



Effective source rock selection and oil–source correlation in the Western Slope of the northern Songliao Basin, China

He Bi¹ · Peng Li² · Yun Jiang¹ · Jing-Jing Fan¹ · Xiao-Yue Chen³

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Abstract

This study considers the Upper Cretaceous Qingshankou Formation, Yaojia Formation, and the first member of the Nenjiang Formation in the Western Slope of the northern Songliao Basin. Dark mudstone with high abundances of organic matter of Gulong and Qijia sags are considered to be significant source rocks in the study area. To evaluate their development characteristics, differences and effectiveness, geochemical parameters are analyzed. One-dimensional basin modeling and hydrocarbon evolution are also applied to discuss the effectiveness of source rocks. Through the biomarker characteristics, the source–source, oil–oil, and oil–source correlations are assessed and the sources of crude oils in different rock units are determined. Based on the results, Gulong and Qijia source rocks have different organic matter primarily derived from mixed sources and plankton, respectively. Gulong source rock has higher thermal evolution degree than Qijia source rock. The biomarker parameters of the source rocks are compared with 31 crude oil samples. The studied crude oils can be divided into two groups. The oil–source correlations show that group I oils from Qing II–III, Yao I, and Yao II–III members were probably derived from Gulong source rock and that only group II oils from Nen I member were derived from Qijia source rock.

Keywords Songliao basin · Western slope · Source rock distinction · Crude oils · Biomarker · Oil–source correlation

1 Introduction

As an important part of oil and gas geology research, the detailed study of source rock characteristics and oil–source correlation can provide direct guidance during oil and gas exploration and can assist with the further deployment of oil and gas exploration. Oil–source correlation is based on the concept that some components of oil and gas discharged from the source rocks are relatively similar to those of the oil and gas remaining in the source rocks (Peters and Moldowan 1993; Peters et al. 2005; Bennett et al. 2006). The source of oil and gas directly affects oil and gas accumulation and

reservoir distribution (Dow 1974; Schoell 1980; Tissot 1984; Magoon and Dow 1991). The detailed study of oil–source correlation can provide favorable information on the genesis and migration direction of oil and gas, thus leading to the discovery of new exploration fields (Tissot 1984). As one of the most important preliminary factors in oil–source correlation, the selection of source rocks and determining oil–oil correlations are indispensable steps (Dong et al. 2015; Mashhadi and Rabbani 2015; Brito et al. 2017).

The study of oil and gas sources can be performed with many available comparative indicators, including terpenes, steranes, *n*-alkanes, cycloalkanes, aromatic hydrocarbons and typical abiogenic heteroatomic compounds, such as thiophene and carbazole derivatives (Dong et al. 2015; Mashhadi and Rabbani 2015; Brito et al. 2017; Zhang et al. 2019). When an oil–source correlation is made, multifactor comparison is often adopted. Intuitively, the relationship between the source rock and crude oil can be judged by observing the degree of similarity in the biomarker "fingerprint" and by comparing the parameter ratios of light hydrocarbon compounds and biomarkers.

This study considers the Western Slope of the northern Songliao Basin, China, as an example to investigate the

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✉ He Bi
gloria_hebi@126.com

¹ Research Institute of Petroleum Exploration and Development, PetroChina, Beijing 100083, China

² Petroleum Exploration and Production Research Institute, SINOPEC, Beijing 100083, China

³ China National Oil and Gas Exploration and Development Co. Ltd., Beijing 100034, China

source rock characteristics and to differentiate the sources of the oils in different formations based on the analysis of oil–source correlation. To date, most research has focused on the Qijia-Gulong sag as a single unit to explore the source rock characteristics in this region, and little attention has been focused on the differences and effectiveness of the source rocks between Gulong and Qijia sags (Zou et al. 2004; Feng et al. 2005; Zhou et al. 2006; Meng et al. 2014). Thus, the oils of different formations in the Western Slope and their sources remain little attention to distinguish. Therefore, this study investigates the source rock characteristics between the Gulong and Qijia sags, determines the differences and effectiveness of the hydrocarbon supply, and discusses the source–source, oil–oil and oil–source correlations. This study is essential for understanding the sources of the discovered oils in different formations and for helping to explore the direction of oil and gas migration and significant exploration potential in the study area.

2 Geological setting

The Songliao Basin occupies an area of $2.6 \times 10^5 \text{ km}^2$, and six first-order tectonic zones are developed in the basin, namely the Western Slope, Northern pitching, Northeast uplift, Central depression, Southwest uplift, and Southeast uplift. To better investigate the sources of crude oils in the Western Slope tectonic zone, the eastern edge of the study area extends to the axis of Central depression, which contains the main source rocks. The study area covers an area

of $2.4 \times 10^4 \text{ km}^2$ and is divided into five tectonic zones: the Western overlap zone, Taikang uplift belt, Longhupao–Da’an terrace, Qijia sag and Gulong sag (Fig. 1). The oil resources in the study area are abundant and are distributed in multiple sets of strata (Meng et al. 2014).

Several sets of strata are drilled in the study area (Fig. 2). In this study, the Upper Cretaceous Qingshankou Formation (Qing I and Qing II–III members), Yaojia Formation (Yao I and Yao II–III members), and the first member of the Nenjiang Formation (Nen I member) are considered target strata based on the main distribution of the oil system in the study area. Source rocks are mainly present in the Qing I member and the Nen I member of Gulong and Qijia sags, and oil reservoirs are mainly present in the Qing II–III member, the Yaojia Formation and the Nen I member of Western Slope (Gao and Cai 1997; Fu et al. 2001; Kang et al. 2005; Zhou et al. 2006; Hou et al. 2009; Meng et al. 2014; Zhu et al. 2016; Bi et al. 2017).

3 Samples and methods

A detailed study was conducted based on the geochemical analyses of mudstone and oil-bearing sandstone samples. Total organic carbon (TOC), Rock–Eval pyrolysis, vitrinite reflectance, and *n*-alkanes data were obtained from the database of the Exploration and Development Research Institute of Daqing Oilfield, China. In addition, five mudstone cores (from Wells Y78, G572, G3 and D37, with depth of 1796–2322 m) were collected from the Gulong and Qijia

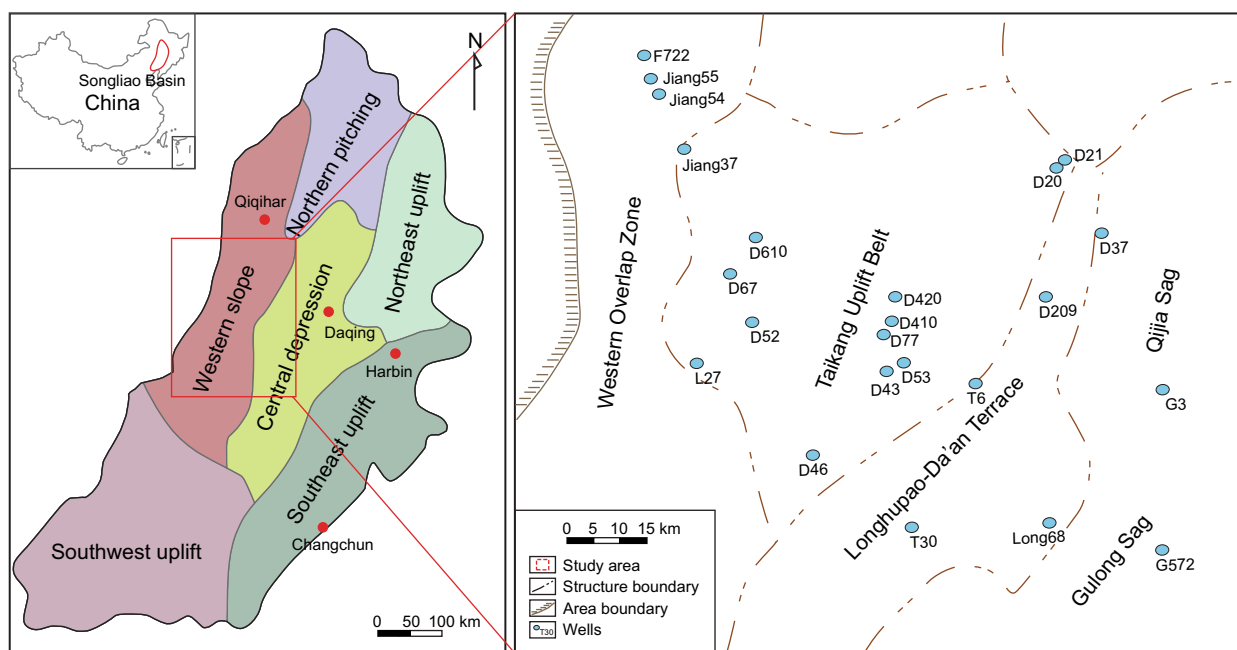


Fig. 1 Tectonic unit division, including the Western Slope of the Songliao Basin with location of well samples

