. They are pointed and only slightly curved posteriorly. The mesial and distal cutting edges are well developed in the apical section of the tooth height. The medial surface of the apex bears very faint striae which continue through the surface of the cutting edge.

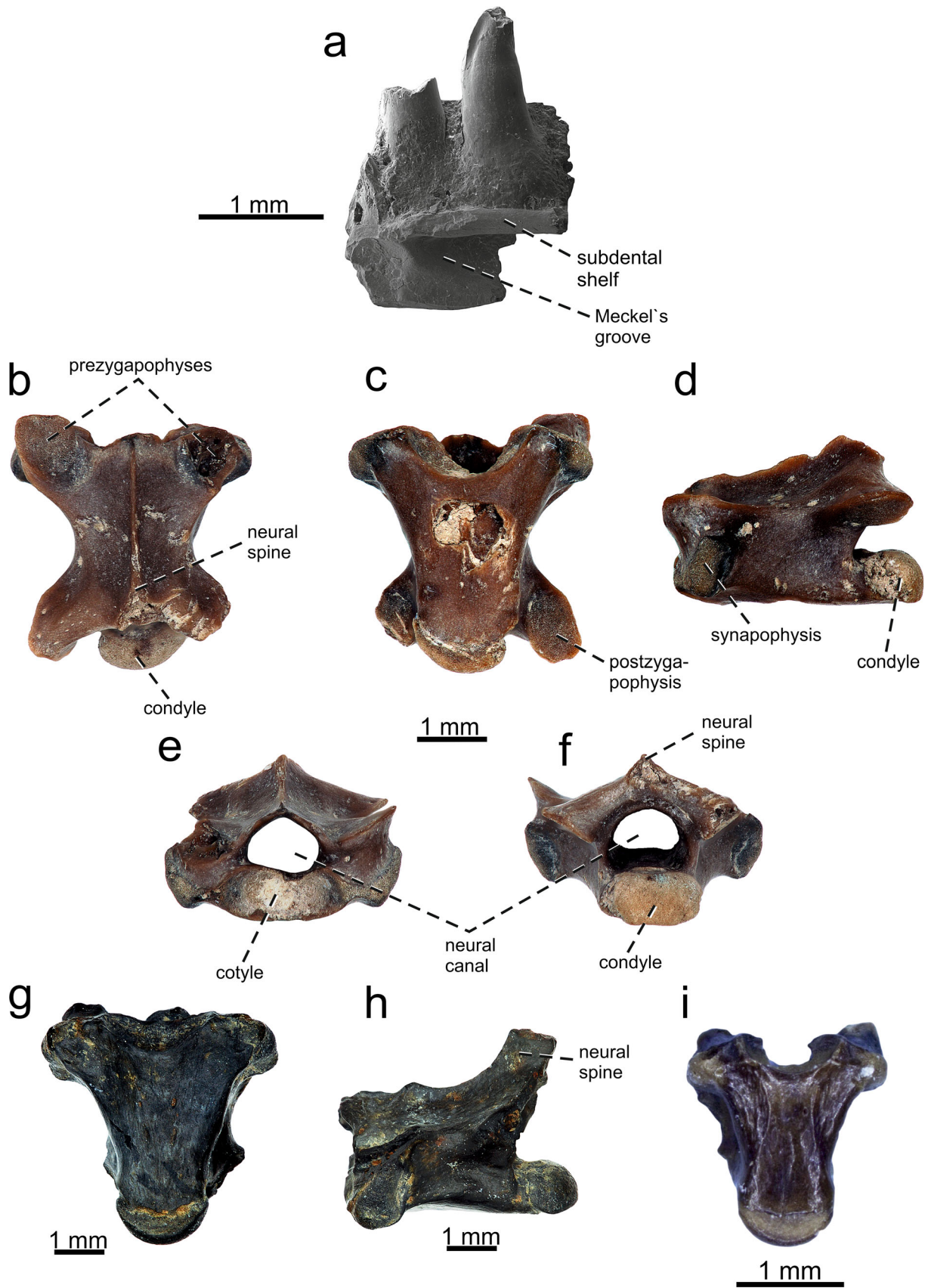
Dorsal vertebrae The description of the dorsal vertebrae is based mainly on the best preserved specimen, UU KE 5101, from Keseköy (Fig. 2b–f). The vertebra is markedly anteroposteriorly elongated relative to its height. In anterior view, the vertebra is approximately 1.5 times wider than high. The median ridge is gradually ascending posteriorly, forming a dorsally elevated neural spine in the posterior-most section. Unfortunately, the dorsal portion of