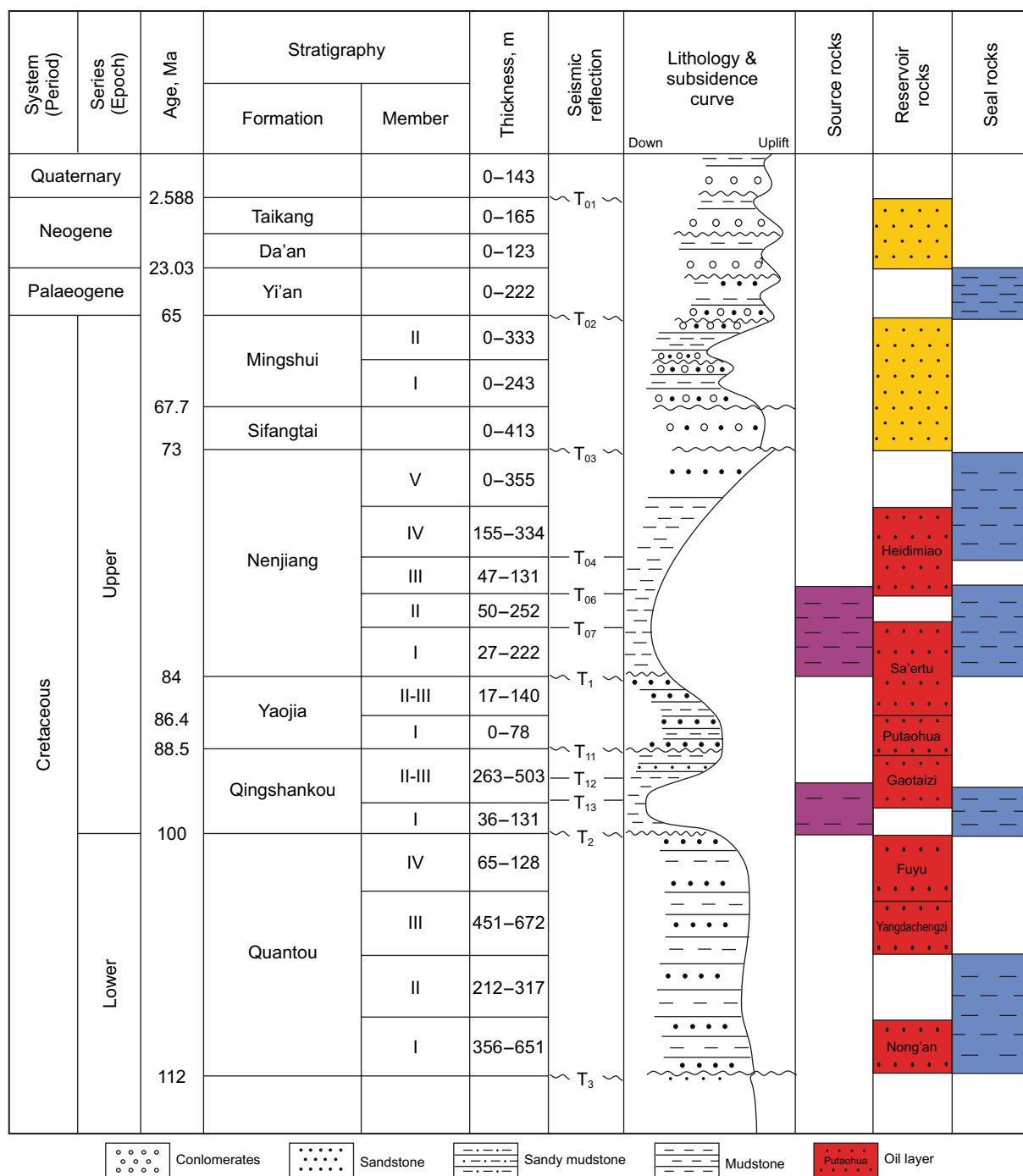


Fig. 2 Stratigraphic division and the major petroleum system elements of the study area in the Songliao Basin (after Bi et al. 2017)

sags, and 31 oil-bearing sandstone cores were collected from the Western Slope (Fig. 1). These core samples were subjected to bitumen extraction and gas chromatography–mass spectrometry (GC–MS) to define their molecular biomarker compositions. The analyses were carried out at Petroleum Geology Research and Laboratory Center, RIPED.

Asphaltenes were removed from oil-bearing sandstone samples by precipitation with *n*-hexane. Saturated and

aromatic hydrocarbons were separated by sequential elution with *n*-hexane and a solution of *n*-hexane and dichloromethane (2:1 v:v) on activated silica gel/alumina columns. The saturated fractions were analyzed using a Thermo-Trace GC Ultra-DSQ II GC–MS device. The GC uses an HP-5MS capillary column (60 m × 0.25 mm × 0.25 mm) with helium as the carrier gas (1.0 mL/min). The temperature of the GC oven was set to 100 °C for 5 min, increased to 220 °C at

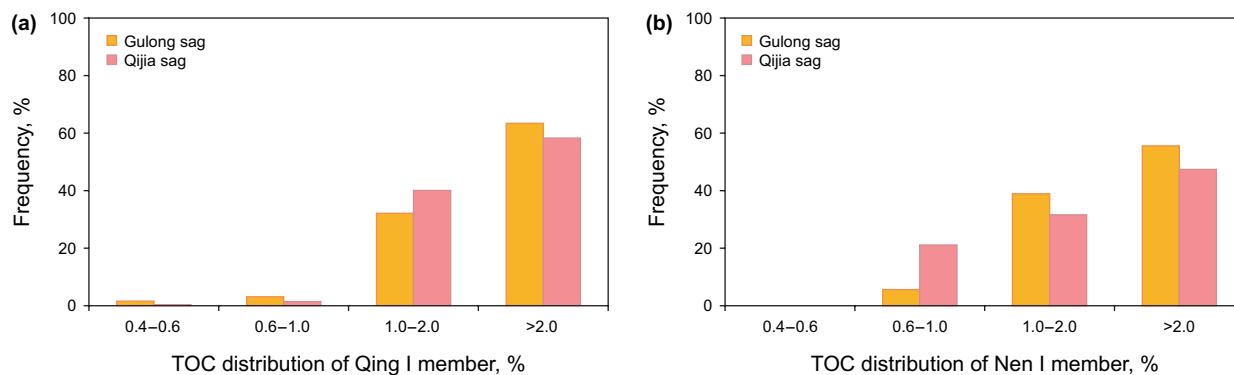


Fig. 3 Frequency distribution of the TOC values in the (a) Qing I and (b) Nen I source rocks in the Gulong and Qijia sags

4 °C/min and then to 320 °C at 2 °C/min, and finally held for 20 min. The mass spectrometer was operated in Electron Impact (EI) mode with an ionization energy of 70 eV and an ion source temperature of 280 °C. Based on selected ion monitoring, the mass-to-charge ratios of 177, 191, 217, 218 and 221 were determined. In this study, the fractions were analyzed using the following ion monitoring modes: m/z 191 and m/z 217.

4 Results and discussion

4.1 Source rock evaluation and generation potential

The evaluation of the Upper Cretaceous source rocks in the Gulong and Qijia sags is discussed based on the TOC, *n*-alkane, Rock–Eval pyrolysis, vitrinite reflectance and molecular analysis data. This evaluation assisted in selecting the effective source rocks and defining the differences among the source rocks in different tectonic units and formations.

4.1.1 Organic matter abundance

The TOC values of the source rocks range from 0.55% to 8.31%, with an average value of 2.41%, indicating fair-to-excellent petroleum potential (Diasty et al. 2019). The TOC values of the source rock in the Qing I member range from 0.55% to 8.31%, with an average value of 2.36%. The TOC values of the source rock in the Nen I member range from 0.72% to 8.06%, with an average value of 2.80%. According to Fig. 3, the TOC distributions of the source rocks in different tectonic units and formations are analyzed. The Qing I source rock has the highest TOC content. 97.5% of the Qing I samples have TOC > 1.0% and 59.6% of the samples have TOC > 2.0%, indicating good-to-excellent potential (Fig. 3a). The unit with the second highest TOC values is the Nen I source rock, in which 89.1% of samples have

TOC > 1.0% and 52.7% of the samples have TOC > 2.0% indicating good-to-excellent potential (Fig. 3b).

4.1.2 Organic matter type

Plots of Rock–Eval hydrogen index (HI) versus T_{\max} (Fig. 4) for Qing I and Nen I samples in the Gulong and Qijia sags show the organic matter type of the source rocks. HI values of Qing I samples in Gulong sag range from 71 to 455 mg/g TOC and those in Qijia sag range from 222 to 823 mg/g TOC, demonstrating that the Qing I source rock in the Gulong sag is primarily Type II–III kerogen and that in Qijia sag is mainly Type I–II kerogen (Fig. 4a). HI values of Nen I samples in Gulong sag range from 227 to 762 mg/g TOC and those in Qijia sag range from 318 to 799 mg/g TOC, demonstrating that the Nen I source rocks in the Gulong and Qijia sags are mixed Type I–II kerogen (Fig. 4b).

4.1.3 Depositional environment and provenance

Isoprenoids are commonly used to determine the sedimentary environment. Pristane (Pr) and phytane (Ph) are the most abundant and ubiquitous isoprenes and are often used to determine whether the original depositional environment was oxidizing or reducing. Oxidation preferentially promotes the conversion of phytols into pristane, while reduction promotes the conversion of dihydrophytol to phytane (Brooks et al. 1969; Ten Haven et al. 1987; Rowland 1990; Rontani et al. 2013; Schouten et al. 2013). The Pr/Ph ratios of the source rocks in the sags are generally 0.8–1.8, showing a suboxic condition (Fig. 5). In addition, relatively low Pr/*n*-C₁₇ and Ph/*n*-C₁₈ ratio values (Fig. 6) represent non-biodegraded hydrocarbons generated from the source rocks, which were deposited in a suboxic condition.