◀**Fig. 2** *Ophisaurus* sp. Right dentary from Kargi 2 (UU KAR2 1102) in **a** medial aspect; vertebrae from **b** to **f** Keseköy (UU KE 5101), **g**, **h** Kargi 2 (UU KAR2 1101) and **i** Süleymanlı (EUNMH PV-14000), in **b** dorsal; **c**, **g**, **i** ventral; **d**, **h** lateral, **e** anterior and **f** posterior aspects

the neural spine is broken off in the Keseköy specimen. In dorsal view, the ridge is thin and sharp, well defined along its entire length. It rather slightly widens posteriorly. The posterior-most section is distinctly wider than the anterior rest of the median ridge (in UU KAR2 1101, the neural spine forms a dorsally high structure, rectangular in lateral aspect, posteriorly inclined, ovoid in cross section, but frustratingly, its dorsal tip and the anterior portion are broken). The neural canal is moderately sized and roughly pentagonal in shape. The pre- and postzygapophyses are well expanded laterally. A strong interzygapophyseal constriction is developed between them, forming a deep lateral notch on both sides in dorsal view. The prezygapophyses bear roughly square articulation surfaces. They are inclined from the horizontal plane in an angle of approximately 36°. The synapophyses are rectangular in shape. The cotyle is not completely preserved in UU KE 5101, but it is well preserved in UU KAR2 1101 (Fig. 2g, h). In all specimens, the cotyle is strongly depressed and its height is slightly larger than the height of the neural canal. The condyle is elliptical and also depressed. The ventral surface of the centrum is flat. In UU KE 5101, the ventral surface of the centrum is partially damaged. However, this portion is well preserved in EUNMH PV-14000 or UU KAR2 1101. In EUNMH PV-14000 (Fig. 2i), the ventral surface of vertebra possesses two asymmetrically located subcentral foramina. In this specimen, almost parallel edges, limiting the flat ventral portion of the centrum, run from the lateral limits of the cotyle posteriorly to the condyle, and appear to be sharper and more strongly developed than those in the Keseköy and Kargi 2 specimens. The lateral margins (subcentral ridges) of the centrum are concave and the centrum gradually narrows posteriorly.

4.1.2 Remarks

Among extant anguine genera, the tooth morphology of UU KAR 1102 resembles that of *Ophisaurus*, specifically *O. koellikeri*, *O. harti*, *O. ventralis* and *O. attenuatus*. However, in all these extant species the striations are much more distinctly developed (Klembara et al. 2014). Only in small specimens of *Ophisaurus*, the striae on the apices of teeth are very faint (e.g., *O. acuminatus*; Klembara et al. 2014: Fig. 13E–G). According to its size, UU KAR2 1102 specimen corresponds to adult specimens of extant *Ophisaurus* (e.g., *O. koellikeri*). However, the morphology of



the tooth of UU KAR2 1102 specimen is very similar to the one recently described from the late Middle Miocene of Kazakhstan (Vasilyan et al. 2017).

A morphology of the teeth similar to that of UU KAR 1102 is also present in *Ophisauromimus* from the European Oligocene (Čerňanský et al. 2016). However, in well preserved teeth of *Ophisauromimus*, it is clearly visible that the cutting edges of the teeth are much longer in comparison to the length of the tooth and the cutting edges extend to the tooth base (Čerňanský et al. 2016: Fig. 14G, G1).

The concave lateral walls of the centrum of the dorsal vertebrae allow their taxonomic assignment to *Ophisaurus*. In contrast, straight lateral walls on the centrum are present in the dorsal vertebrae of adult specimens of *Pseudopus* (Klembara 1981: Fig. 3). Additionally, in *Pseudopus* the cotyle is significantly dorsoventrally higher than the height of the neural canal (see Čerňanský et al. 2016).

Anguis LINNAEUS, 1758.

4.2 *Anguis* sp. (Fig. 3)

Material, locality and horizon Dorsal vertebra UU BAG 5101, Bağıci, late Middle Miocene, MN 7/8; dorsal vertebra EUNMH PV-14001, Süleymanlı, Upper Miocene, MN 13.