Sterane, as a biogenic indicator in sediments, has been widely used to obtain information about the sedimentary environment and the source properties (Peters et al. 2005). Sterane compounds are mostly derived from algae,

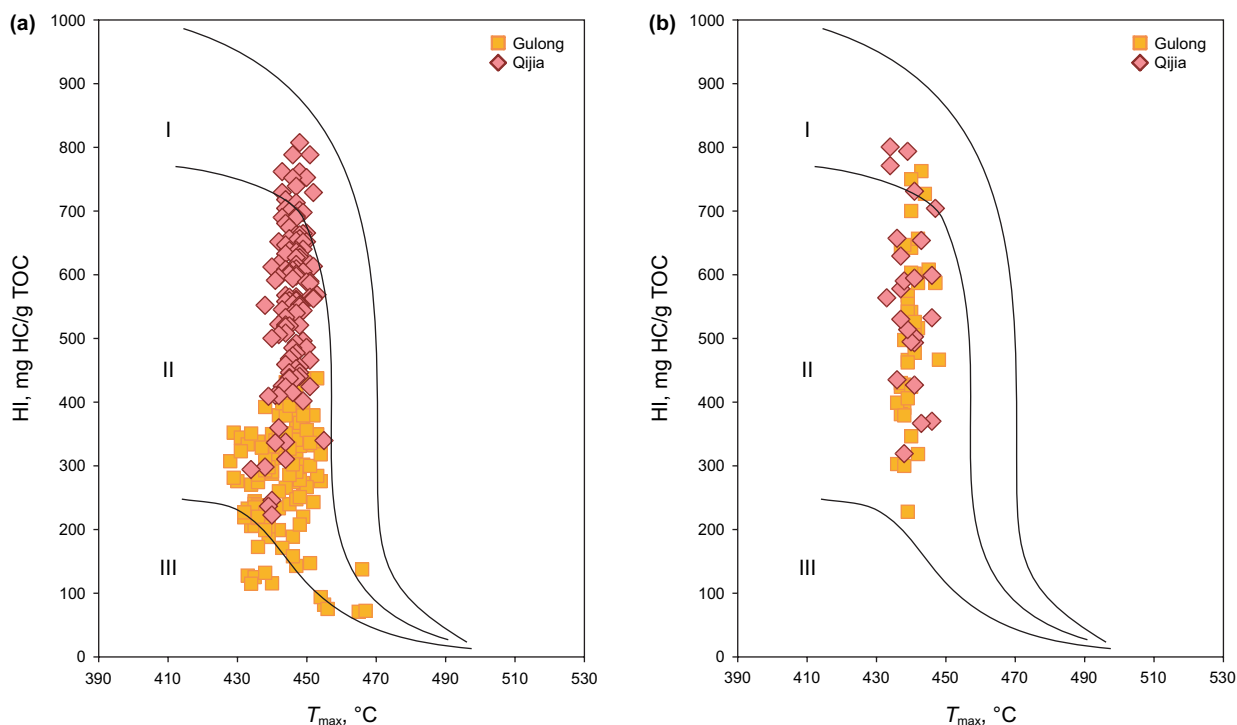


Fig. 4 Variation of Rock–Eval hydrogen index with T_{max} for samples from the (a) Qing I and (b) Nen I source rocks in the Gulong and Qijia sags, showing organic matter type

phytoplankton and higher plants. After deposition, the organic matter is converted from sterol to sterane by steroidal diene during heating (Peters et al. 2005). In general in source rocks and crude oils, the most common steranes include C_{27} – C_{29} steranes. It is generally recognized that C_{27} and C_{28} regular steranes originate from the lower aquatic organisms and algae, and C_{29} steranes are typically from terrestrial higher plants (Li et al. 2017). Therefore, a ternary diagram of 5α , 14α , 17α (20R) C_{27} , C_{28} , and C_{29} steranes can be used to identify the organic matter source (Cai et al. 2009). The relative distribution of C_{27} , C_{28} , and C_{29} steranes were plotted on a sterane ternary diagram (Fig. 7). In comparison with samples from Gulong and Qijia sags, samples from Gulong sag are toward C_{29} sterane and those from Qijia sag are toward C_{27} sterane. These show that the organic matter in the Gulong sag is primarily derived from mixed sources, while that in the Qijia sag is primarily derived from plankton (Fig. 7).

4.1.4 Thermal maturity

Vitrinite reflectance (R_o), as a maturity parameter, is commonly used to evaluate the degree of evolution. The R_o values of the source rocks in the study area range from

0.34% to 2.02%, with an average of 0.80%, which corresponds to the mature oil–generating stage. The R_o values of the Qing I member range from 0.44% to 2.02%, with an average of 0.94%. The R_o values of the Nen I member range from 0.34% to 1.18%, with an average of 0.58%. It can be clearly seen that the thermal maturity of the source rocks in the Qing I member is relatively high, while that in the Nen I member is low.

The planar R_o distribution maps of the study area show that the mature source rocks are widely distributed (Fig. 8). The rocks with a high thermal maturity of organic matter are distributed in the sags, where the R_o values are generally higher than 0.9%. The source rocks with the highest maturity are located in the Gulong sag, and the highest R_o is greater than 2.0%, followed by the source rocks in Qijia sag, with R_o values greater than 1.5%. From the depression to the slope, the maturity gradually decreases. In the west of the Taikang uplift belt and the Western overlap zone, the R_o values are less than 0.7%, and the organic matter is in the immature stage. In terms of planar distribution maps of each source rock, the mature source rocks of the Qing I member are widely distributed from the sags to the west of the Taikang uplift belt, while

the mature source rocks of the Nen I member are present only in sags and the Longhupao-Da’an terrace (Fig. 8).

In addition, this study also compares the thermal evolution degree of the source rocks in the two sags based on the $\sum C_{21}^-/\sum C_{22}^+$ n-alkane ratio. In the process of continuous thermal evolution of organic matter in source rocks, n-alkanes with high molecular weights gradually decrease in abundance, while those with low molecular weights increase in abundance. These results indicate that the $\sum C_{21}^-/\sum C_{22}^+$ n-alkane ratio has increased. According to the statistics of the $\sum C_{21}^-/\sum C_{22}^+$ ratios in each source rock, the ratios in the Gulong sag are obviously higher than those in the Qijia sag (Fig. 9). This indicates that the thermal evolution degree of the source rocks in the Gulong sag is higher than that in the Qijia sag.

4.1.5 Main source rock

The moderately to shallowly buried source rocks in the Songliao Basin generated and expelled hydrocarbons relatively late, within a narrow range. The source rocks began to expel hydrocarbons with 12% conversion rate when the R_o value reached 0.75%, and the conversion rate reached 70% when the R_o value reached 0.90% (Li 2019). Therefore, based on the planar R_o distribution, the area with R_o values greater than 0.75% represents the area with effective source rocks in the study area, and these areas are mainly within the sags. Some source rocks in the slope area have reached the mature stage but failed to exhibit effective hydrocarbon expulsion. The effective source rocks in the Qing I member have the largest distribution range, reaching the eastern edge of the Taikang uplift. The distribution ranges of the effective source rocks in the Nen I members are smaller and are limited to the Longhupao-Da’an terrace, the Qijia sag and the Gulong sag.

The moderately to shallowly buried strata of the northern Songliao Basin experienced two periods of large-scale uplift and denudation during later periods corresponding

to the deposition of the Nenjiang Formation and Mingshui Formation. These two periods of large-scale tectonic movement affected the hydrocarbon generation and expulsion of source rocks in the Gulong and Qijia sags (Zhou et al. 2006; Wen 2012). Two typical wells in the Gulong and Qijia sags (Wells G572 and G3, respectively) are selected to simulate the burial history and maturity evolution (pressure gradient is 1.03 MPa/100 m and thermal gradient is 4.36 °C/100 m) (Figs. 10 and 11). As shown in Fig. 10, the Qing I source rock in the Gulong sag reached the mature stage during the period of the first uplift, and the R_o value was greater than 0.75%, indicating that hydrocarbon generation and expulsion could have initiated. In contrast, the Nen I source rocks did not reach the mature stage during the period of the first

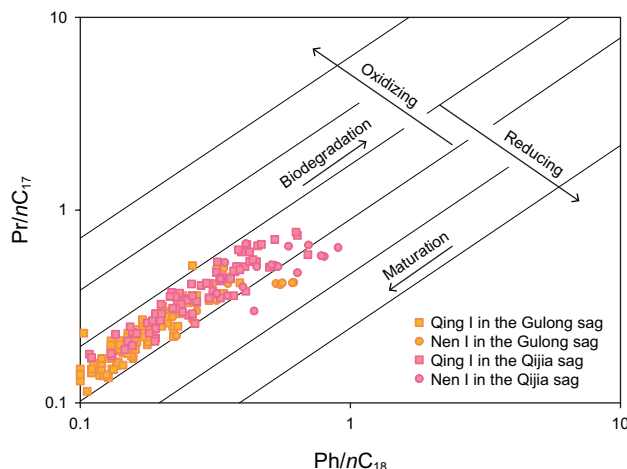


Fig. 6 Plot of Pr/nC_{17} versus Ph/nC_{18} showing maturation and depositional environment in the studied source samples

uplift and failed to generate and discharge hydrocarbon during that period. As shown in Fig. 11, the source rocks in the Qijia sag did not reach the hydrocarbon generation and

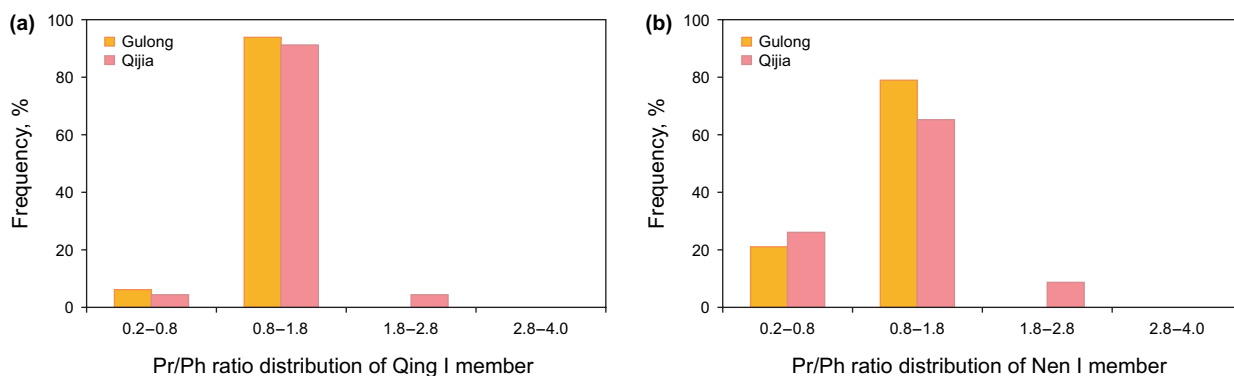


Fig. 5 Frequency distribution of Pr/Ph ratios in the (a) Qing I and (b) Nen I source rocks in the Gulong and Qijia sags

Slope of the northern Songliao Basin and represents the most important hydrocarbon generation and expulsion center.

4.1.6 Source–source correlation

The differences in provenance, maturity, and generation of the organic matter and in the expulsion evolution of the source rocks in the Gulong and Qijia sags are responsible for the differences in hydrocarbon generation. In addition, the comparison of the characteristics of biomarkers in the source rocks in the Gulong and Qijia sags can also intuitively reflect the differences between these areas. Therefore, tricyclic terpanes, hopanes, pregnane, and C₂₇, C₂₈, and C₂₉ regular steranes are selected for comparison of the biomarkers in the source rocks in the Gulong and Qijia sags.

According to the distribution characteristics of the terpanes, tricyclic terpanes are widely distributed in the extracts from the source rock and are generally dominated by C₁₉–C₃₀ terpanes. It can be seen from the m/z 191 mass chromatograms of the source rock samples from the two sags that the main differences between the source rocks in the two areas are associated with the tricyclic terpanes. The content of tricyclic terpanes of the Gulong source rocks is high (Fig. 14). The main peak corresponds to C₂₁ tricyclic terpane, showing an obvious predominance of C₂₁ tricyclic terpane. The Ts hopance contents are relatively higher than Tm hopance of Gulong source rocks, showing predominance of Ts hopane. Qijia source rocks have low contents of tricyclic terpanes and relative equilibrium between Ts and Tm hopanes contents (Fig. 15). The terpane distributions suggest that the source rocks in Gulong sag are more mature than those in Qijia sag.

The distribution characteristics of sterane compounds can be seen in the m/z 217 mass chromatogram. By comparing the m/z 217 mass chromatograms of the source rock samples from the two sags, it is found that the main differences

between the two sets of source rocks are in the concentrations of C₂₁ pregnane, C₂₇ regular sterane, C₂₈ regular sterane and C₂₉ regular sterane. According to the comparison of the sterane distributions of the two source rock regions (m/z 217), the content of C₂₁ pregnane in the Gulong source rocks is obviously high (Fig. 16). The C₂₉ regular sterane is relatively more abundant than the C₂₇ regular sterane and the C₂₈ sterane, showing an obvious predominance (Figs. 7 and 16). The C₂₁ pregnane content in the Qijia source rock is obviously low (Fig. 17). The content of the C₂₇ regular sterane is obviously higher than that of the C₂₉ regular sterane and the abundance of the C₂₈ sterane is the lowest, showing an obvious predominance of C₂₇ sterane (Figs. 7 and 17).

Therefore, based on the provenance, maturity, generation and expulsion evolution, and biomarkers distribution, the studied source rocks can be separated in two sets. One set is located in the Qing I member in Gulong sag. The other set is the Qing I source rock in Qijia sag.

4.2 Crude oil geochemistry

4.2.1 Physical properties

The density, viscosity and wax content of crude oils from Western Slope are given in Table 1. Great variations in physical properties of crude oils can be observed in the different rock units. Crude oils from Nen I member are characterized by high density, averaging 0.9020 g/cm³, high viscosity, averaging 121.8 mPa s, and low wax content, averaging 23.0%. Crude oils from Yao II–III member are similar by high density, averaging 0.9058 g/cm³, high viscosity, averaging 140.5 mPa s, and low wax content, averaging 22.0%. Due to the intense denudation, Yao I member mainly develops in the Longhupao-Da’an terrace. The short migration distance has less effects on crude oils, with lowest density (averaging 0.8459 g/cm³), lowest viscosity (averaging 35.1 mPa.s), and highest wax content (averaging 26.9%). Compared with oils from the

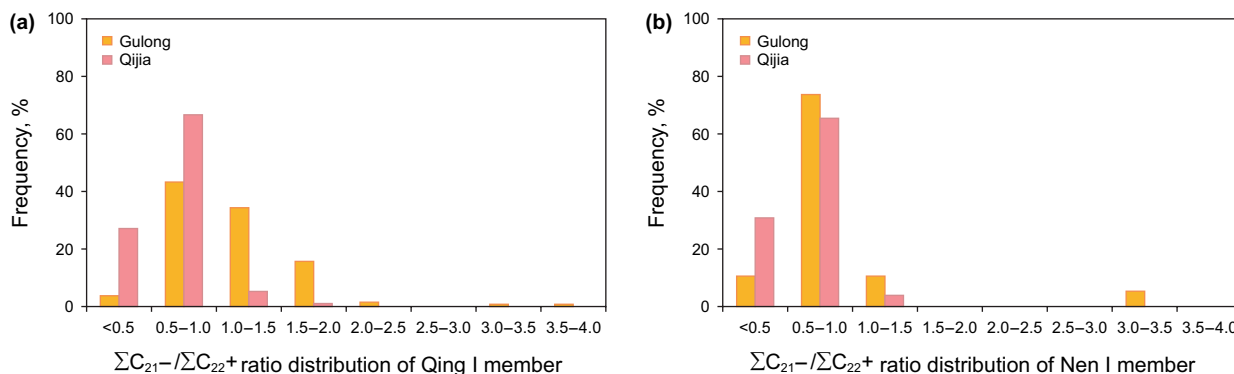


Fig. 9 Frequency distribution of the $\Sigma C_{21-}/\Sigma C_{22+}$ ratio in the (a) Qing I and (b) Nen I source rocks in the Gulong and Qijia sags

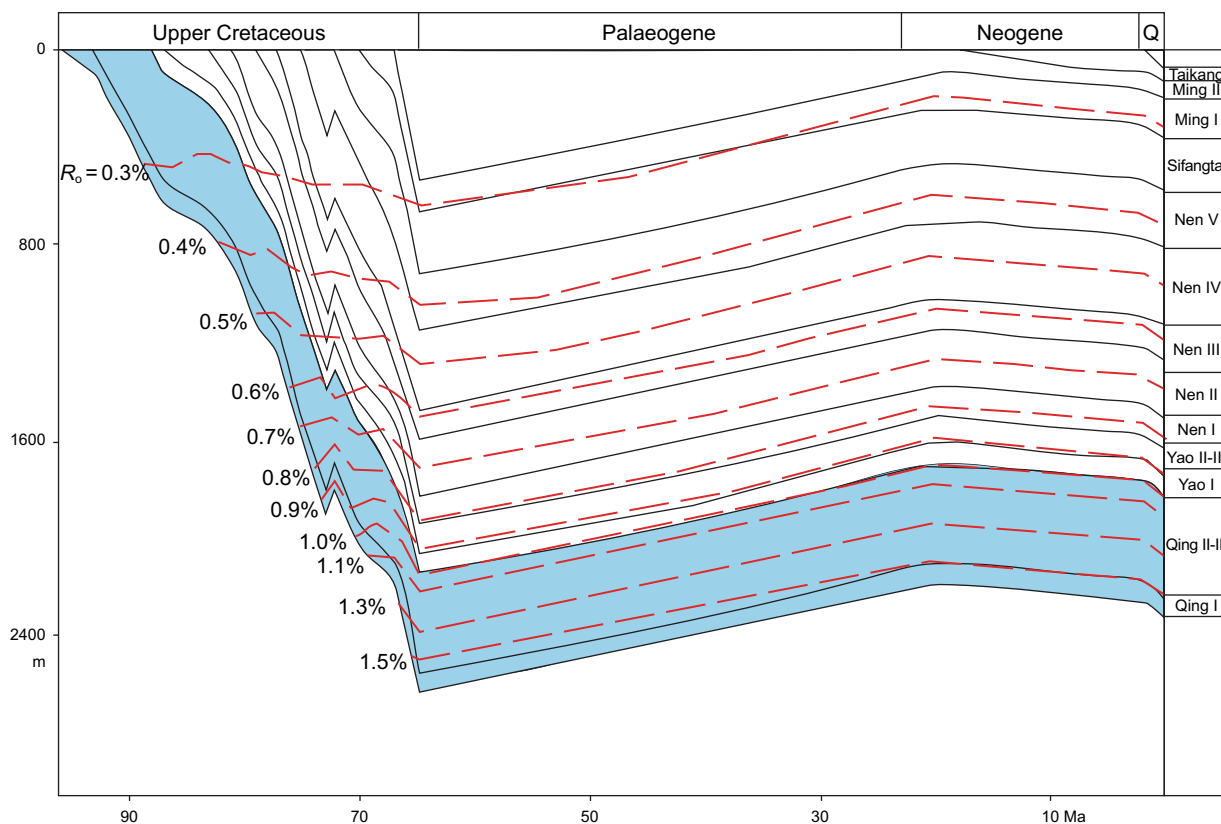


Fig. 10 Burial history and maturity evolution of well G572 in the Gulong sag

overlying Nen I and Yao II–III members, crude oils from Qing II–III member have the relatively lower density with average value of 0.8708 and lower viscosity with average value of 86.8.

4.2.2 Gross compositions

The gross compositions of the oils from the Western Slope are given in Table 2. Saturated hydrocarbon fractions for crude oils range from 47.88% to 81.40%, with an average of 63.34% (Table 2). Crude oils are also characterized by high fractions of resins and asphaltenes ranging from 9.26% to 38.66%, with an average of 24.42% (Table 2). In view of high density and viscosity, the crude oils might be affected by slight biodegradation.

4.2.3 Oil maturity

The maturity of crude oil can be judged by biomarker distributions (Mackenzie 1984). The ratios of $\alpha\alpha\alpha$ C₂₉ 20S/(20S + 20R) and C₂₉ $\beta\beta/(\alpha\alpha + \beta\beta)$ are commonly used indicators for estimation of maturity (Waples and Machihara 1990; Peters and Moldowan 1993; Chen et al. 1997; Peters et al. 2005; Hu 2012; Bai et al. 2013). Oils from Qing II–III, Yao I and Yao II–III member have $\alpha\alpha\alpha$ C₂₉ 20S/

(20S + 20R) sterane ratios ranging from 0.42 to 0.50 and C₂₉ $\beta\beta/(\alpha\alpha + \beta\beta)$ sterane ratios ranging from 0.43 to 0.52 (Table 3 and Fig. 18). Oils from Nen I member have relatively low $\alpha\alpha\alpha$ C₂₉ 20S/(20S + 20R) sterane ratios ranging from 0.37 to 0.44 and relatively low C₂₉ $\beta\beta/(\alpha\alpha + \beta\beta)$ sterane ratios ranging from 0.28 to 0.35 (Table 3 and Fig. 18). These results indicate that the crude oils from Western Slope are mature and that there are no immature and low-maturity oil mixes (Fig. 18).