4.2.1 Description

Dorsal vertebra The description is based on the well preserved specimen UU BAG 5101 (Fig. 3a–e). The vertebral centrum is anteroposteriorly elongated. The height of the vertebra slightly gradually increases posteriorly. The neural spine is markedly low. It forms a ridge running along almost the entire length of the dorsal section of the neural arch. In dorsal aspect, the ridge is thin, becoming less distinct in the anterior section. In the posterior two thirds of its length, it is well defined and widens at its posterior end.

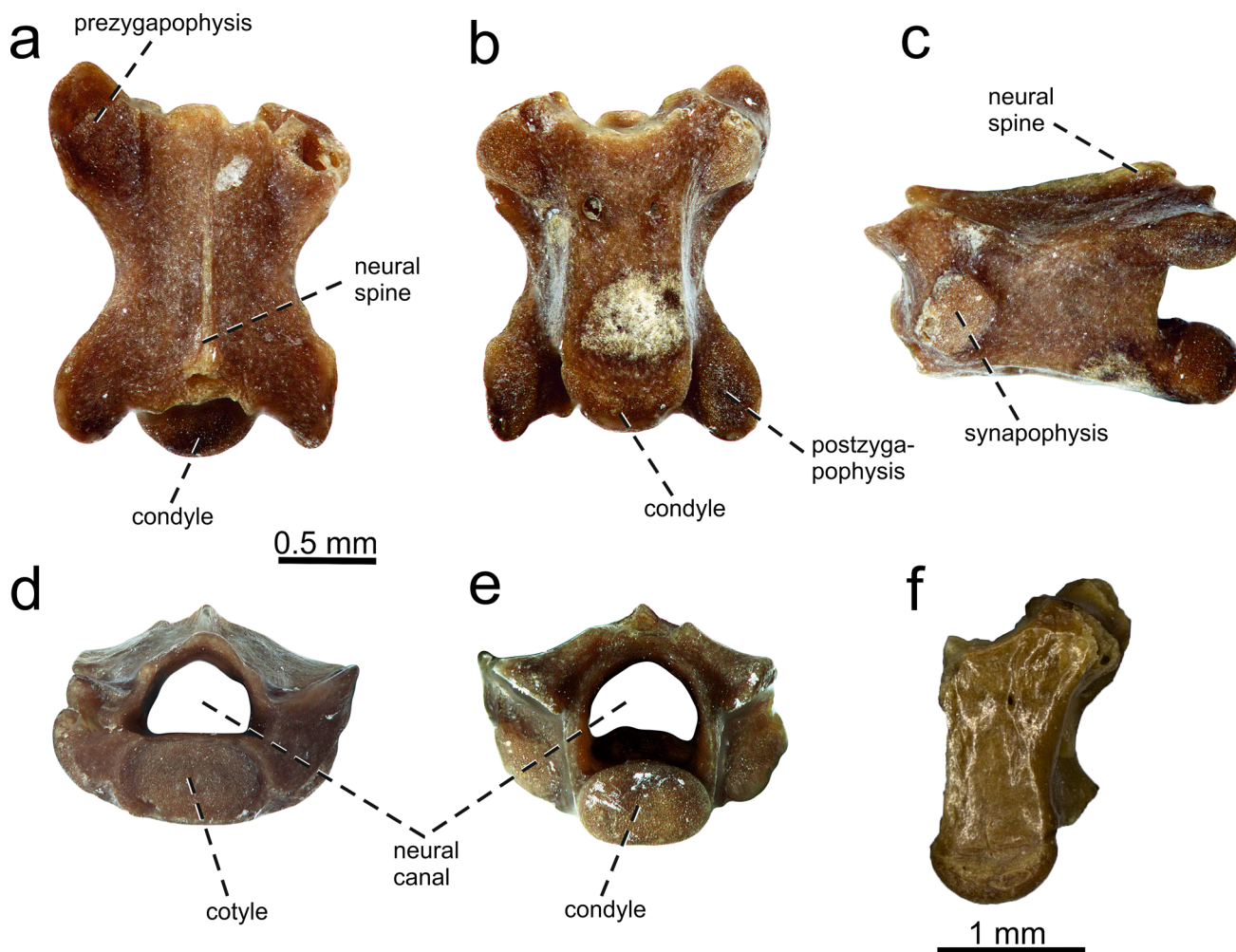


Fig. 3 *Anguis* sp. Vertebrae from a to e Bağıci (UU BAG 5101) and f Süleymanlı (EUNMH PV-14001), in a dorsal; b, f ventral; c lateral; d anterior; and e posterior aspects