More detailed information about the mature crude oils from different rock units can be provided by hopane parameters, which are sensitive to change in thermal maturity. $T_s/(T_s + T_m)$ and C₂₉ T_s/C_{29} hopane ratios are reliable maturity indicators increasing with maturity (Fowler and Brooks 1990; Farrimond et al. 1996; Peters et al. 2005; Cheng et al. 2013; Dong et al. 2015). As shown in Table 3, crude oils from Nen I member have low $T_s/(T_s + T_m)$ and C₂₉ T_s/C_{29} hopane ratios, ranging from 0.45–0.51 and 0.30–0.41, respectively. Crude oils from Yao II–III, Yao I and Qing II–III have high $T_s/(T_s + T_m)$ (0.57–0.66, 0.54–0.63 and 0.58–0.68, respectively) and C₂₉ T_s/C_{29} hopane ratios (0.61–1.10, 0.63–0.87 and 0.67–1.18, respectively). These indicate that maturity of Nen I oils is lower than that of oils from Yao II–III, Yao I and Qing II–III.

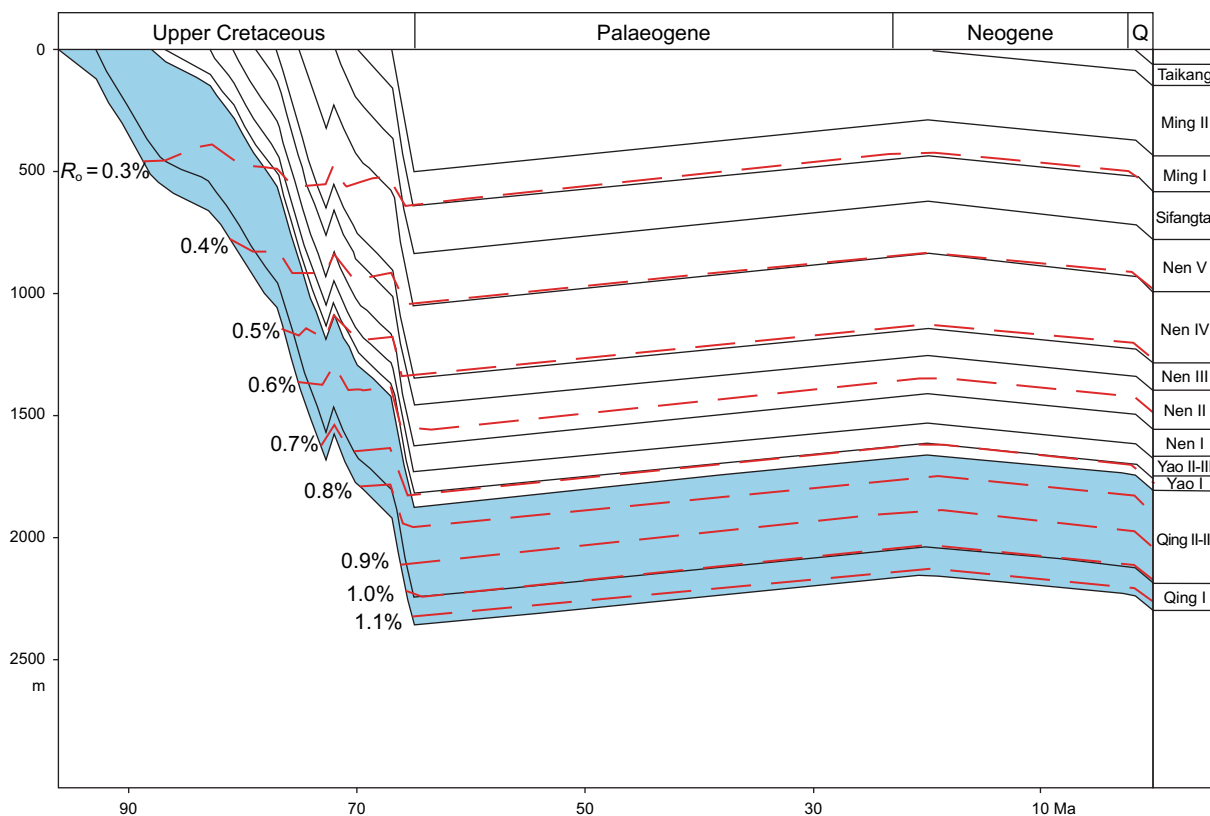


Fig. 11 Burial history and maturity evolution of well G3 in the Qijia sag

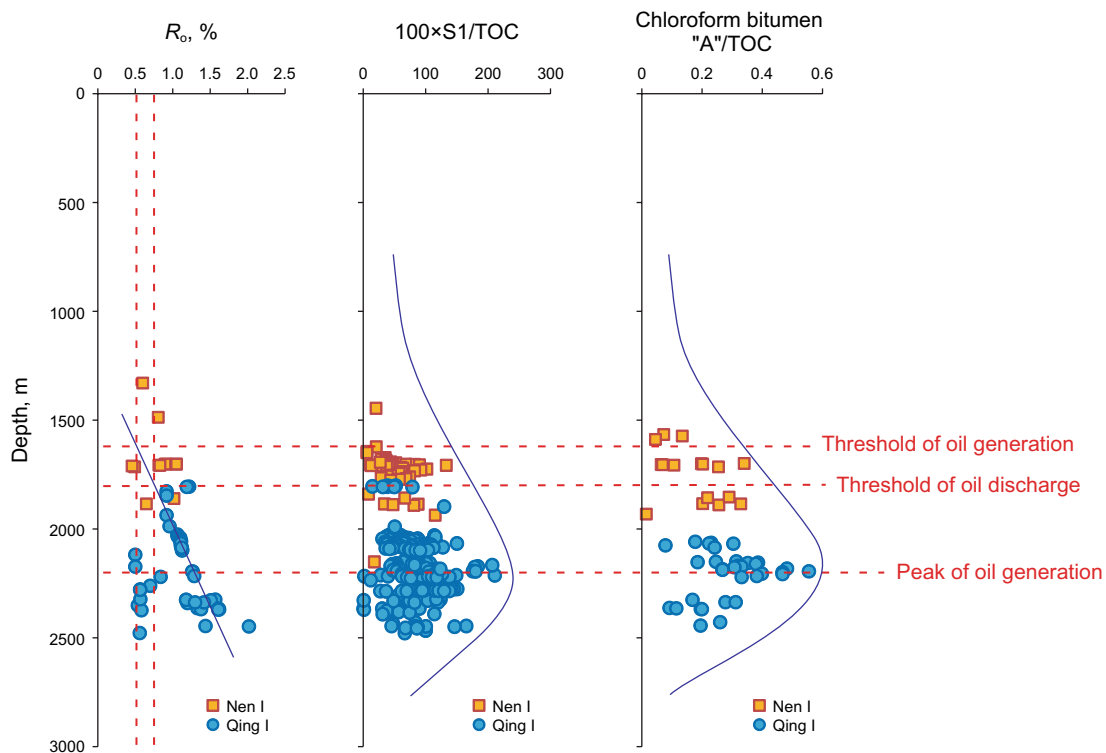


Fig. 12 Hydrocarbon evolution profiles of the source rocks in the Gulong sag

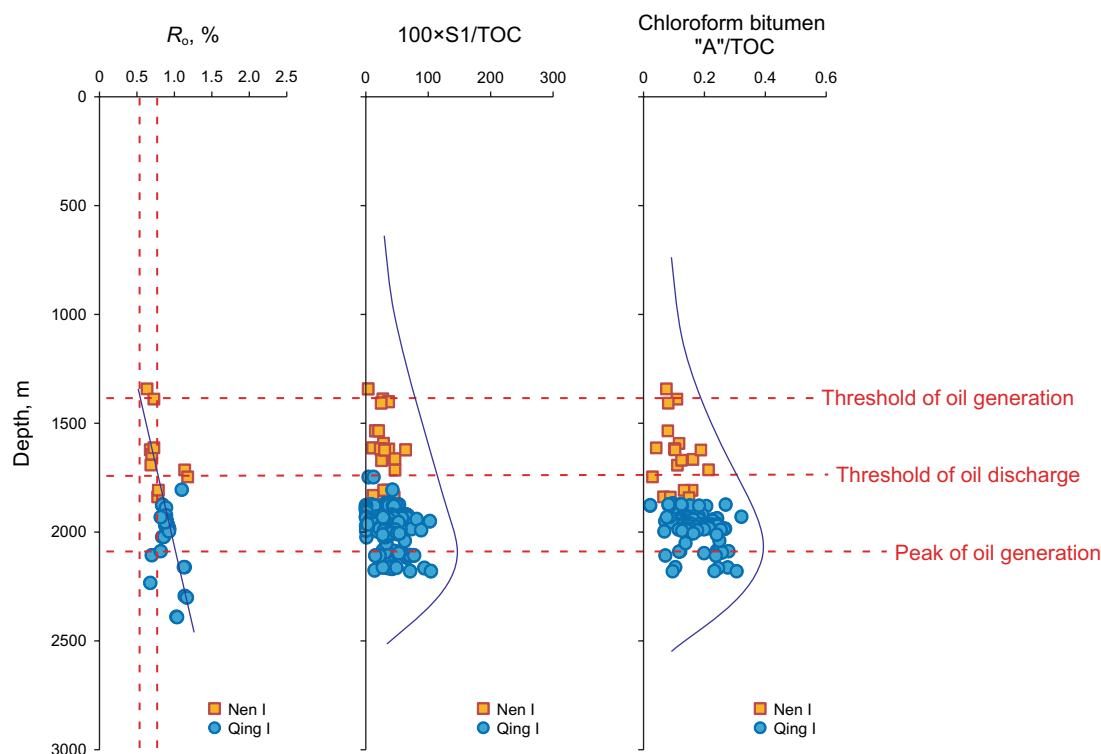


Fig. 13 Hydrocarbon evolution profiles of the source rocks in the Qijia sag

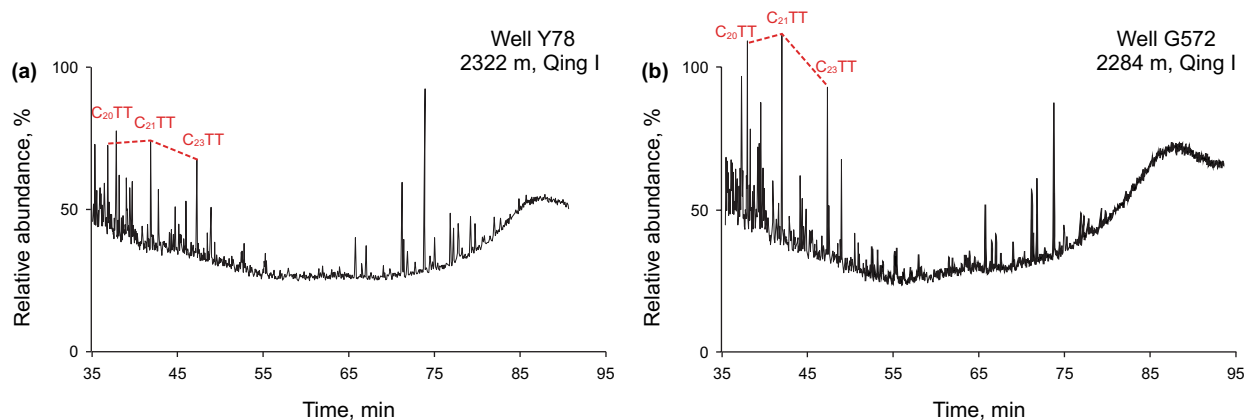


Fig. 14 Terpane distributions (m/z 191) in the source rocks in the Gulong sag

4.2.4 Oil–oil correlation

The distribution characteristics of tricyclic terpanes in the Qing II–III, Yao I and Yao II–III crude oils are significantly different from those in the Nen I crude oils (Fig. 19). The contents of tricyclic terpanes are high in the crude oils of Qing II–III, Yao I and Yao II–III members. The main peak corresponds to the C_{21} tricyclic terpane, showing the obvious predominance of the C_{21} tricyclic terpane. Qing II–III, Yao I and Yao II–III oils have relatively higher T_s hopane contents than T_m hopane contents, showing predominance

of T_s hopane (Fig. 19a, b, and c). Nen I oils have low contents of tricyclic terpanes and relative equilibrium between T_s and T_m hopanes contents (Fig. 19d). The terpane distributions suggest that crude oils in Qing II–III, Yao I and Yao II–III members are more mature than those in Nen I member.

The abundance of steranes in m/z 217 mass fragmentograms is also used to differentiate crude oils. The sterane distributions of the Qing II–III, Yao I and Yao II–III crude oils are similar and have relatively high C_{21} pregnane contents (Fig. 20a, b, and c). The C_{27} , C_{28} and C_{29} regular steranes are characterized by low C_{27} sterane and a predominance of

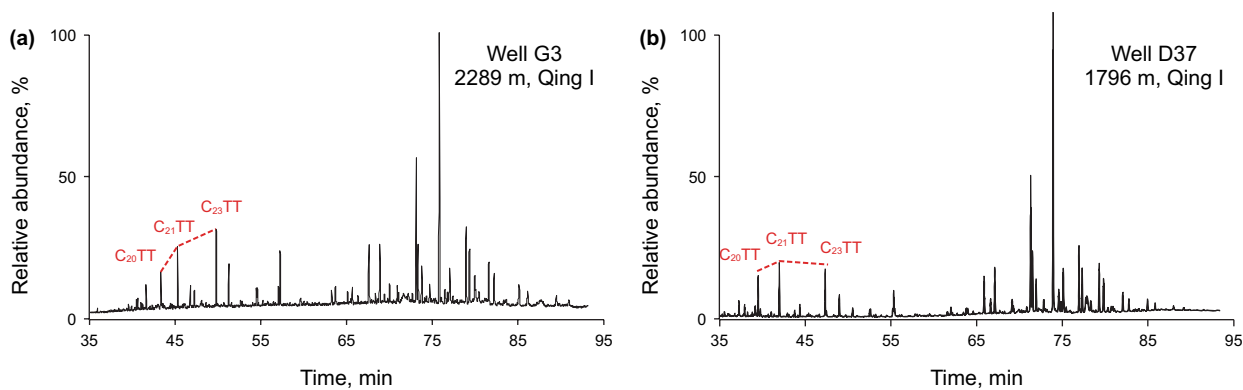


Fig. 15 Terpane distributions (m/z 191) in the source rocks in the Qijia sag

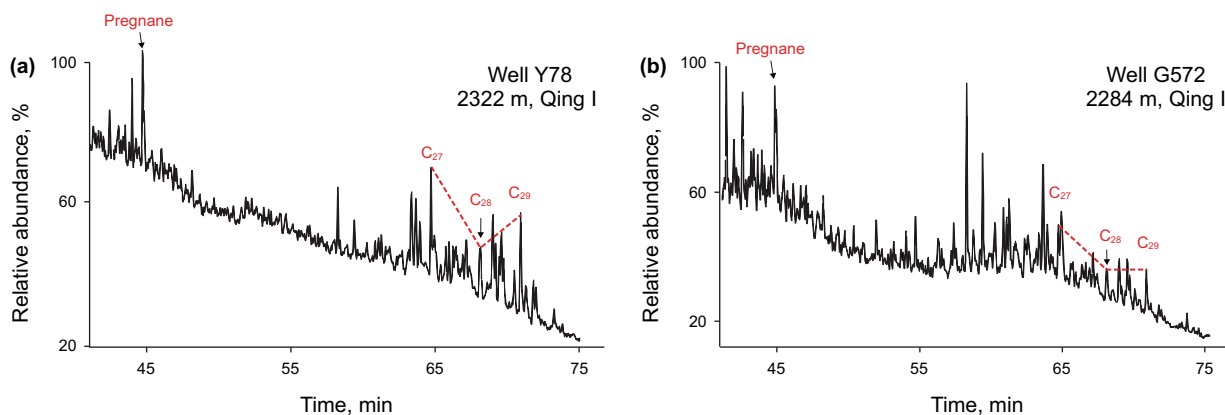


Fig. 16 Sterane distributions (m/z 217) in the source rocks in the Gulong sag

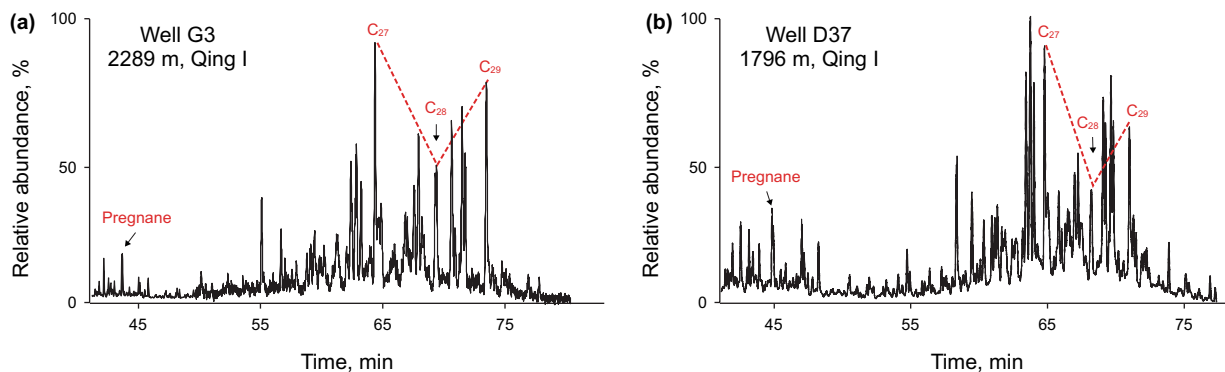


Fig. 17 Sterane distributions (m/z 217) in the source rocks in the Qijia sag

C_{29} sterane contents ($C_{29} > C_{28} > C_{27}$) (Table 3 and Fig. 21). The sterane distributions of the Nen I crude oils are different and features lower C_{21} pregnane contents (Fig. 20d). The C_{27} , C_{28} and C_{29} regular steranes are characterized by low C_{28} sterane and a predominance of C_{27} sterane contents ($C_{27} > C_{29} > C_{28}$) (Table 3 and Fig. 21). The different sterane

distributions of the two kinds of oils may be related to differences between their kerogen precursors.

Based on the physical properties, gross compositions, maturity and molecular compound, the crude oils are classified into two groups. Group I oils consist of oils from Qing II–III, Yao I and Yao II–III members. Group II oils consist of oils from Nen I member. The oils with systematic variations

Table 1 Statistics on physical properties of crude oils from Western Slope

Member	Density, g/cm ³	Viscosity, mPa.s	Wax, %
Nen I	0.8360–0.9382 0.9020(48)	6.7–424.5 121.8(26)	11.5–65.1 23.0(45)
Yao II–III	0.8319–0.9389 0.9058(76)	8.9–547.1 140.5(37)	5.4–43.5 22.0(78)
Yao I	0.8139–0.9072 0.8459(48)	6.9–113.6 35.1(28)	16.4–48.5 26.9(49)
Qing II–III	0.8211–0.9317 0.8708(93)	3.4–842.7 86.8(63)	10–35.3 23.9(88)
Minimum–maximum			
Average (sample number)			

suggest that oils might produce from the different source rock units.