The neural canal is large and roughly triangular in shape. The pre- and postzygapophyses are well expanded laterally, bearing elliptical articulation surfaces. A well developed interzygapophyseal constriction is located between them. The prezygapophyses are inclined from the horizontal plane in an angle of approximately 45°. The synapophyses are square in shape. The cotyle is large and its height is slightly bigger than the height of the neural canal (Fig. 3d), a typical diagnostic feature of *Anguis* and *Pseudopus*. The cotyle is depressed, as well as the condyle, which is markedly protruded posteriorly. The ventral surface of the centrum is flat, possessing two subcentral foramina in the anterior one third of the centrum. The lateral margins of the centrum run almost parallel, but the centrum is constricted at the level of the subcentral foramina.

4.2.2 Remarks

These vertebrae can be attributed to *Anguis* based on the course of the lateral margin of the vertebral centrum in ventral aspect. In extant *Anguis*, the lateral margins are almost parallel in ventral view, with a slight constriction immediately posterior to the subcentral foramina (Klembara 1981, Fig. 3).

4.3 Anguinae indet. (Fig. 4)

Material, locality and horizon Caudal vertebra UU KE 5102, Keseköy, Lower Miocene, MN3; osteoderm EUNMH PV-14003 Kocayarma, Oligocene MP 23–MP 27 or ? MP 25; osteoderm EUNMH PV-14021, Kavakdere, Oligocene MP 26/27; osteoderm EUNMH PV-14021, Kılçak 3b, Lower Miocene, MN 1; osteoderm EUNMH PV-14004, Sabuncubeli, Lower Miocene, MN 3; osteoderm UU KE 5103, Keseköy, Lower Miocene, MN 3; osteoderm UU KAR2 1103, Kargı 2, Oligocene/Miocene, MP 30/MN 1; osteoderms UU KAR3 1101 and UU KAR3 1102, Kargı 3, Lower Miocene, MN 1; osteoderms UU CD 5102, 5103, and 5104, Çandır, Middle Miocene, MN 6; osteoderm EUNMH PV-14005-12, Çandır HW, Middle Miocene, MN 6; osteoderms UU BAG 5102 and 5103, Bağıcı, late Middle Miocene, MN 7/8; osteoderm EUNMH PV-14013-20, Süleymanlı, Upper Miocene, MN 13.

4.3.1 Description

Caudal vertebra The caudal vertebra (KE 5102) is narrow, anteroposteriorly elongated and roughly cylindrical (Fig. 4a–d). The postzygapophysis is small, posteriorly directed. Only the right one is preserved. The neural canal is large, pentagonal in shape. In lateral aspect, the condyle protrudes well posteriorly. In posterior view, it is markedly

depressed. The haemapophyses are fused to the centrum, occupying its posterior region, right anterior to the condyle. Only the bases of the haemapophyses are preserved. The anterolaterally oriented transverse processes are robust, but most of their length is broken. The neural spine and the entire posterodorsal region of the vertebra are broken.

Osteoderms The osteoderms are of several morphotypes. The first morphotype is represented by slender, rectangular osteoderms, with a low medial ridge running along their entire central regions, or sometimes being restricted to the sculpture region (e.g. EUNMH PV-14003, Fig. 4e). In some osteoderms of the same morphotype, the medial ridge runs obliquely (this reflects a body location) (EUNMH PV-14013, Fig. 4j). The anterior overlap surface occupies almost one third of the external surface, and the lateral bevel is highest close to the overlap surface. The posterior portion of the external surface is ornamented. The ornamentation is formed by short grooves, tubercles, pits and ridges diverging from the central region. The central part of the internal surface is pierced by three foramina. Shallow grooves are located on the periphery of the internal surface.

Osteoderms of the second morphotype are roughly square in shape and possess no medial ridge (Fig. 4k). In several osteoderms, the lateral bevel is present on both sides.

The third morphotype consists of irregular, or roughly rhomboidal-shaped osteoderms (Fig. 4l). The medial ridge is absent.

5 Discussion

The specimens of *Ophisaurus* and *Anguis* described herein represent the first unambiguous evidence that anguines used the region of Anatolia as a passage on their dispersal route from Europe to East Asia. Although there is only one small fragment of dentary with only one tooth in our material, the morphology of this tooth is extremely similar with that of *Ophisaurus* sp. described recently from the Middle Miocene of Kazakhstan (Vasilyan et al. 2017).