4.3 Oil–source correlation

It is well-known that the organic matter in source rocks cracks into oil and gas under certain temperature and pressure conditions. Subsequently, some of these hydrocarbons undergo primary and secondary migrations to appropriate traps, forming oil and gas reservoirs, whereas the rest remains in the source rocks. Therefore, the extractable components in the source rocks are related to the oil and gas derived from the same source rocks and these two substances share some chemical composition similarities. The similarities and differences in chemical composition among oil and gas from the same and different source rocks are the fundamental basis of oil–source correlation. However, oil

Table 2 Gross composition of the crude oils from Western Slope

Well	Depth, m	Member	Saturates, %	Aromatics, %	Resins, %	Asphaltenes, %
D410	1036	Nen I	72.50	7.10	17.37	3.04
D43	1065	Nen I	64.54	15.69	19.51	0.26
D46	1045.75	Nen I	67.15	16.30	16.55	–
D52	810	Nen I	54.81	6.59	38.16	0.45
D53	1063.4	Nen I	53.43	14.29	32.28	–
D610	720	Nen I	59.50	7.29	33.21	–
D67	752	Nen I	68.42	8.90	21.29	1.40
D77	1027.3	Nen I	67.80	14.06	18.15	–
D20	1250.37	Yao II–III	63.44	12.09	24.46	–
D209	1272	Yao II–III	68.46	10.67	20.52	0.35
D43	1094.6	Yao II–III	63.19	14.00	22.82	–
D46	1070.37	Yao II–III	63.34	16.56	20.10	–
D77	1052.5	Yao II–III	64.88	12.64	21.09	1.39
F722	455	Yao II–III	47.88	13.69	35.55	2.88
Jiang54	484	Yao II–III	63.14	12.42	24.13	0.31
L27	676	Yao II–III	47.89	23.72	27.70	0.69
D209	1341.5	Yao I	66.03	6.23	27.74	–
D46	1085.57	Yao I	63.61	20.58	15.58	0.24
Long68	1762.98	Yao I	81.40	9.03	9.58	–
T30	1039	Yao I	80.58	10.15	9.26	–
T6	1245.85	Yao I	74.83	7.62	17.30	0.25
D20	1404	Qing II–III	56.18	15.43	26.98	1.41
D209	1626	Qing II–III	72.55	7.47	19.98	–
D21	1450	Qing II–III	61.69	12.22	26.09	–
D410	1115	Qing II–III	59.11	11.24	29.66	–
D420	1106	Qing II–III	60.48	20.08	19.44	–
D610	770	Qing II–III	49.32	12.02	38.66	–
Jiang37	596	Qing II–III	62.81	11.92	25.28	–
Jiang55	466	Qing II–III	49.49	12.04	35.27	3.20
L27	687	Qing II–III	58.27	14.17	27.56	–
T6	1300	Qing II–III	76.70	3.48	19.39	0.44

and gas can be contaminated in the process of migration and accumulation, which increases the difficulty of oil–source correlation to some extent. Therefore, we need to conduct a comprehensive comparative analysis of various parameters and select specific and stable parameters for oil–source correlation.

A successful oil–source rock correlation aims to find the genetic relationship between source rocks and oils and relies on biomarker composition (Peters and Moldowan 1993; Peters et al. 2005; Curiale 2008). As discussed earlier, based on the physical properties, chemical compositions, and biomarker fingerprints, the investigated oils from the Western Slope belong to group I and II. The Qing I member in the both Gulong and Qijia source rocks with different characteristics are considered be effective sources for the oils. Based

on the m/z 191 and m/z 217 mass chromatograms, the crude oil to source rock correlation is discussed in order to determine genetic relationships.

4.3.1 Group I oils

Terpane parameters that can be used to perform oil–source rock correlation. Group I oils have high abundance in tricyclic terpanes and show obvious predominance of the C_{21} tricyclic terpane. The T_s hopane contents are higher than T_m hopane, revealing predominance of T_s hopane. Additionally, $T_s/(T_s + T_m)$ and $C_{29} T_s/C_{29}$ hopane ratios for group I oils are relatively high, suggesting relatively high maturity. The Gulong source rock extracts have the similar characteristics of terpane parameters, indicating group I oils are related.

Table 3 Molecular parameters of crude oils

Well	Depth, m	Member	Group	$\alpha\alpha\alpha C_{29}$ 20S/ (20S + 20R)	$C_{29} \beta\beta/(\beta\beta + \alpha\alpha)$	$T_s/(T_s + T_m)$	$C_{29} T_s/$ C_{29} hopane	$C_{27}/\sum C_{27-29}$	$C_{28}/\sum C_{27-29}$	$C_{29}/\sum C_{27-29}$
D410	1036	Nen I	II	0.40	0.28	0.45	0.30	41.33	23.68	34.99
D43	1065	Nen I	II	0.43	0.33	0.49	0.39	37.76	23.8	38.44
D46	1045.75	Nen I	II	0.37	0.32	0.47	0.33	41.64	23.2	35.16
D52	810	Nen I	II	0.40	0.35	0.48	0.39	39.75	24.16	36.09
D53	1063.4	Nen I	II	0.44	0.34	0.47	0.32	39.63	24.82	35.55
D610	720	Nen I	II	0.38	0.31	0.50	0.39	42.45	24.19	33.36
D67	752	Nen I	II	0.38	0.31	0.51	0.41	41.4	24.22	34.38
D77	1027.3	Nen I	II	0.42	0.28	0.45	0.30	38.98	24.56	36.46
D20	1250.37	Yao II–III	I	0.44	0.44	0.57	0.75	25.67	32.41	41.93
D209	1272	Yao II–III	I	0.44	0.45	0.66	1.10	27.36	31.82	40.82
D43	1094.6	Yao II–III	I	0.45	0.43	0.66	1.06	24.02	33.4	42.58
D46	1070.37	Yao II–III	I	0.45	0.43	0.60	0.61	27.09	29.33	43.58
D77	1052.5	Yao II–III	I	0.47	0.46	0.62	0.67	28.31	30.53	41.16
F722	455	Yao II–III	I	0.45	0.49	0.59	0.64	21.66	34	44.34
Jiang54	484	Yao II–III	I	0.47	0.44	0.61	0.68	27.83	30.68	41.49
L27	676	Yao II–III	I	0.43	0.46	0.62	0.83	26.07	30.69	43.25
D209	1341.5	Yao I	I	0.42	0.44	0.63	0.82	24.09	30.8	45.11
D46	1085.57	Yao I	I	0.47	0.45	0.59	0.63	24.39	30.22	45.39
Long68	1762.98	Yao I	I	0.44	0.45	0.63	0.87	26.42	33.5	40.08
T30	1039	Yao I	I	0.46	0.45	0.54	0.68	26.55	29.22	44.23
T6	1245.85	Yao I	I	0.46	0.44	0.61	0.75	24.32	34.09	41.59
D20	1404	Qing II–III	I	0.46	0.45	0.58	0.71	22.45	32.86	44.68
D209	1626	Qing II–III	I	0.45	0.44	0.66	1.10	23.27	32.72	44.01
D21	1450	Qing II–III	I	0.46	0.45	0.59	0.75	22.01	34.18	43.81
D410	1115	Qing II–III	I	0.45	0.52	0.65	0.91	22.03	33.2	44.77
D420	1106	Qing II–III	I	0.46	0.45	0.68	1.12	23.02	32.06	44.92
D610	770	Qing II–III	I	0.47	0.45	0.62	0.79	25.6	31.63	42.77
Jiang37	596	Qing II–III	I	0.47	0.44	0.61	0.67	29.97	30.35	39.68
Jiang55	466	Qing II–III	I	0.50	0.45	0.62	0.72	27.23	31.25	41.52
L27	687	Qing II–III	I	0.46	0.46	0.65	1.04	25.44	30.73	43.82
T6	1300	Qing II–III	I	0.50	0.47	0.68	1.18	26.26	33.64	40.1

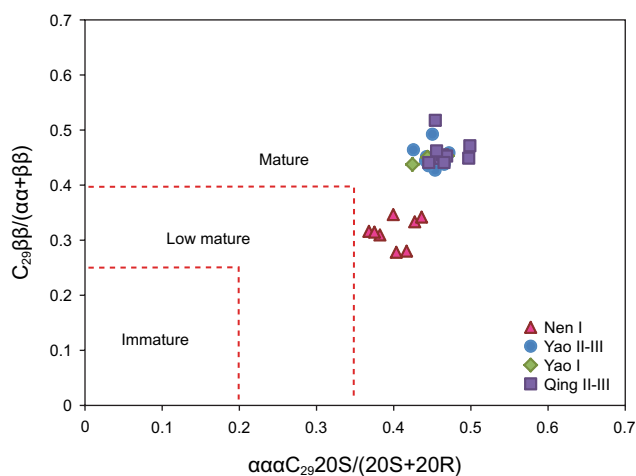


Fig. 18 Sterane isomerization index distribution of crude oils in the Western Slope of Northern Songliao Basin

Group I oils and Gulong source rock are dominated by C_{29} steranes and have similar distributions of C_{27} - C_{28} - C_{29} steranes ($C_{29} > C_{28} > C_{27}$), indicating that they are genetically related. $C_{29} \beta/(\alpha + \beta)$ and $\alpha\alpha C_{29} 20S/(20S + 20R)$ sterane ratios in group I oils are relatively high, indicating that they were derived from source rocks with maturity relatively

high. In terms of the sterane parameters, group I oils correlate with Gulong source rock extracts.

4.3.2 Group II oils

Oils from Nen I member, defined as group II, and extracts of source rock from Qijia generally show similar biomarker fingerprints. Group II oils and Qijia source rock extracts have low abundance in tricyclic terpanes and relative equilibrium between T_s and T_m hopanes. In addition, group II oils have low $T_s/(T_s + T_m)$ and $C_{29} T_s/C_{29}$ hopane ratios, suggesting the maturity are relatively low.

Group II oils are dominated by C_{27} steranes and correlated with Qijia source rock extracts which have similar distributions of C_{27} - C_{28} - C_{29} steranes ($C_{27} > C_{29} > C_{28}$). Group II oils have low values in $C_{29} \beta/(\alpha + \beta)$ and $\alpha\alpha C_{29} 20S/(20S + 20R)$ sterane ratios, suggesting that they were derived from source rocks with maturity relatively low. All of these results demonstrate that oils in the Nen I member were mainly derived from Qing I source rock within Qijia sag.