The material from the localities of Kocayarma (Oligocene, MP 23–MP 27 or ? MP 25) and Kavakdere (MP 26/27) represents the oldest known occurrences of anguid lizards not only from Turkey, but from the whole south-eastern Europe. The presence of anguines in the Oligocene of Turkey is consistent with the closing of the Turgai Strait and the establishment of the terrestrial connection between Europe and Asia that opened a dispersal corridor for land-dwelling animals (Haq et al. 1987; Rögl 1999). Many lizard lineages appeared in Asia during the Oligocene, among them e.g. lacertids (Böhme 2007), whereas several Asian newcomers appeared for the first time in Europe (Rage 2013). The anguine material described from these

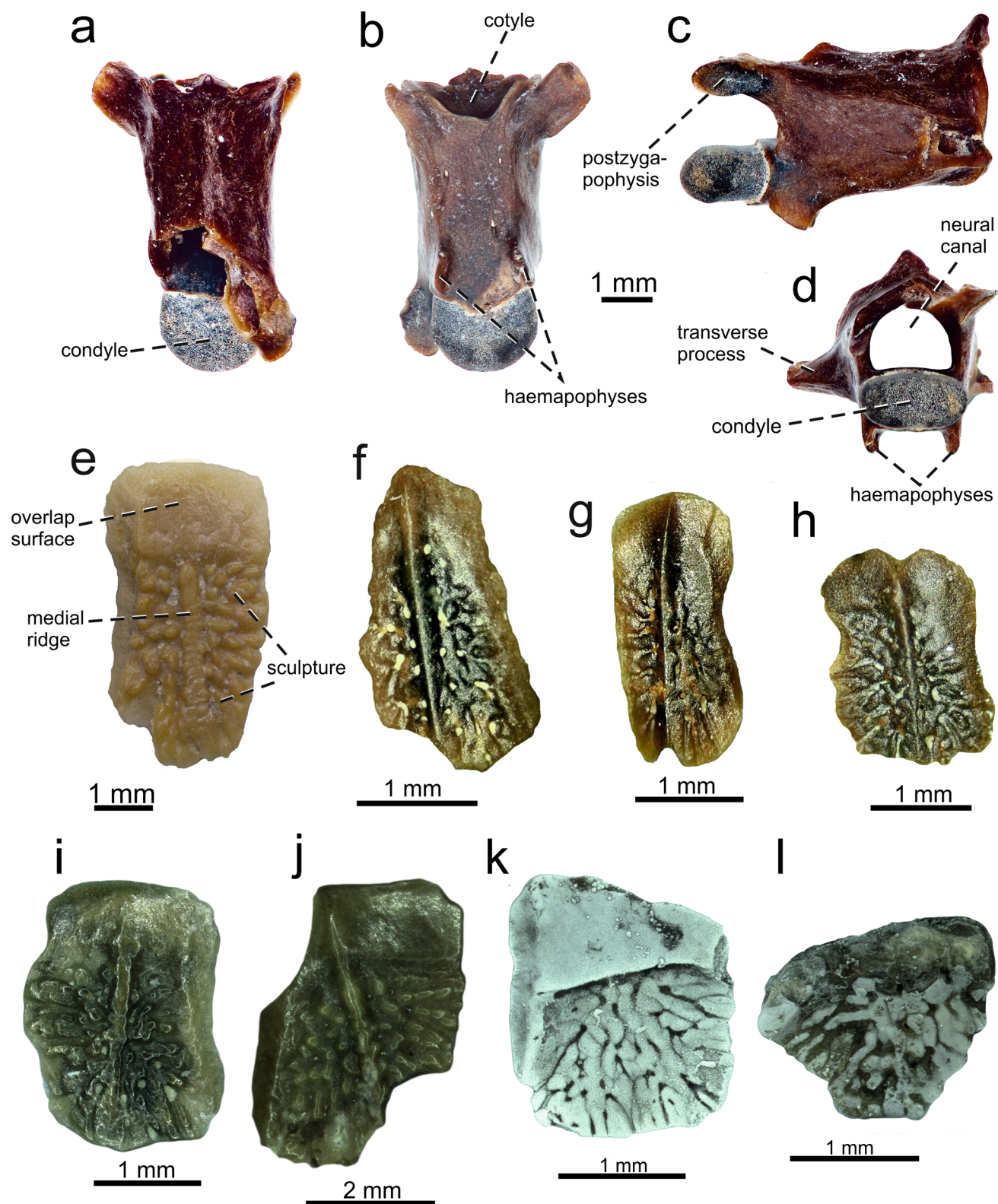


Fig. 4 Anguinae indet. Caudal vertebra from Keseköy (UU KE 5102) in **a** dorsal; **b** ventral; **c** lateral; and **d** posterior aspects. Body osteoderms from **e** Kocayarma (EUNMH PV-14003); **f** Kavakdere

(EUNMH PV-14021); **g** Kılçak 3b (EUNMH PV-14021), **h** Sabuncubeli EUNMH PV-14004, **i** Candır HW (EUNMH PV-14005) and **j**–**l** Süleymanlı (EUNMH PV-14013–15) in external aspects

two Oligocene localities of European Turkey consists only of isolated osteoderms, so its determination to the alpha taxonomy level is currently impossible. Nevertheless, they resemble osteoderms of *Ophisaurus*, but other taxa cannot be excluded (osteoderms of *Ophisauromimus* are unknown; Čerňanský et al. 2016). Osteoderms similar to that from Kocayarma (although with a less pronounced sculpturing pattern) have also been described from the European Oligocene, e.g. from Germany (see Čerňanský et al. 2016). In the locality of Kargı 2 (Oligocene/Miocene, MP 30/MN 1), the allocation of the material to *Ophisaurus* is without any doubt. Based on this fossil record, *Ophisaurus* appeared in the Asiatic part of Turkey during the terminal Oligocene (and perhaps even slightly earlier). The Oligocene and Miocene rhinocerotoids and associated faunas from several Turkish localities (Antoine et al. 2008; Sen et al. 2011), including the snake *Bavarioboa* (Szyndlar and Hoşgör 2012), indicate that Anatolia had close terrestrial connections with both Asia and Europe during the late Oligocene and Early Miocene. Therefore, the anguine material described herein from the Asiatic part of Turkey brings additional evidence for links between the terrestrial faunas of Europe and southwestern Asia.

In contrast to *Ophisaurus*, *Anguis* and *Pseudopus* still inhabit Turkey; see e.g. Baran et al. 1988). *Pseudopus* was not found in our collection of anguines from the available Turkish localities, so very little can be said about its first appearance in the region. Moreover, *Anguis* seems to have appeared much later in this region, as the oldest fossil record is herein shown to occur in the terminal Middle Miocene (Bağıcı). An hypothetical explanation for this might be the possible absence of suitable environments preferred by this taxon in Turkey (or along its dispersal route) till the end of the Middle Miocene. Members of *Anguis* are restricted to temperate and humid habitats, such as woodlands, grassland and heathland. Usually *Anguis* spp. live up to 1000–1500 m above sea level, and avoid open, unshielded areas (see e.g., Gvoždík and Moravec 2015). However, it is worth noting that the fossil finds of *Anguis* are in general extremely rare (see Klembara and Rummel 2016) so an earlier date of dispersal cannot be fully excluded.

Moreover, the Turkish area is at the crossing of Eurasia and Africa; Northern Africa is part of the Palearctic region as is Eurasia. It should be noted that in Northern Africa, the oldest known anguine lizards are found in the Middle Miocene of Beni Mellal (Morocco; see Rage 1976). *Ophisaurus* probably migrated to North Africa during the Miocene, as supported by material from the Late Miocene deposits of Morocco (Blain et al. 2013). According to Blain et al. (2013), *Ophisaurus* may have entered North Africa during or prior to the Middle Miocene via the land bridge between the Afro-Arabian Plate and Eurasia established

from the Early Miocene onwards, following a route that is the opposite from the direction of other squamate clades, such as chamaeleonids and cordylids (Čerňanský 2010, 2012; Georgalis et al. 2016b).

6 Conclusion

Fossil anguine lizard finds from several Turkish Oligocene and Miocene localities are described here for the first time. These rare records consist of dentary fragment, vertebrae and osteoderms. These skeletal and integumentary elements provide important information about the dispersal routes of anguines from Europe to Asia. The Kocayarma material represents the oldest occurrences of anguine lizards not only from Turkey, but also from the south-eastern Europe and the Eastern Mediterranean Basin. The fossils show that *Ophisaurus* appeared in the Asiatic part of Turkey during the Late Oligocene (and perhaps even slightly earlier), whereas the first appearance of *Anguis* seems to have happened much later in this region, in the late Middle Miocene.

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