Therefore, based on the geological background, the distribution characteristics of biomarkers and the maturity differences between crude oils and source rocks, it can be concluded that the crude oils in the Qing II–III, Yao I and Yao II–III members were derived from Gulong source rock

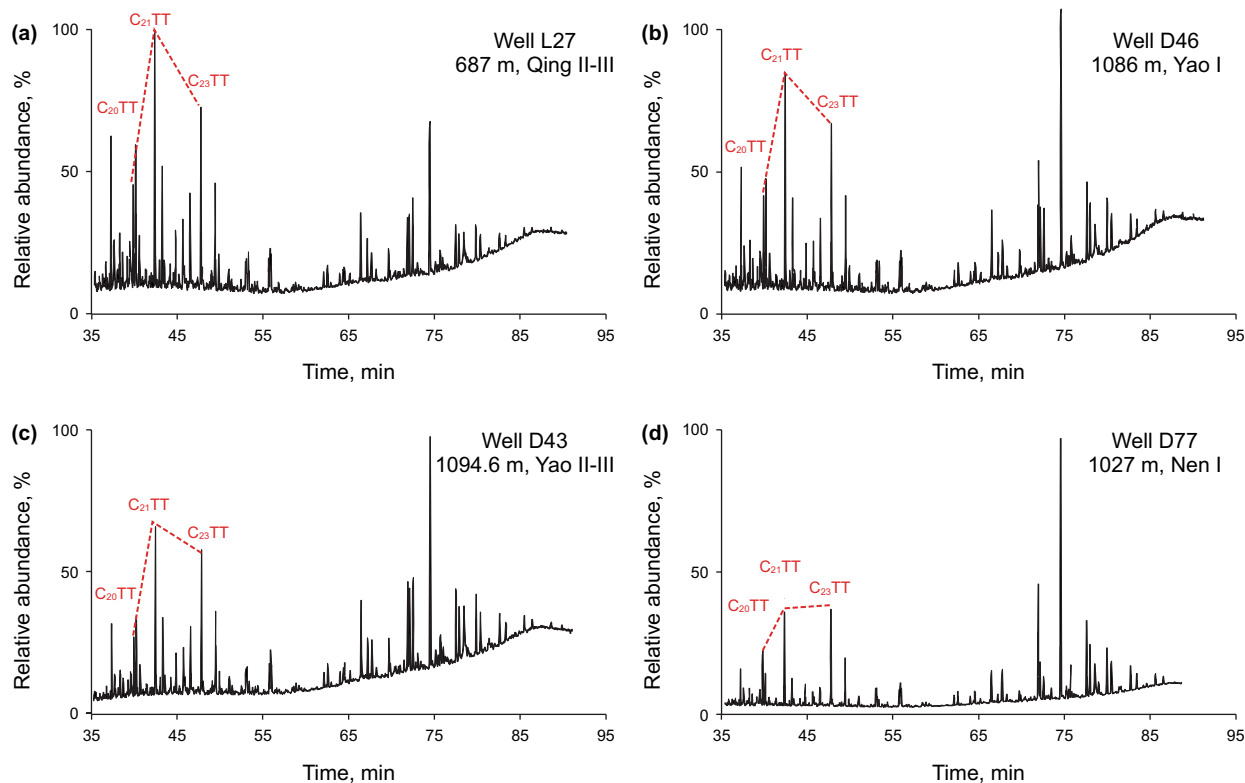


Fig. 19 Terpene distributions (m/z 191) in the oils from different members in the Western Slope

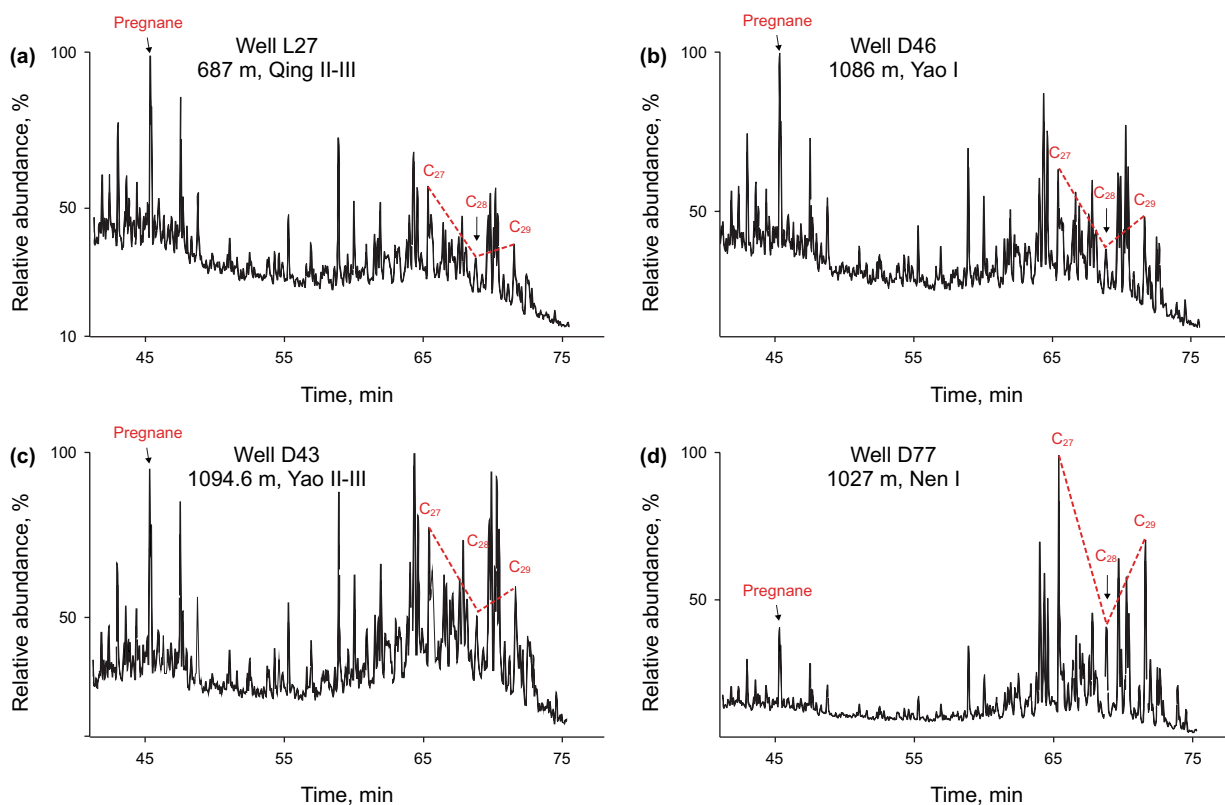


Fig. 20 Sterane distributions (m/z 217) in the oils from different members in the Western Slope

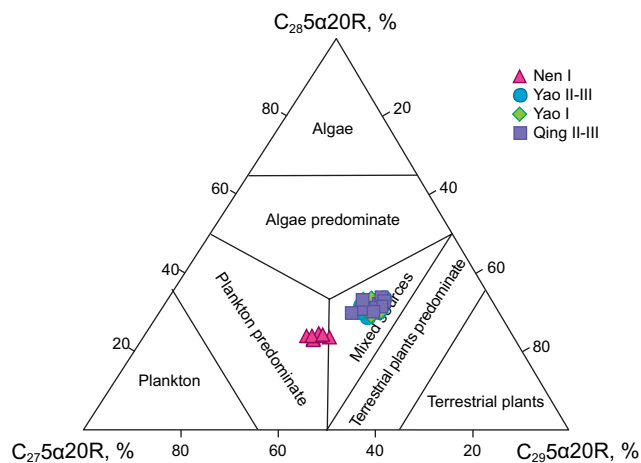


Fig. 21 Ternary diagram showing the relative abundance of C₂₇, C₂₈ and C₂₉ regular steranes in saturate hydrocarbon fraction of crude oils

and that the crude oils in the Nen I member were derived from Qijia source rocks.

The differences in crude oil sources in the Western Slope are mainly controlled by the differences in source rock development, hydrocarbon expulsion and distribution of migration paths. The differences in organic matter provenance, thermal maturity, generation and expulsion evolution

between the Gulong and Qijia source rocks controlled the differences in hydrocarbon generation and expulsion and the important contribution of the Gulong source rocks. The late expulsion of hydrocarbons in the Qijia sag and the existence of surrounding faults formed in the late stage of the deposition of the Mingshui Formation may have led to hydrocarbon migration and accumulation in the sandstone of the Nen I member, resulting in difference and delamination between hydrocarbon sources.

5 Conclusions

Multiple geochemical parameters were employed for source rock characteristic analysis, such as organic matter abundance, type, provenance, thermal maturity and saturated hydrocarbon chromatography, to recognize the differences between the Gulong source rocks and the Qijia source rocks. Furthermore, multiple analytical parameters were employed for oil–source correlation, such as molecular characteristics and biomarker fingerprints of source rock extracts and crude oils, to differentiate among the sources. The results and analyses allow the following conclusions:

- (1) The possible source rocks are mainly distributed in the Qing I and Nen I members of Gulong and Qijia sags with a high abundance of organic matter. The organic matter of Gulong source rocks is primarily Type II–III kerogen and derived from mixed sources, while that of Qijia source rocks is mixed Type I–II kerogen and primarily derived from plankton. Gulong source rocks feature a high degree of thermal evolution and a broad range of mature source rocks. The Qing I member in the Gulong sag has the highest thermal maturity. The hydrocarbon evolution profiles indicate that the Qing I source rocks have been the most effective at producing and expelling hydrocarbons, especially in the Gulong sag, where the rocks have a high conversion rate.
- (2) Based on physical properties and molecular compositions, crude oils in the target strata are generally classified into two groups. Group I oils exist in the Qing II–III, Yao I, and Yao II–III members with maturity relatively high, while group II oils originate from Nen I member with maturity relatively low.
- (3) Great differences in biomarker characteristics exist between the Gulong source rocks and the Qijia source rocks, as primarily evidenced by differences in the distribution characteristics of tricyclic terpanes, hopanes, pregnane, and C_{27} , C_{28} , C_{29} regular steranes. Correlations of these biomarker fingerprints between the source rocks and the crude oils suggest that the group I oils from the Qing II–III, Yao I, and Yao II–III members are mainly derived from Gulong source rocks, while group II oils from the Nen I member are derived from Qijia source rocks.

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